



## Article

# Prediction of Future Spatial and Temporal Evolution Trends of Reference Evapotranspiration in the Yellow River Basin, China

Shengqi Jian <sup>1</sup>, Aoxue Wang <sup>1</sup>, Chengguo Su <sup>1,\*</sup> and Kun Wang <sup>2</sup><sup>1</sup> Yellow River Laboratory, Zhengzhou University, Zhengzhou 450001, China<sup>2</sup> China Institute of Water Resources and Hydropower Research, Beijing 100038, China

\* Correspondence: cgsu@zzu.edu.cn

**Abstract:** Reference evapotranspiration ( $ET_0$ ) is an integral part of the regional hydrological cycle and energy balance and is extremely sensitive to climate change. Based on temperature data from 24 global climate models (GCMs) in the Coupled Model Intercomparison Project Phase 6 (CMIP6), this study developed a multi-model ensemble based on delta statistical downscaling with multiple interpolation methods and evaluation indicators to predict the spatial and temporal evolution trends of  $ET_0$  in the Yellow River Basin (YRB) under four emission scenarios (SSP126, SSP245, SSP370, and SSP585) for the near- (2022–2040), mid- (2041–2060), and long- (2081–2100) term future. Results demonstrate that regional data generated based on delta statistical downscaling had good simulation performance for the monthly mean, maximum, and minimum temperatures in the YRB, and the developed multi-model ensemble had better simulation capability than any single model. Compared to the historical period (1901–2014), the annual  $ET_0$  showed a highly significant increase for different future emission scenarios, and the increase is faster with increasing radiative forcing. The first main cycle of  $ET_0$  change was 52, 53, 60, and 48 years for the SSP126, SSP245, SSP370, and SSP585, respectively.  $ET_0$  in the YRB had positive values for EOF1 under all four emission scenarios, responding to a spatially consistent trend across the region. Compared to the historical period, the spatial distribution of  $ET_0$  under different future emission scenarios was characterized by being larger in the west and smaller in the east. As the radiative forcing scenario increased and time extended,  $ET_0$  significantly increased, with a maximum variation of 112.91% occurring in the western part of the YRB in the long-term future under the SSP585 scenario. This study can provide insight into the water cycle patterns of watersheds and scientific decision support for relevant departments to address the challenges of climate change.

**Keywords:** reference evapotranspiration; CMIP6; delta statistical downscaling; Hargreaves model; Yellow River Basin; EOF analysis



**Citation:** Jian, S.; Wang, A.; Su, C.; Wang, K. Prediction of Future Spatial and Temporal Evolution Trends of Reference Evapotranspiration in the Yellow River Basin, China. *Remote Sens.* **2022**, *14*, 5674. <https://doi.org/10.3390/rs14225674>

Academic Editors: Alban Kuriqi and Luis Garrote

Received: 15 October 2022

Accepted: 7 November 2022

Published: 10 November 2022

**Publisher's Note:** MDPI stays neutral with regard to jurisdictional claims in published maps and institutional affiliations.



**Copyright:** © 2022 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (<https://creativecommons.org/licenses/by/4.0/>).

## 1. Introduction

Climate change and its impacts on the water cycle, particularly on regional hydrological systems, are major global challenges in the 21st century [1–3]. As an important factor in the regional hydrological cycle and energy balance, reference evapotranspiration ( $ET_0$ ) can be used to make total energy estimates of actual evapotranspiration [4], and is the component of the water cycle that is directly affected by climate change. Changes in  $ET_0$  have a significant impact on the global water cycle and water resources [5], thereby leading to droughts and floods, water scarcity, and ecosystem degradation. In the context of climate change,  $ET_0$  is an important guide for understanding the hydrological cycle and formulating water resource plans in watersheds [6–8].

Although studies on  $ET_0$  have been conducted recently [9–14], most existing studies focused on the historical period. With the development of global climate models (GCMs), exploring the future  $ET_0$  of watersheds based on historical data has become a topic of research interest in the context of climate change. GCMs are the most powerful tools for

climate change modeling and future predictions [15,16], and the modeling results can provide valuable data to support studies on climate change-induced impacts at regional and continental scales. Nevertheless, low-resolution data will lead to large biases in the prediction of regional climate change when climate studies are conducted at regional scales. Downscaling is an effective method for transforming large-scale, low-resolution outputs from GCMs into small-scale, high-resolution regional ground information [17,18]. Current mainstream downscaling methods include dynamic downscaling [19] and statistical downscaling [20]. Compared to statistical downscaling, dynamic downscaling requires a large number of complex inputs and computational requirements [21,22], and sometimes fine and reliable climate data at regional scales are not available [22,23]. Statistical downscaling is the most widely used and established downscaling technique in basin climate change studies because of its low computational cost, easy model construction, multiple implementation methods, ease of operation, and lack of consideration of the influence of boundary conditions on prediction results [24].

Two methods are widely used for  $ET_0$  prediction under future climate scenarios: (1) input of future meteorological data from GCMs into  $ET_0$  models [25,26]; and (2) directly predicting future  $ET_0$  via downscaling methods based on historical  $ET_0$  [1,27]. Liu et al. [28] used Coupled Model Intercomparison Project Phase 5 (CMIP5) and Coupled Model Intercomparison Project Phase 6 (CMIP6) climate models to compare global potential evapotranspiration and found that both models could effectively simulate the increasing trend; they also revealed that CMIP6 multi-model results simulated higher values of global potential evapotranspiration than CMIP5 for the same emission scenario. Nistor et al. [29] assessed the impact of climate change on  $ET_0$  in Turkey in the 21st century based on the Thornthwaite equation and the CMIP5 dataset. They revealed that  $ET_0$  will increase in southern and southeastern Turkey and along the Mediterranean coast in the coming period owing to climate warming.

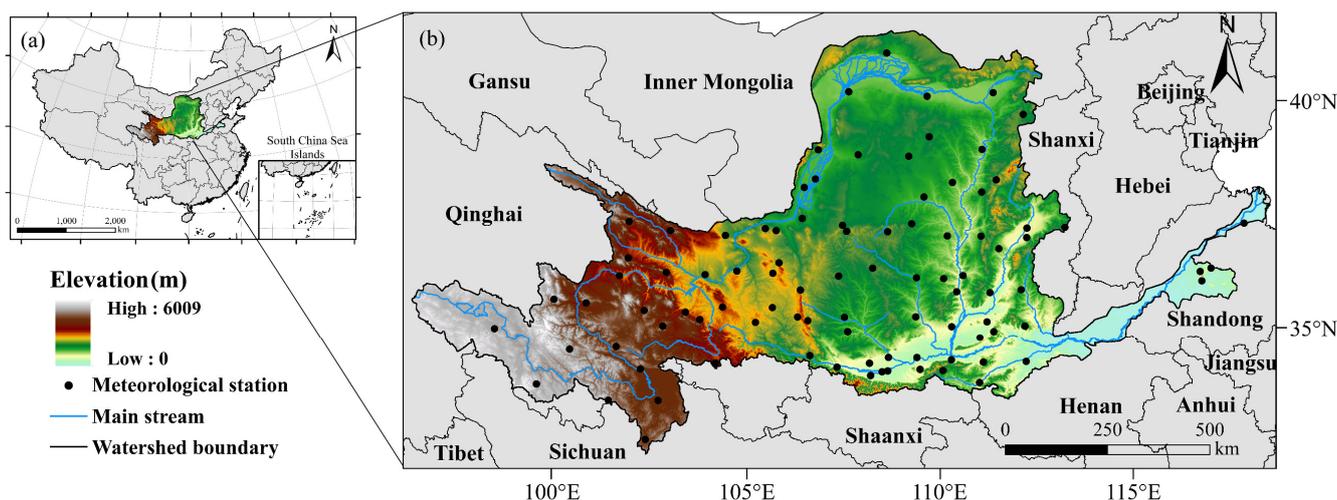
The Yellow River Basin (YRB) is an important component of China's strategic ecological security pattern, and most of the YRB is an arid and semi-arid region. Because of its unique geographical location, its environment is fragile and highly sensitive to global climate change [21,30], making it a good indicator of climate change. Despite the high sensitivity of the region to climate change, studies on the evolution of  $ET_0$  in the YRB in the context of future climate are limited, and most of the existing studies on future  $ET_0$  in other regions are at the CMIP5 stage [21,30,31], with no downscaling treatment [1], a single spatial interpolation method [21], or a single indicator for climate model preferences [32]. Therefore, against the backdrop of global warming, the  $ET_0$  predictions in the YRB can provide a theoretical reference basis for water resource planning and management, as well as a scientific basis for relevant authorities to formulate future climate change response strategies.

This study used the YRB as the study area and developed a multi-model ensemble based on the delta statistical downscaling using multiple interpolation methods and multiple evaluation indicators to predict the spatial and temporal evolution characteristics of  $ET_0$  in the YRB under different CMIP6 emission scenarios. Studies on  $ET_0$  not only enhance the understanding of hydrological processes in the YRB but also provide data to support and guide future water resource management and drought mitigation. The specific objectives are to: (1) obtain monthly mean, maximum, and minimum temperature datasets in the YRB with a resolution of  $1 \times 1$  km based on CMIP6 climate model data and delta statistical downscaling; (2) select the best simulated climate model and multi-model ensemble by evaluating and validating historical measured data; and (3) predict the spatial and temporal changes in  $ET_0$  under different emission scenarios in the future based on the Hargreaves formula and downscaled temperature data from 2022 to 2100.

## 2. Materials and Methods

### 2.1. Study Area

The Yellow River, the second-largest river in China, starts from the Bayankara Mountains in the west, crosses the Qinghai–Tibetan Plateau, Loess Plateau, and Huang-Huai-Hai Plain, and flows into the Bohai Sea in Shandong Province. The total length of the basin is 5464 km, covering an area of approximately  $79.5 \times 10^4 \text{ km}^2$  ( $95^\circ 53' \text{E}$ – $119^\circ 05' \text{E}$ ,  $32^\circ 10' \text{N}$ – $41^\circ 50' \text{N}$ ), accounting for 8% of China's land area (Figure 1). The YRB is in the mid-latitude zone, with complex natural conditions and undulating basin topography, and is influenced by atmospheric and monsoonal circulation, making the climate different from that of the other basins in China [33,34]. The average annual precipitation in the basin is 495 mm, with concentrated and highly variable interannual precipitation and an evident downward trend from the southeast to the northwest [35]. The average annual temperature ranges from  $-4$  to  $14$  °C, varying with latitude and altitude [36]. The basin's evapotranspiration varies markedly, with an average annual  $ET_0$  of 700–1800 mm and an increasing trend from the southeast to the northwest. As the basin straddles arid, semi-arid, and semi-humid zones, it is in the transition zone between semi-arid and semi-humid climates, rendering it extremely sensitive to climate change [21]. Climate change has exacerbated the uneven spatial and temporal distributions of water resources in the YRB, and the contradiction between water resource supply and demand has become evident, seriously affecting the production and life of human society and restricting the high-quality economic development of the region.



**Figure 1.** (a) Location and digital elevation model of the Yellow River Basin in China and (b) the distribution of 93 national meteorological stations in the Yellow River Basin.

### 2.2. Data Collection

#### 2.2.1. Ground-Based Observation Data

In this study, monthly monitoring data from 93 national meteorological stations in the YRB from 1980 to 2014 were obtained from the National Meteorological Information Centre–China Meteorological Data Network (<http://data.cma.cn/> (accessed on 11 March 2022)), including monthly mean temperature ( $tas$ ), monthly mean maximum temperature ( $tasmax$ ), monthly mean minimum temperature ( $tasmin$ ), and monthly pan evaporation. Some of the missing data were reasonably interpolated via the hydrologic analogy method and the linear interpolation method. The  $tas$ ,  $tasmax$ , and  $tasmin$  were used to assess the accuracy of the climate model simulations, and the converted value based on pan evaporation data [4] were used to assess the  $ET_0$  values based on the multi-model ensemble and Hargreaves formula.

### 2.2.2. Reference Data on Downscaling

The national 30-year cumulative mean, mean maximum, and mean minimum temperature datasets with a resolution of 1 km from 1971 to 2000 were selected as the regional high-resolution reference data to construct the delta statistical downscaling model in this study. Data were obtained from the National Ecosystem Science Data Center (NESDC) (<http://www.nesdc.org.cn/> (accessed on 9 May 2022)).

### 2.2.3. Future Climate Data

In this study, 24 GCMs were selected from CMIP6 (<https://esgf-node.llnl.gov/search/cmip6/> (accessed on 13 May 2022)) for the historical period (1901–2014) and three future periods (near-term 2022–2040, mid-term 2041–2060, and long-term 2081–2100). The tas, tasmax, and tasmin data of the models were presented, which contained 21, 19, and 21 GCMs, respectively; the basic details about each model and variable are summarized in Table 1. For future forcing scenarios, the recent shared socioeconomic pathways (SSPs), such as SSP1-2.6 (low-forcing scenario, SSP126), SSP2-4.5 (medium-forcing scenario, SSP245), SSP3-7.0 (medium to high-forcing scenario, SSP370), and SSP5-8.5 (high-forcing scenario, SSP585), were selected [37]. Notably, the future scenarios of the climate model were set for the 2015–2100 period; the historical period in this study did not extend back to 2021, and the future period did not extend forward to 2015 to ensure the reasonability of the data. The selection of the periods for downscaling the simulation accuracy and Hargreaves model validation were based on these considerations.

**Table 1.** Introduction to climate models with temperature variables.

Serial Number	Climate Models	Variables	Research Institution, Country	Spatial Resolution
1	ACCESS-CM2	tasmax, tasmin	ACCESS, Australia	1.9° × 1.3°
2	ACCESS-ESM1-5	tas, tasmax, tasmin	ACCESS, Australia	1.9° × 1.3°
3	AWI-CM-1-1-MR	tas, tasmin	AWI, Germany	0.9° × 0.9°
4	AWI-ESM-1-1-LR	tasmax, tasmin	AWI, Germany	1.9° × 1.9°
5	BCC-CSM2-MR	tas, tasmax, tasmin	BBC, CMA, China	1.125° × 1.125°
6	BCC-ESM1	tasmax, tasmin	BBC, CMA, China	2.8° × 2.8°
7	CanESM5	tas, tasmax, tasmin	CCCMA, Canada	2.8125° × 2.8125°
8	CMCC-CM2-SR5	tas	CMCC, Italy	1.250° × 0.938°
9	CMCC-ESM2	tas, tasmin	CMCC, Italy	1.25° × 0.9375°
10	E3SM-1-0	tas	LLNL, ANL, LANL, LBNL, ORNL, PNNL, SNL, U.S.A	1° × 1°
11	EC-Earth3	tas, tasmax, tasmin	EC-Earth, 10 European countries	0.7° × 0.7°
12	EC-Earth3-Veg	tas, tasmax, tasmin	EC-Earth, 10 European countries	0.703° × 0.703°
13	FGOALS-f3-L	tas	IAP, CAS, China	1° × 1.25°
14	FIO-ESM-2-0	tas, tasmax, tasmin	FIO, China	0.9424° × 1.25°
15	GFDL-ESM4	tas, tasmax, tasmin	GFDL, U.S.A	1° × 1.25°
16	GISS-E2-1-G	tas, tasmax, tasmin	NASA-GISS, U.S.A	1° × 1.25°
17	INM-CM5-0	tas, tasmax, tasmin	INM, Russia	2° × 1.5°
18	IPSL-CM6A-LR	tas, tasmax, tasmin	IPSL, France	1.2676° × 2.5°
19	MIROC6	tas, tasmax, tasmin	MIROC, Japan	1.389° × 1.406°
20	MPI-ESM-1-2-HAM	tas, tasmax, tasmin	MPI, Germany	1.865° × 1.875°
21	MPI-ESM1-2-HR	tas, tasmax, tasmin	MPI, Germany	0.9375° × 0.9375°
22	MPI-ESM1-2-LR	tas, tasmax, tasmin	MPI, Germany	1.875° × 1.875°
23	MRI-ESM2-0	tas, tasmax, tasmin	MRI, Japan	1.124° × 1.125°
24	NESM3	tas, tasmax, tasmin	NUIST, China	1.865° × 1.875°

Note: In the variable column, tas is the average temperature, tasmax is the average maximum temperature, and tasmin is the average minimum temperature.

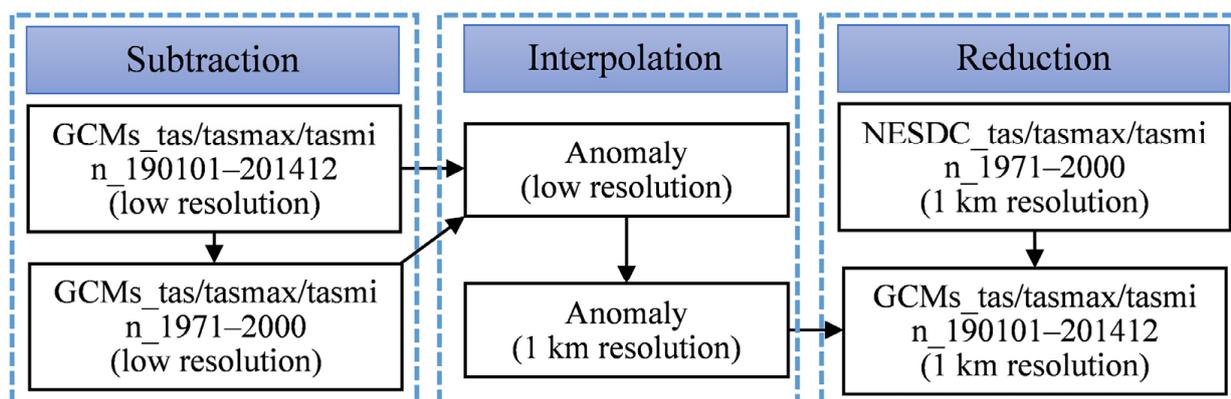
### 2.3. Research Methodology

#### 2.3.1. Delta Statistical Downscaling

The delta statistical downscaling method is a simple bias correction technique recommended by the U.S. Global Change Research Program (see <http://www.nacc.usgcrp.gov> (accessed on 6 June 2022)) that is easy to understand and operate, requires fewer factors, and is widely used in a wide range of fields related to climate change impact studies [21,38,39]. For the temperature variables used in this study, the delta method was used to compare the temperature of different periods of each simulation grid with the simulated average temperature of the base period, calculate the absolute change in temperature in each period of each simulation grid, and add the measured average temperature of each base period with the change in the grid based on the spatial interpolation of the change to obtain the temperature scenarios of different periods in the reconstruction grid [21]. The calculation equation is as follows:

$$T_f = T_0 + (T_{Mf} - T_{M0}) \quad (1)$$

where  $T_f$  is the grid temperature data reconstructed by the delta method,  $T_{Mf}$  is the simulated grid temperature data for a certain period,  $T_{M0}$  is the simulated grid multi-year average temperature data for the base period, and  $T_0$  is the measured multi-year average temperature data for the base period. In this study, five interpolation methods were considered: bilinear interpolation (BI), inverse distance weighted (IDW), kriging, natural neighbor interpolation (NNI), and spline. The delta statistical downscaling process is shown in Figure 2.



**Figure 2.** The delta downscaling process of the climate models over the Yellow River Basin.

#### 2.3.2. Climate Model Accuracy Assessment and Multi-Model Ensemble

To effectively assess the accuracy and applicability of climate model forecasts in the YRB, the evaluation metrics used were mean absolute error (MAE) [40], Taylor diagram-based quantile S [41], spatial skills score (SS) [42], and temporal skills score (TS) [43]. The closer the MAE and TS are to 0, the better the simulation ability of the model. The closer S and SS are to 1, the better the simulation ability of the model.

The downscaling results of different GCMs differ, and the performance of multi-model averaging is considered to be better than that of individual models [44,45]. In this study, multi-model ensemble averaging of preferred climate models was performed using the equally weighted ensemble averaging (MME) method commonly used in multi-model prediction studies.

#### 2.3.3. $ET_0$ Calculation Model

The Hargreaves model, which can reveal the physics of the evaporative process, was used for calculating monthly  $ET_0$  based on a future climate. It has been widely demonstrated to be able to provide reliable estimations [31,46,47]. The Food and Agriculture Organization of the United Nations (FAO) [4] also suggests it as the simplified standard al-

gorithm for  $ET_0$  under the missing meteorological observations as required by the Penman–Monteith formula. Several improvements were made to the original equation [48,49]. The form used in this study was published in 1985 [50] and expressed as follows:

$$ET_{0(Har)} = 0.0023R_a(T_{max} - T_{min})^{0.5}(T_{mean} + 17.8) \quad (2)$$

where  $ET_{0(Har)}$  is the  $ET_0$  calculated by the Hargreaves empirical formula (mm/d);  $T_{max}$  and  $T_{min}$  are the mean maximum and minimum temperatures for the calculation time interval ( $^{\circ}\text{C}$ ), respectively;  $T_{mean}$  is the mean temperature for the calculation time interval ( $^{\circ}\text{C}$ ); and  $R_a$  is the zenith radiation (mm/d); the applicable time scale is 30 or 10 days.

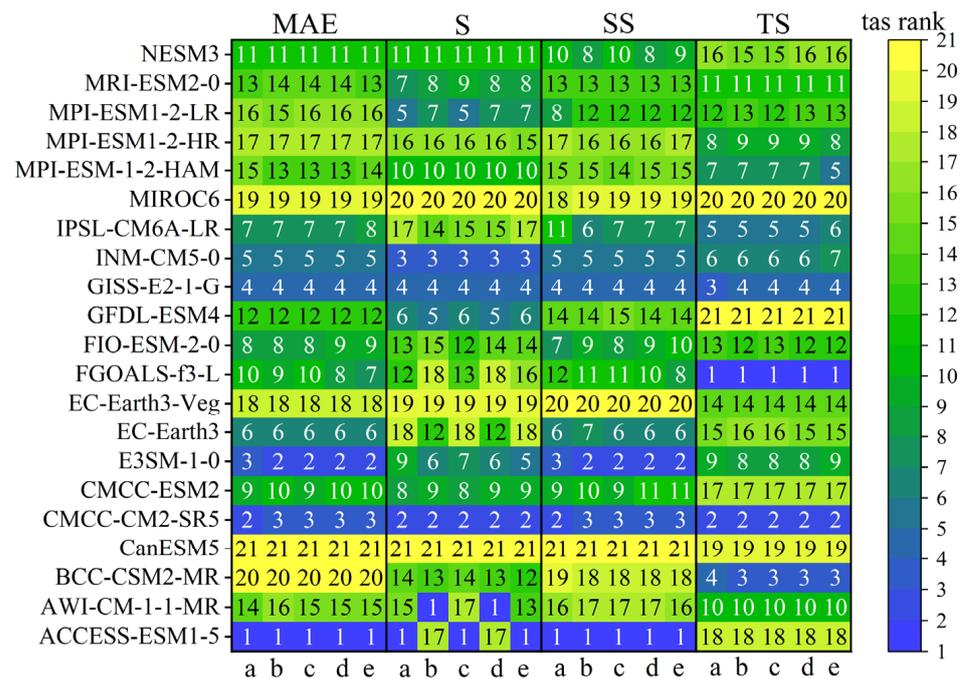
#### 2.3.4. Methods for Spatial and Temporal Trend Analysis

In this study, Morlet wavelet analysis was used to study the significant cycle changes in  $ET_0$  time series at different time scales, which has significant advantages in revealing the multiscale configuration and main climate change cycle [51]. Empirical orthogonal function (EOF) analysis produces a set of orthogonal spatial and temporal patterns in the order of explained variance, reducing the dimensionality of the analyzed system and finding relatively few independent variables that can provide comprehensive information about the variability of the raw data [52]. EOF analysis, supplemented by the North test [53], was used to study the spatial and temporal patterns of the  $ET_0$  climate field in the YRB. In addition, spatial changes in the near-, mid-, and long-term future relative to the historical period were estimated by comparing historical long-term (1901–2014) annual  $ET_0$  averages for different emission scenarios of  $ET_0$ .

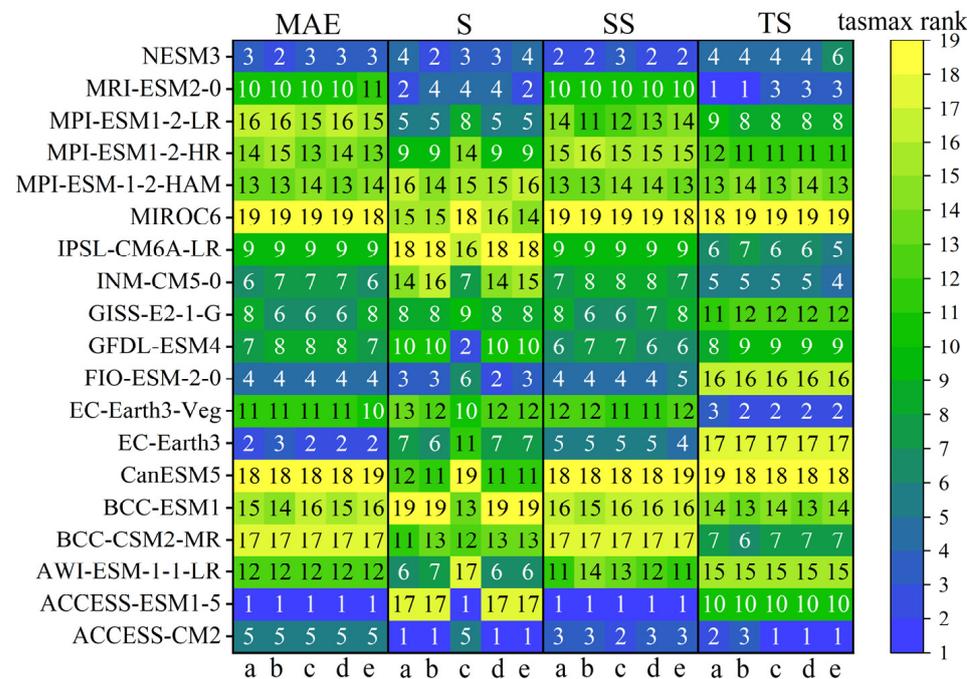
### 3. Results

#### 3.1. Simulation Accuracy Assessment of Regional Temperatures and Multi-Model Ensemble

For the three temperature variables, *tas*, *tasmax*, and *tasmin*, out of all interpolation methods, IDW exhibited the lowest overall error for all four different evaluation metrics, and therefore, it was used to further assess the effectiveness of the temperature simulations for all climate models (Figures 3–5). For *tas*, the climate models ACCESS-ESM1-5, CMCC-CM2-SR5, and GISS-E2-1-G were ranked high for each evaluation index; for *tasmax*, ACCESS-CM2, NESM3, and ACCESS-ESM1-5 were ranked high for each evaluation index; for *tasmin*, ACCESS-CM2, ACCESS-ESM1-5, and GISS-E2-1-G were ranked high for each evaluation index. Because some climate models do not have data on future emission scenarios, considering the data integrity and actual fitting performance of climate models, *tas* selected ACCESS-ESM1-5, CMCC-CM2-SR5, and INM-CM5-0; *tasmax* selected ACCESS-CM2, ACCESS-ESM1-5, and MRI-ESM2-0; and *tasmin* selected ACCESS-CM2, ACCESS-ESM1-5, and MPI-ESM1-2-LR for the ensemble analysis of subsequent temperature patterns in the YRB. The fitting results of the simulated data of the selected climate model for each variable and MME of the 1995–2014 period to the observed data from 93 ground-based meteorological stations (Figure 6) show that, although the simulated *tas*, *tasmax*, and *tasmin* of the YRB by CMIP6 were slightly lower than the observed values, most of the points were near the 1:1 line, and the  $R^2$  and regression coefficients were greater than 0.9, with all passing the 99% confidence level test. The simulated and observed values of the MME dataset were more concentrated than those of the three independent climate models, and the  $R^2$  values of the three variables were 0.9418, 0.9226, and 0.9362, respectively (Figure 6d,h,l), which reduced the errors caused by outlier points and slightly improved the fit of the simulated data. The above analysis reveals that climate models have high application potential in the YRB, and the CMIP6 multi-model ensemble is a good reference value for predicting  $ET_0$  trends in the YRB under future climate scenarios.



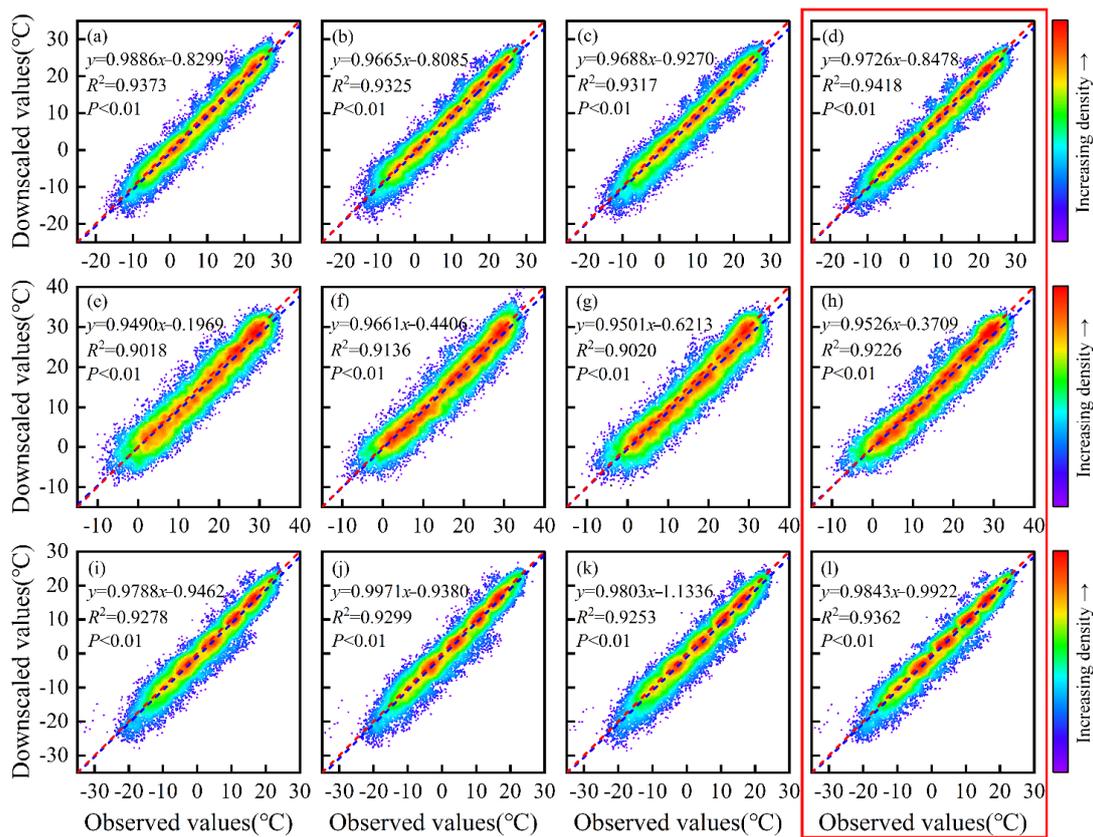
**Figure 3.** Fitting performance ranking of 21 climate models simulating tas monthly series of the Yellow River Basin from 1995 to 2014 under different interpolation methods and evaluation metrics. a, b, c, d, and e under each evaluation method are the evaluation rankings of the five interpolation methods: BI, IDW, kriging, NNI, and spline, respectively. The higher the ranking, the higher the model’s fitting accuracy.



**Figure 4.** Fitting performance ranking of 19 climate models simulating tasmax monthly series of the Yellow River Basin from 1995 to 2014 under different interpolation methods and evaluation metrics. a, b, c, d, and e under each evaluation method are the evaluation rankings of the five interpolation methods: BI, IDW, kriging, NNI, and spline, respectively. The higher the ranking, the higher the model’s fitting accuracy.

	MAE					S					SS					TS					tasmin rank
NESM3	9	10	15	11	11	9	8	10	9	8	8	7	8	7	6	15	15	15	15	15	21
MRI-ESM2-0	10	9	10	8	7	5	4	7	6	5	9	8	11	8	7	14	14	14	14	14	20
MPI-ESM1-2-LR	5	5	6	5	4	4	3	5	5	4	4	4	5	4	4	13	13	13	13	13	19
MPI-ESM1-2-HR	13	11	9	10	10	11	10	12	12	10	6	10	10	9	8	1	2	3	2	2	18
MPI-ESM1-2-HAM	7	8	8	9	9	6	6	6	8	6	10	9	13	10	9	9	8	8	8	8	17
MIROC6	16	16	16	16	16	18	16	17	18	16	16	16	16	16	15	4	6	6	6	6	16
IPSL-CM6A-LR	4	4	4	4	6	13	12	11	14	14	5	5	4	5	12	12	11	10	12	12	15
INM-CM5-0	15	14	13	15	15	10	9	8	11	12	13	14	14	14	14	8	9	9	9	9	14
GISS-E2-1-G	3	3	3	3	3	3	2	3	4	3	3	2	2	2	2	7	7	7	7	7	13
GFDL-ESM4	11	12	11	12	13	8	7	9	10	9	15	15	15	15	16	21	21	21	21	21	12
FIO-ESM-2-0	12	13	7	13	12	14	13	14	15	13	11	11	7	11	10	2	1	1	1	1	11
EC-Earth3-Veg	20	20	19	20	18	19	18	19	20	19	20	20	20	20	20	20	20	19	19	19	10
EC-Earth3	6	6	5	6	5	17	17	18	19	18	12	12	9	12	11	3	3	2	3	3	9
CMCC-ESM2	14	15	14	14	14	7	5	4	7	7	14	13	12	13	13	18	17	17	17	17	8
CanESM5	17	17	17	17	17	20	19	20	21	20	19	17	17	18	19	5	5	4	4	4	7
BCC-ESM1	18	18	18	18	20	16	15	16	17	17	17	18	18	17	17	19	19	20	20	20	6
BCC-CSM2-MR	19	19	20	19	19	15	14	15	16	15	18	19	19	19	18	10	10	11	10	10	5
AWI-ESM-1-1-LR	8	7	12	7	8	12	11	13	13	11	7	6	6	6	5	6	4	5	5	5	4
AWI-CM-1-1-MR	21	21	21	21	21	21	20	21	1	21	21	21	21	21	21	16	16	16	16	16	3
ACCESS-ESM1-5	1	1	1	1	1	1	2	1	1	1	1	1	1	1	1	17	18	18	18	18	2
ACCESS-CM2	2	2	2	2	2	2	1	2	3	2	2	3	3	3	3	11	12	12	11	11	1

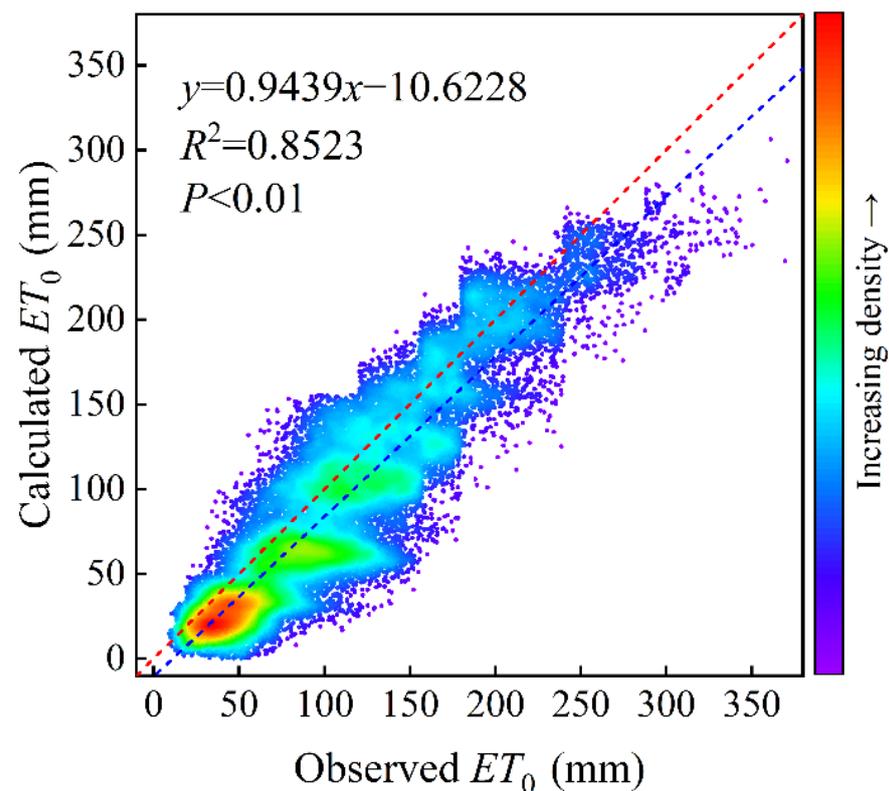
**Figure 5.** Fitting performance ranking of 21 climate models simulating the tasmin monthly series of the Yellow River Basin from 1995 to 2014 under different interpolation methods and evaluation metrics. a, b, c, d, and e under each evaluation method are the evaluation rankings of the five interpolation methods: BI, IDW, kriging, NNI, and spline, respectively. The higher the ranking, the higher the model’s fitting accuracy.



**Figure 6.** Scatter density plots of measured and selected climate models and multi-model ensembles simulating monthly temperatures in the Yellow River Basin at stations from 1995 to 2014. (a–c) correspond to the fitting effects of ACCESS-ESM1-5, CMCC-CM2-SR5, and GISS-E2-1-G, respectively, under tas variable; (d) corresponds to the fitting effect of  $MME_{tas}$ ; (e–g) correspond to the fitting effects of ACCESS-CM2, ACCESS-ESM1-5, and MRI-ESM2-0, respectively, under  $tas_{max}$  variable; (h) corresponds to the fitting effect of  $MME_{tas_{max}}$ ; (i–k) correspond to the fitting effects of ACCESS-CM2, ACCESS-ESM1-5, and MPI-ESM1-2-LR, respectively, under  $tas_{min}$  variable; and (l) corresponds to the fitting effect of  $MME_{tas_{min}}$ .

### 3.2. Simulation Accuracy Assessment of Regional $ET_0$

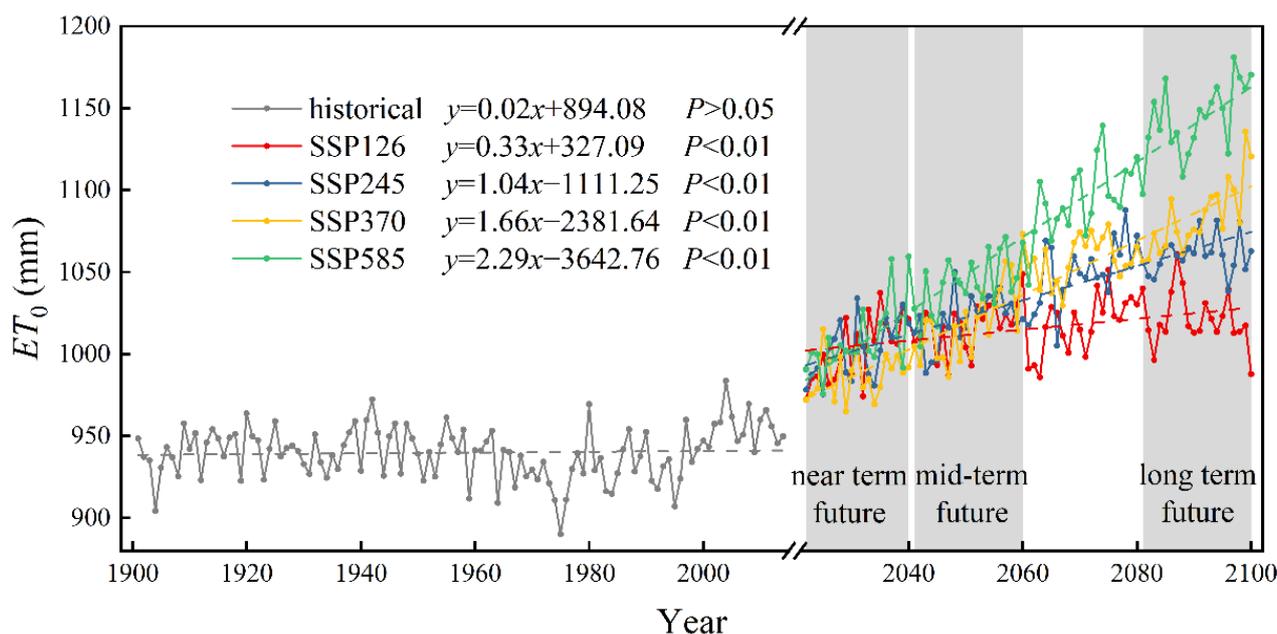
The scatter density plots of the actual monthly pan evaporation converted values ( $ET_{0(pan)}$ ) and simulated  $ET_0$  values based on the multi-model ensemble-Hargreaves formula ( $ET_{0(Har)}$ ) for the YRB from 1980 to 2014 (Figure 7) showed that the CMIP6-simulated  $ET_{0(Har)}$  for the YRB correlated well with the evaporation dish converted  $ET_{0(pan)}$  at the monthly scale, with a fitted line regression coefficient of 0.9439 and  $R^2$  of 0.8523, passing the 99% confidence level test. Although the Hargreaves formula partially underestimated  $ET_0$ , most of the points were near the 1:1 line, making them good reference values for calculating the monthly  $ET_0$  under the conditions of missing meteorological observations in the future. Therefore, future  $ET_0$  can be simulated using the Hargreaves formula.



**Figure 7.** Scatter density plot of monthly observed pan evaporation converted values ( $ET_{0(pan)}$ ) and multi-model ensemble simulations ( $ET_{0(Har)}$ ) in the Yellow River Basin from 1980 to 2014.

### 3.3. Temporal Trends and Cyclic Characteristics of $ET_0$

Based on the 1 km high-resolution regional climate data generated by the delta statistical downscaling and multi-model ensemble, annual  $ET_0$  trends were estimated for the YRB from 1901 to 2100 (Figure 8). The annual  $ET_0$  of the basin under the different future emission scenarios (2022–2100; SSP126, SSP245, SSP370, and SSP585) substantially increased with time relative to the historical period (1901–2014), and all passed the 99% confidence test. Among all scenarios, the SSP585 scenario had the most pronounced upward trend in  $ET_0$ , increasing at a rate of 22.9 mm/10a, reaching 1170.39 mm in 2100. It was followed by the SSP370 scenario, which increased at a rate of 16.6 mm/10a, reaching 1120.42 mm in 2100. However, SSP245 and SSP126 scenarios had relatively small trends in  $ET_0$ , increasing at rates of 10.4 and 3.3 mm/10a, reaching 1062.71 and 987.68 mm in 2100, respectively. In general, the annual  $ET_0$  in the YRB will rapidly increase with increasing levels of radiative forcing in the future.

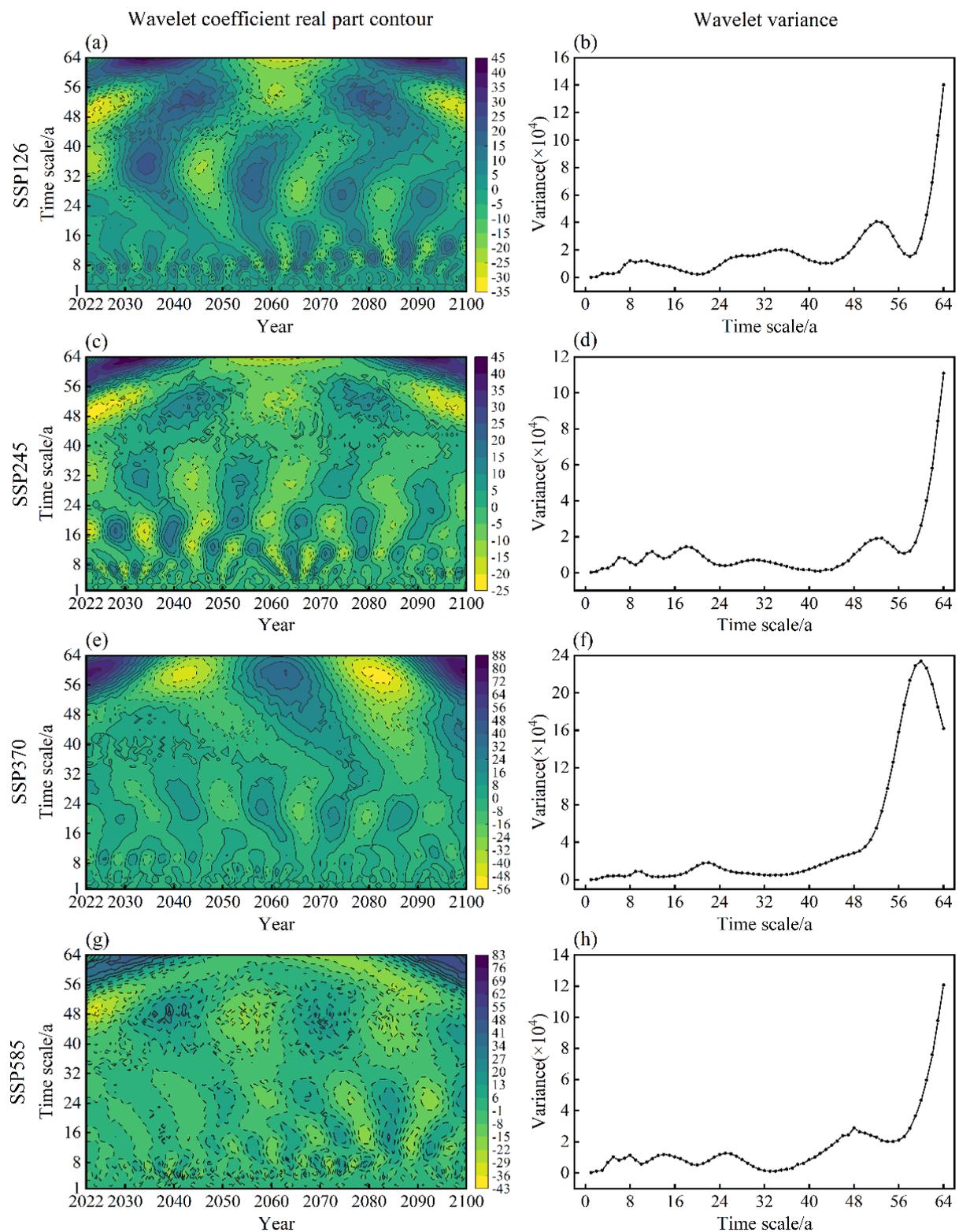


**Figure 8.** Interannual variations in  $ET_0$  in the Yellow River Basin over the historical period (1901–2014) and under different future emission scenarios (2022–2100; SSP126, SSP245, SSP370, and SSP585) ( $p$  is an indicator of significance obtained by the F-test).

To comprehensively understand the temporal  $ET_0$  variations in the basin, Morlet wavelet analysis was used to investigate the cyclic variations in  $ET_0$  in the YRB under the four future emission scenarios (Figure 9). In the SSP126 scenario, three distinct peaks were observed in the wavelet variance of  $ET_0$  in the YRB, corresponding to the time scales of 52, 35, and 8 years, with the most pronounced periodic oscillation of approximately 52 years and a cyclic pattern of 34–38 years for  $ET_0$  in the basin under this time scale (Figure 9a,b). In the SSP245 scenario, the  $ET_0$  in the YRB considerably oscillated for approximately 53 years, and there was a cyclic pattern of approximately 34 years in the basin  $ET_0$  at this time scale (Figure 9c,d). In the SSP370 scenario, the  $ET_0$  in the YRB considerably oscillated for approximately 60 years, and there was a cyclic pattern of approximately 39 years in the basin  $ET_0$  at this time scale (Figure 9e,f). In the SSP585 scenario, the  $ET_0$  in the YRB substantially oscillated for approximately 48 years, and there was a cyclic pattern of approximately 27–32 years in the basin  $ET_0$  at this time scale (Figure 9g,h).

### 3.4. Spatial Evolution Characteristics of $ET_0$

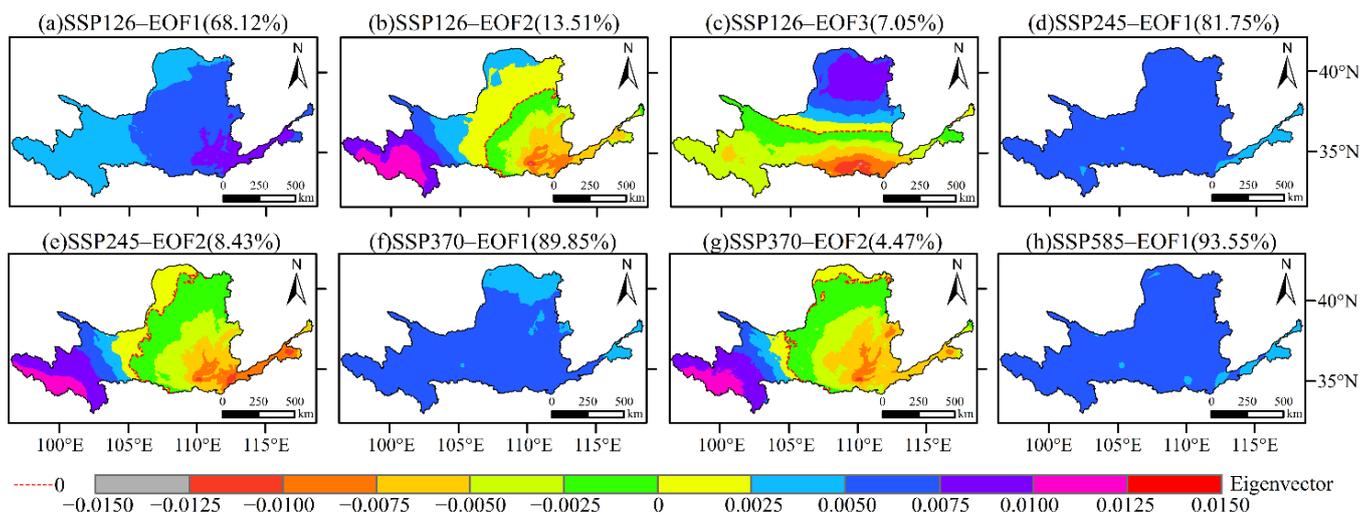
The EOF analysis of  $ET_0$  was conducted under four future emission scenarios in the YRB for 2022–2100, and the results were tested for modal significance using the North test, which are presented in Table 2 and Figures 10 and 11. As presented in Table 2, the first three modes of  $ET_0$  in SSP126 passed the North test with a cumulative variance contribution of 88.68%; the first two modes of  $ET_0$  in SSP245 passed the North test with a cumulative variance contribution of 90.18%; the first two modes of  $ET_0$  in SSP370 passed the North test with a cumulative variance contribution of 94.32%; and the first mode of  $ET_0$  in SSP585 passed the North test with a cumulative variance contribution of 93.55%.



**Figure 9.** Wavelet analysis of  $ET_0$  in the Yellow River Basin under four future emission scenarios (SSP126, SSP245, SSP370, and SSP585), including wavelet coefficient contour plots of the real part (a,c,e,g) and wavelet variance plots (b,d,f,h).

**Table 2.** Main modes and contributions of  $ET_0$  under future climate scenarios in EOF analysis.

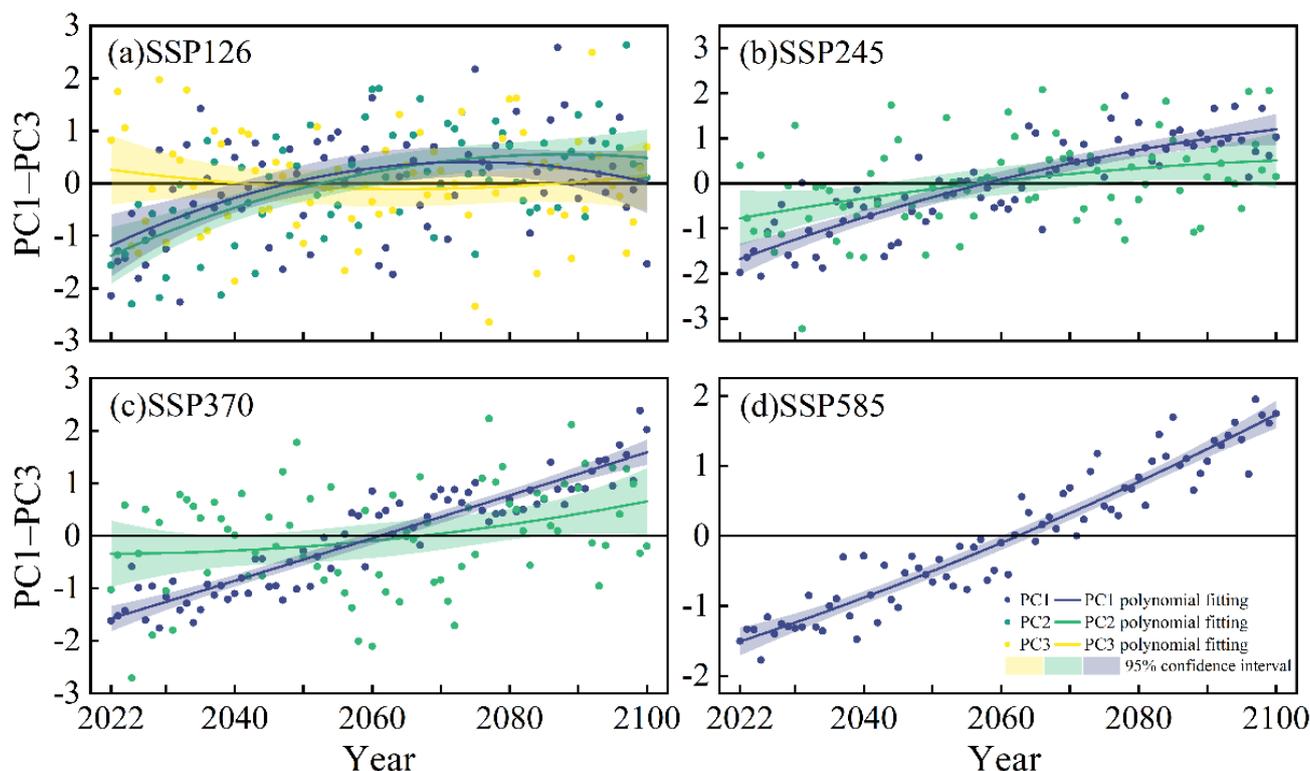
Climate Scenarios	Corresponding Modes	Variance Contribution	Cumulative Variance Contribution	North Test
SSP126	EOF1	68.12%	68.12%	pass
	EOF2	13.51%	81.63%	pass
	EOF3	7.05%	88.68%	pass
SSP245	EOF1	81.75%	81.75%	pass
	EOF2	8.43%	90.18%	pass
SSP370	EOF1	89.85%	89.85%	pass
	EOF2	4.47%	94.32%	pass
SSP585	EOF1	93.55%	93.55%	pass



**Figure 10.** Main modal eigenvectors of the EOF of the  $ET_0$  spatial field in the Yellow River Basin for 2022–2100 under scenarios SSP126 (a–c), SSP245 (d,e), SSP370 (f,g), and SSP585 (h), where SSP126/245/370/585–EOF $_i$  (contribution/%) in each figure represents the  $i$ -th modal eigenvector (EOF $_i$ ) of  $ET_0$  under scenarios SSP126/245/370/585, with the contribution of each mode in parentheses. The red dashed lines in (b,c,e,g) are the positive and negative dividing lines of the eigenvectors.

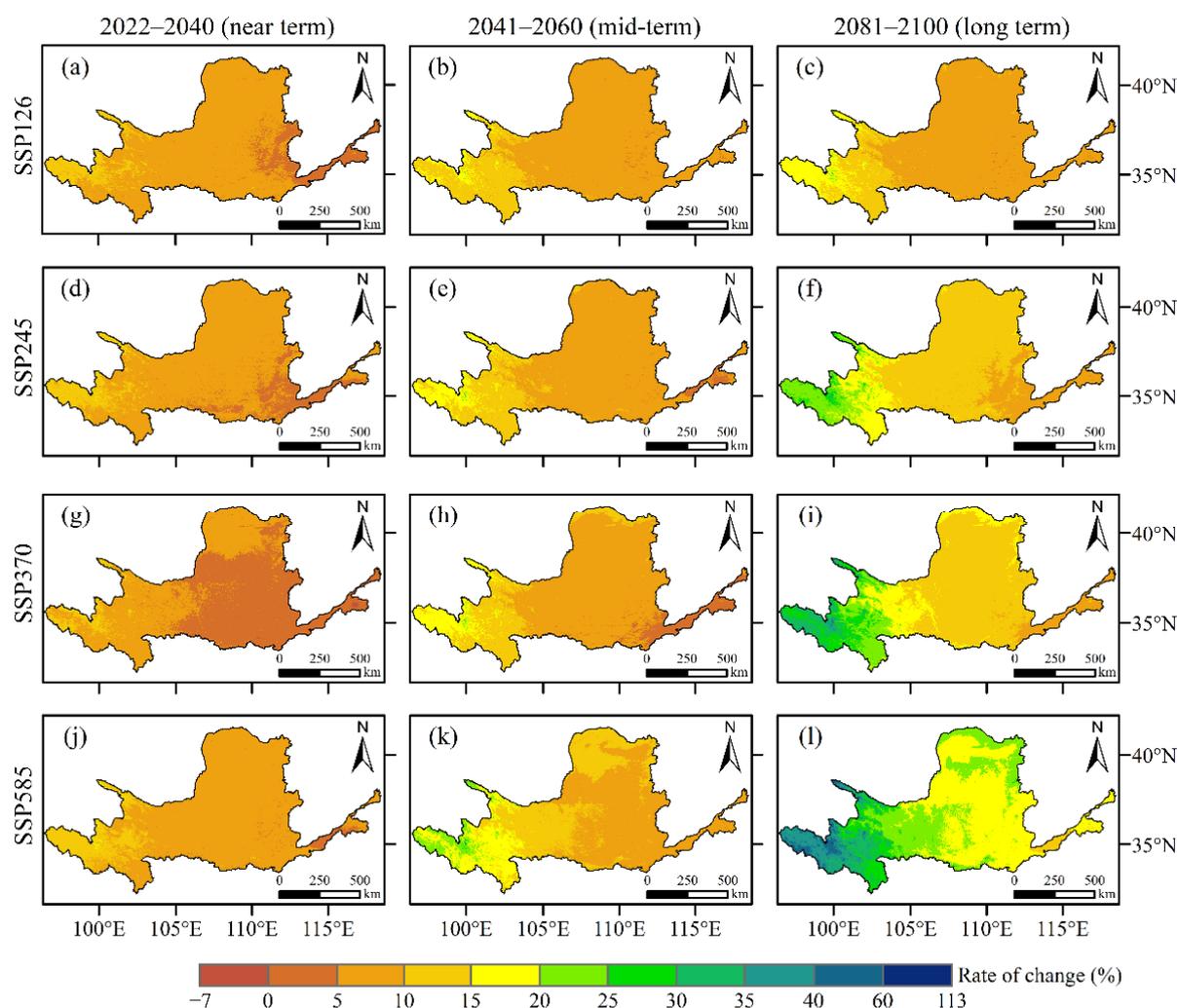
Under the SSP126 scenario, the first EOF modal eigenvector (EOF1) of  $ET_0$  in the YRB was positive, reflecting a spatially consistent trend of  $ET_0$  across the region, and exhibited an increasing trend from the northwest to the southeast, indicating a more pronounced increase in  $ET_0$  in the lower YRB (Figure 10a). EOF2 and EOF3, which explained 20.56% of the variations, reflected the secondary spatial characteristics of  $ET_0$  with opposite trends from northwest to southeast and from north to south (Figure 10b,c). Combined with the temporal coefficients (Figure 11a), PC1 and PC2 exhibited roughly the same trend, with an increasing trend from 2022 to 2100, particularly after the 2150s when PC1 and PC2 remained positive, indicating that  $ET_0$  remained high throughout this period. PC3 fluctuated at a value of approximately 0, reflecting no significant trend in  $ET_0$ . Under SSP245 and SSP370, the EOF1 eigenvectors of  $ET_0$  in the basin were all positive, exhibiting spatial trends of larger values in the upper and middle reaches and smaller values in the lower reaches, as well as larger values in the central and western parts and smaller values in the northern and eastern parts (Figure 10d,f). All the EOF2 eigenvectors exhibited a secondary spatial trend of positive in the northwest and negative in the rest of the basin, with a relatively larger increase in  $ET_0$  near the source area in the upper part of the basin and a relatively larger decrease in  $ET_0$  in the south. Combined with the temporal coefficients (Figure 11b,c), PC1 and PC2 exhibited an increasing trend from 2022 to 2100 under the SSP245 and SSP370 scenarios, and PC1 increased more than PC2, particularly after the 2060s, when PC1 and PC2 always maintained positive values, indicating that  $ET_0$  remained high during this period. The distribution of EOF1 eigenvectors for  $ET_0$  in the basin under the SSP585

scenario was similar to that of EOF1 under the SSP245 scenario (Figure 10h), with an increasing trend in the time coefficient PC1 (Figure 11d), indicating an increasing trend in  $ET_0$  in the basin and a significant increase in  $ET_0$  after the 2060s.



**Figure 11.** Principal component time coefficients (PC1–PC3) and their polynomial fits for the EOF analysis of future  $ET_0$  in the Yellow River Basin under four emission scenarios ((a) SSP126, (b) SSP245, (c) SSP370, and (d) SSP585).

As observed from the spatial variations in annual  $ET_0$  in the YRB in the near-, mid-, and long-term future relative to historical periods in the 21st century (Figure 12), the near-annual  $ET_0$  growth was generally low on an annual scale, and the rate of  $ET_0$  change was even negative in parts of Tai'an, Shandong Province, located in the lower reaches of the YRB, at  $-6.09\%$  under the SSP370 scenario. In the mid- and long-term future scenarios, the  $ET_0$  rate of change gradually increased in the whole basin, and the areas with high  $ET_0$  variations were primarily concentrated in the YRB source area and a small part of the northern basin. In the lower reaches, the  $ET_0$  change rate was low, and the variations were spatially distributed as high in the west and low in the east. As the radiative forcing increased, the increase in  $ET_0$  became more significant, ranging from  $-3.08$  to  $50.78\%$  under SSP126, from  $-1.32$  to  $68.88\%$  under SSP245, from  $-6.09$  to  $89.30\%$  under SSP370, and from  $-1.27$  to  $112.91\%$  under SSP585. A maximum variation of  $112.91\%$  was observed in the western part of the YRB in the long-term future (2081–2100) under the SSP585 scenario.



**Figure 12.** Spatial variations in the near (2022–2040; **a,d,g,j**), mid- (2041–2060; **b,e,h,k**), and long (2081–2100; **c,f,i,l**) term future annual  $ET_0$  of the Yellow River Basin relative to the historical period (1901–2014) under four SSP scenarios (SSP126, SSP245, SSP370, and SSP585).

## 4. Discussion

### 4.1. Influence of the $ET_0$ Model

The models commonly used internationally for  $ET_0$  estimation can generally be divided into four major categories [54]: temperature methods, such as the Hargreaves [50] and Blaney–Criddle models [55]; mass transfer methods, such as the Rohwer [56] and Penman models [57]; radiation methods, such as the Priestley–Taylor [58] and Jensen–Haise models [59]; and synthesis methods, such as the Penman–Monteith model [4]. Although the Penman–Monteith model is considered the standard method for calculating  $ET_0$ , accurate predictions of future  $ET_0$  in watersheds require more reliable meteorological data, and the outputs of existing GCMs and downscaling methods typically have low modeling accuracy for meteorological parameters such as wind speed, relative humidity, and radiation. Compared to other models, temperature-based models have lower data requirements and are computationally simple, and downscaling methods have a clear advantage in modeling temperature [60,61], with correlation coefficients generally above 0.90 (Figure 6), making them more widely applicable [31]. In addition, several studies have revealed strong relationships between future  $ET_0$  and temperature in other basins or regions. Xing et al. [27] attributed the increase in  $ET_0$  in the 21st century to an increase in temperature, and Ding and Peng [31] found that global warming led to a change in the main sensitive factor for potential evapotranspiration in the Loess Plateau from the average temperature in the his-

torical period to the maximum temperature in the future period. The results of both these studies theoretically supported the use of the temperature-based  $ET_0$  model. Ahmadi and Baaghdeh [25] explored the effect of climate change on  $ET_0$  in pistachio cultivation areas in Iran using the Hargreaves model. Yan and Mohammadian [47] evaluated the performance of the evaporation model based on the Hargreaves formulation using various fitting methods, and the results showed that the simulation was satisfactory. Among the many  $ET_0$  simplification methods, this study also selected the Hargreaves model, a temperature method recommended by the FAO [4] and researchers [62,63], to estimate the spatial and temporal characteristics of  $ET_0$  in the YRB under four future emission scenarios. The fitted scatter density plot had an  $R^2$  of 0.8523 at the monthly scale (Figure 7). The Hargreaves model can effectively simulate basin  $ET_0$ ; reflecting the rationality of the formula selection in this study. Notably, this study only considered the effect of temperature on  $ET_0$ , thus, there may be some bias in the estimation results. The applicability of temperature models in specific regions [64,65] should be improved, and the effect of other meteorological elements on future  $ET_0$  [66] should be explored in future studies.

#### 4.2. Spatial and Temporal Variations in Future $ET_0$

As a result of global climate change, the hydrometeorological elements and hydrological environment of the YRB have been significantly affected [67]. Although many studies have examined the trends and attribution of  $ET_0$  in the YRB [36,68], previous studies were based on site-scale and historical data. Studies on future water balance and hydrological cycles in the YRB are relatively weak owing to the lack of studies on  $ET_0$ .

This study predicted  $ET_0$  trends in the YRB for different periods under four future emission scenarios based on CMIP6 temperature data with high spatial and temporal resolution generated by the delta statistical downscaling method and the Hargreaves model. This study revealed an overall significant increase in  $ET_0$  in the YRB from 2022 to 2100 ( $p < 0.01$ ) (Figure 8), without the “evaporation paradox” [69], which is similar to the future  $ET_0$  trends predicted by many studies [28,69,70]. Based on the Hargreaves model (i.e., Equation (2)), the change in  $ET_0$  is proportional to  $t_{as}$  and the difference between  $t_{asmax}$  and  $t_{asmin}$ . With general global warming,  $t_{as}$ ,  $t_{asmax}$ , and  $t_{asmin}$  in the YRB in the future period showed sudden increases relative to the historical period (Figure S1). Therefore, the abrupt increase of  $ET_0$  can be attributed to the abrupt increase of temperature-like variables in the future period [27]. Radiative forcing is expected to stabilize at  $2.6 \text{ W/m}^2$  by 2100 for SSP126 and  $8.5 \text{ W/m}^2$  by 2100 for SSP585, and radiation values and temperature are positively related to  $ET_0$  [4]. The positive effects of increasing climatic factors, such as temperature and radiation, on  $ET_0$  in the YRB were greater than the negative effects of other factors, and therefore, the latter trend was greater than the former as emission concentrations increased in the YRB (Figure 8). The  $ET_0$  changes in the YRB in the near-, mid-, and long-term future under different future scenarios exhibited high spatial heterogeneity (Figures 10 and 12), with a spatial distribution high in the west and low in the east, and the  $ET_0$  increase became more significant as the radiative forcing scenario increased. Consistent with the results of Ding and Peng [31], the increase in  $ET_0$  was generally greater at higher elevations than at lower elevations in the basin, with the most pronounced change in  $ET_0$  in the western part of the basin, reaching a maximum variation of 112.91% compared to that in the historical period (Figure 12). According to the  $ET_0$  equation (i.e., Equation (2)) and Figure S2, this change can be attributed to the largest temperature difference between  $t_{asmax}$  and  $t_{asmin}$  in the western part of the basin. In addition, as shown in the change in future precipitation in the YRB relative to the historical period (Figure S3), precipitation in the western part of the basin showed less growth overall and even negative growth in some phases. Warming and decreasing precipitation caused an increase in dryness in the western part of the basin, and the warm-dry trend intensified. Wang et al. [71] found that the evapotranspiration process was more sensitive to relative humidity in the western part of the basin, and a decrease in relative humidity caused an increase in evapotranspiration. Therefore,  $ET_0$  predictions based on the Hargreaves model

were greater in these areas. Water loss in the YRB is likely to accelerate in the future than in historical periods, which will threaten the food and ecological security of the region; thus, developing flexible mitigation strategies tailored to local conditions is critical to coping with climate change [72].

#### 4.3. Climate Model Uncertainty Analysis

Because of the differences in the feedback processes of different GCMs, a certain degree of uncertainty exists in their response to future greenhouse gas emissions, and the actual generalized optimal climate models and  $ET_0$  models cannot be determined [73]. Inevitable uncertainties exist in future  $ET_0$  predictions stemming from climate scenarios, climate models, and  $ET_0$  models [74], which greatly affect the confidence of the prediction results.

In response to the uncertainty of climate scenarios, this study selected the CMIP6 data, which had the largest number of participating models, the richest design of numerical experiments, and the largest amount of simulated data available than other CMIP generations for more than 20 years of the CMIP [37,75], initiated by the current Working Group on Coupled Models (WGCM). Compared to previous generations (CMIP3, CMIP5, etc.), CMIP6 uses a new scenario combining shared socio-economic pathways and typical concentration pathways to constrain multi-model predictions of key climate change indicators such as global surface temperature and ocean heat content based on historical observations, climate simulations, and climate sensitivity awareness, reducing uncertainty in predictions and providing higher resolution and reliability [76], thereby making the results more informative and time-sensitive than those based on CMIP5 for future  $ET_0$  studies, such as in Ahmadi and Baaghideh [25], Ding and Peng [31], Kundu et al. [6] and Le and Bae [77]. In response to the uncertainty of climate models, this study selected 24 GCMs with historical and future emission scenarios, which is more extensive than the studies of Li et al. [70], which only used one climate model (HadCM3) under two emission scenarios (A2 and B2), and Nooni et al. [1], which used only one climate model (CNRM-CM6). In addition, this study reduced the uncertainty of future temperature data by preferentially selecting climate models based on multiple interpolation methods, multiple evaluation indicators, and equal weight sets on a downscaling basis. The MAE was controlled within 2.5 mm, S and SS were approximately 1, and TS was approximately 0, indicating very high simulation accuracy (Table 3). Wang and Chen [24] reduced the spatial resolution of GCMs' data to 0.5° based on the delta method, and the MAE of tas was in the range of 1.6–5.7 °C. However, the MAE of tas in this study was controlled in the range of 2.2–2.6 °C and had a higher spatial resolution of 1 km.

Although this study provides a comprehensive theoretical basis for future  $ET_0$  assessments, the uncertainties in the downscaling of GCMs [78] and in the selection and accuracy of  $ET_0$  models [74] may impact the prediction results. To improve  $ET_0$  estimates in future studies, consideration should be given to the long-term goal of the United Nations Framework Convention on Climate Change (Paris Agreement) to limit the increase in global average temperature to less than 2 °C compared to the pre-industrial period and to further efforts to limit it to less than 1.5 °C [79], as well as to achieve China's 2060 carbon neutrality and global carbon neutrality. Describing and quantifying the relative importance of various uncertainty sources and the risks they pose in the assessment is important in current and future climate change impact studies and water resource assessments that should be strengthened to reduce prediction uncertainty.

**Table 3.** Evaluation of the fitting error of downscaled climate models for monthly temperatures (tas/tasmax/tasmin) in the Yellow River Basin from January 1995 to December 2014.

Models	Tas				Tasmax				Tasmin			
	MAE	S	SS	TS	MAE	S	SS	TS	MAE	S	SS	TS
ACCESS-CM2	2.196449	0.984745	0.926681	0.001749	2.675489	0.999877	0.889805	$1.94 \times 10^{-6}$	2.338917	0.997536	0.915941	0.001035
ACCESS-ESM1-5	2.442822	0.999548	0.909236	0.00021	2.497431	0.974297	0.901489	0.000464	2.295015	0.963689	0.917158	0.004465
AWI-CM-1-1-MR									2.837847	0.965667	0.881039	0.003347
AWI-ESM-1-1-LR					2.789143	0.987086	0.876886	0.00268	2.470842	0.990019	0.907696	0.000134
BCC-CSM2-MR	2.548259	0.986199	0.904698	$7.59 \times 10^{-6}$	2.886012	0.979498	0.871659	0.000229	2.65175	0.985043	0.89462	0.00095
BCC-ESM1					2.810414	0.954278	0.876325	0.00136	2.617258	0.98485	0.898064	0.005105
CanESM5	2.553748	0.965738	0.900803	0.002406	2.909749	0.980601	0.867035	0.007883	2.572498	0.970803	0.898088	0.000166
CMCC-CM2-SR5	2.273762	0.99501	0.920346	$2.77 \times 10^{-6}$								
CMCC-ESM2	2.393424	0.990247	0.912677	0.001578					2.496331	0.99373	0.905089	0.003913
E3SM-1-0	2.264877	0.990943	0.921241	0.000168								
EC-Earth3	2.345391	0.986413	0.91518	0.001221	2.630697	0.98709	0.887931	0.005839	2.45174	0.980805	0.905882	$3.75 \times 10^{-5}$
EC-Earth3-Veg	2.518135	0.980041	0.901782	0.000944	2.760198	0.9801	0.877507	$5.90 \times 10^{-7}$	2.653931	0.976477	0.887252	0.005208
FGOALS-f3-L	2.392895	0.982742	0.912669	$9.79 \times 10^{-8}$								
FIO-ESM-2-0	2.39042	0.986034	0.913127	0.000604	2.660444	0.989316	0.888524	0.003313	2.483592	0.98829	0.906382	$3.56 \times 10^{-6}$
GFDL-ESM4	2.415156	0.991749	0.911421	0.002598	2.702498	0.984136	0.884943	0.000381	2.479362	0.99217	0.902494	0.006418
GISS-E2-1-G	2.2951	0.992667	0.919373	$2.03 \times 10^{-5}$	2.67767	0.985282	0.885917	0.001097	2.363976	0.997295	0.916792	0.000367
INM-CM5-0	2.313318	0.994196	0.917327	$5.45 \times 10^{-5}$	2.688034	0.976984	0.884853	0.000201	2.494703	0.991961	0.904855	0.000691
IPSL-CM6A-LR	2.366487	0.986157	0.915197	$2.53 \times 10^{-5}$	2.709299	0.959537	0.88337	0.000268	2.433205	0.989767	0.908817	0.001011
MIROC6	2.538982	0.975222	0.903379	0.002419	2.944748	0.977422	0.86446	0.008182	2.534029	0.982993	0.902185	0.000318
MPI-ESM-1-2-HAM	2.426283	0.988153	0.911244	$7.45 \times 10^{-5}$	2.795797	0.97782	0.8774	0.001699	2.472226	0.993353	0.906509	0.000618
MPI-ESM1-2-HR	2.482515	0.985307	0.90934	0.000179	2.813033	0.984801	0.87578	0.000855	2.478992	0.990503	0.906471	$2.68 \times 10^{-5}$
MPI-ESM1-2-LR	2.440139	0.990895	0.912516	0.000608	2.825583	0.987112	0.87801	0.000333	2.44428	0.994861	0.909502	0.001426
MRI-ESM2-0	2.430481	0.99084	0.911775	0.000386	2.754317	0.989183	0.880503	$5.16 \times 10^{-7}$	2.472997	0.993823	0.907291	0.001735
NESM3	2.393744	0.988131	0.913526	0.00115	2.625448	0.989953	0.890104	$8.44 \times 10^{-5}$	2.475125	0.992128	0.907585	0.002775

Note: MAE, S, SS, and TS are mean absolute error, Taylor diagram-based quantifiers, spatial skill scores, and temporal skill scores, respectively. MAE is measured in °C, and the others are dimensionless indicators. The closer MAE and TS are to 0, the better the simulation ability of the model; the closer S and SS are to 1, the better the simulation ability of the model.

## 5. Conclusions

Based on the 24 GCMs in CMIP6 and temperature data with high spatial and temporal resolution generated by the delta statistical downscaling model, this study assessed the evolution of  $ET_0$  in the YRB under four emission scenarios (SSP126, SSP245, SSP370, and SSP585) for the near (2022–2040), mid- (2041–2060), and long (2081–2100) term future. The major conclusions are as follows:

The regionally high-precision climate data generated by delta statistical downscaling based on multiple interpolation methods reduced the uncertainty in the GCM dataset. For the YRB, tas selected the climate models ACCESS-ESM1-5, CMCC-CM2-SR5, and INM-CM5-0; tasmax selected ACCESS-CM2, ACCESS-ESM1-5, and MRI-ESM2-0; and tasmin selected ACCESS-CM2, ACCESS-ESM1-5, and MPI-ESM1-2-LR. The equal-weighted multi-model ensemble had smaller mean absolute errors and higher correlation coefficients than single climate models, and CMIP6 efficiently simulated the temperature and  $ET_0$  in the YRB.

Compared with that of the historical period (1901–2014), the annual  $ET_0$  in the YRB under different emission scenarios (SSP126, SSP245, SSP370, and SSP585) in the future (2022–2100) substantially increased; the rate increased with the increase in emission concentration, and the  $ET_0$  in 2100 under the SSP585 scenario reached 1170.39 mm. Morlet wavelet analysis revealed that  $ET_0$  in the YRB had cyclic patterns of 34–38, 34, 39, and 27–32 years under the SSP126, SSP245, SSP370, and SSP585 scenarios, respectively.

Compared with that in the historical period, the  $ET_0$  variation in the YRB in the near-, mid-, and long-term future under different future scenarios exhibited strong spatial heterogeneity. EOF analysis revealed that  $ET_0$  had positive EOF1 values under all four emission scenarios, exhibiting a spatially consistent trend of  $ET_0$  variation across the region. A maximum variation of 112.91% occurred in the western part of the YRB in the long-term future (2081–2100) under the SSP585 scenario. Without a scientific response, future increases in  $ET_0$  could further reduce the shortage of water resources in the YRB.

**Supplementary Materials:** The following supporting information can be downloaded at: <https://www.mdpi.com/article/10.3390/rs14225674/s1>. Figure S1: Interannual variations in tas (a)/tasmax (b)/tasmin (c) in the Yellow River Basin over the historical period (1901–2014) and under different future emission scenarios; Figure S2: Spatial variations in the near (2022–2040; (a,d,g,j)), mid- (2041–2060; (b,e,h,k)), and long (2081–2100; (c,f,i,l)) term future difference between tasmax and tasmin of the Yellow River Basin relative to the historical period (1901–2014) under four SSP scenarios (SSP126, SSP245, SSP370, and SSP585); Figure S3: Spatial variations in the near (2022–2040; (a,d,g,j)), mid- (2041–2060; (b,e,h,k)), and long (2081–2100; (c,f,i,l)) term future annual precipitation of the Yellow River Basin relative to the historical period (1901–2014) under four SSP scenarios (SSP126, SSP245, SSP370, and SSP585).

**Author Contributions:** Conceptualization, S.J. and A.W.; methodology, C.S.; validation, S.J., A.W. and C.S.; investigation, A.W.; resources, S.J.; data curation, A.W.; writing—original draft, S.J. and A.W.; writing—review and editing, C.S. and K.W.; supervision, A.W. and K.W.; funding acquisition, S.J. and C.S. All authors have read and agreed to the published version of the manuscript.

**Funding:** This research was funded by the Training Program for Young Backbone Teachers in Colleges and Universities of Henan Province (2021GGJS003), the Henan Natural Science Foundation (212300410413), the Henan Youth Talent Promotion Project (2021HYTP030), the China Postdoctoral Science Foundation (2020M672247), the Key Science and Technology Project of Henan Province (No. 222102320108), and the First-class Project Special Funding of Yellow River Laboratory (No. YRL22IR12).

**Institutional Review Board Statement:** Not applicable.

**Informed Consent Statement:** Not applicable.

**Data Availability Statement:** Not applicable.

**Conflicts of Interest:** The authors declare no conflict of interest.

## References

1. Nooni, I.K.; Hagan, D.F.T.; Wang, G.; Ullah, W.; Lu, J.; Li, S.; Dzakpasu, M.; Prempeh, N.A.; Lim Kam Sian, K.T.C. Future Changes in Simulated Evapotranspiration across Continental Africa Based on CMIP6 CNRM-CM6. *Int. J. Environ. Res. Public Health*. **2021**, *18*, 6760. [[CrossRef](#)] [[PubMed](#)]
2. Wang, T.; Tu, X.; Singh, V.P.; Chen, X.; Lin, K. Global data assessment and analysis of drought characteristics based on CMIP6. *J. Hydrol.* **2021**, *596*, 126091. [[CrossRef](#)]
3. Zhang, Y.; Fu, B.; Feng, X.; Pan, N. Response of Ecohydrological Variables to Meteorological Drought under Climate Change. *Remote Sens.* **2022**, *14*, 1920. [[CrossRef](#)]
4. Allen, R.G.; Pereira, L.S.; Raes, D.; Smith, M. *Crop Evapotranspiration-Guidelines for Computing Crop Water Requirements*; FAO Irrigation and Drainage Paper 56; FAO: Rome, Italy, 1998.
5. Liu, Y.; Yue, Q.; Wang, Q.; Yu, J.; Zheng, Y.; Yao, X.; Xu, S. A Framework for Actual Evapotranspiration Assessment and Projection Based on Meteorological, Vegetation and Hydrological Remote Sensing Products. *Remote Sens.* **2021**, *13*, 3643. [[CrossRef](#)]
6. Kundu, S.; Mondal, A.; Khare, D.; Hain, C.; Lakshmi, V. Projecting Climate and Land Use Change Impacts on Actual Evapotranspiration for the Narmada River Basin in Central India in the Future. *Remote Sens.* **2018**, *10*, 578. [[CrossRef](#)]
7. Shi, L.; Feng, P.; Wang, B.; Liu, D.L.; Zhang, H.; Liu, J.; Yu, Q. Assessing future runoff changes with different potential evapotranspiration inputs based on multi-model ensemble of CMIP5 projections. *J. Hydrol.* **2022**, *612*, 128042. [[CrossRef](#)]
8. Zeng, P.; Sun, F.; Liu, Y.; Feng, H.; Zhang, R.; Che, Y. Changes of potential evapotranspiration and its sensitivity across China under future climate scenarios. *Atmos. Res.* **2021**, *261*, 105763. [[CrossRef](#)]
9. Guo, D.; Westra, S.; Maier, H.R. Sensitivity of potential evapotranspiration to changes in climate variables for different Australian climatic zones. *Hydrol. Earth Syst. Sci.* **2017**, *21*, 2107–2126. [[CrossRef](#)]
10. Jhajharia, D.; Dinpashoh, Y.; Kahya, E.; Singh, V.P.; Fakheri-Fard, A. Trends in reference evapotranspiration in the humid region of northeast India. *Hydrol. Process.* **2012**, *26*, 421–435. [[CrossRef](#)]
11. Kuang, X.; Jiao, J.J. Review on climate change on the Tibetan Plateau during the last half century. *J. Geophys. Res. Atmos.* **2016**, *121*, 3979–4007. [[CrossRef](#)]
12. PascoliniCampbell, M.; Reager, J.T.; Chandanpurkar, H.A.; Rodell, M. Retraction Note: A 10 per cent increase in global land evapotranspiration from 2003 to 2019. *Nature* **2022**, *604*, 202. [[CrossRef](#)] [[PubMed](#)]
13. She, D.; Xia, J.; Zhang, Y. Changes in reference evapotranspiration and its driving factors in the middle reaches of Yellow River Basin, China. *Sci. Total Environ.* **2017**, *607–608*, 1151–1162. [[CrossRef](#)] [[PubMed](#)]
14. Yagob, D.; Deepak, J.; Ahmad, F.; Vijay, P.S.; Ercan, K. Trends in reference crop evapotranspiration over Iran. *J. Hydrol.* **2011**, *399*, 422–433.
15. Huang, J.; Zhai, J.; Jiang, T.; Wang, Y.; Li, X.; Wang, R.; Xiong, M.; Su, B.; Fischer, T. Analysis of future drought characteristics in China using the regional climate model CCLM. *Clim. Dyn.* **2018**, *50*, 507–525. [[CrossRef](#)]
16. Torres, R.R.; Marengo, J.A. Climate change hotspots over South America: From CMIP3 to CMIP5 multi-model datasets. *Theor. Appl. Climatol.* **2014**, *117*, 579–587. [[CrossRef](#)]
17. Chen, H.; Guo, J.; Xiong, W.; Guo, S.; Xu, C. Downscaling GCMs using the Smooth Support Vector Machine method to predict daily precipitation in the Hanjiang Basin. *Adv. Atmos. Sci.* **2010**, *27*, 274–284. [[CrossRef](#)]
18. Harding, K.J.; Snyder, P.K. Examining future changes in the character of Central U.S. warm-season precipitation using dynamical downscaling. *J. Geophys. Res.-Atmos.* **2014**, *119*, 13116–13136. [[CrossRef](#)]
19. Giorgi, F.; Mearns, L.O. Approaches to the simulation of regional climate change: A review. *Rev. Geophys.* **1991**, *29*, 191–216. [[CrossRef](#)]
20. Benestad, R.E. A comparison between two empirical downscaling strategies. *Int. J. Climatol.* **2001**, *21*, 1645–1668. [[CrossRef](#)]
21. Peng, S.; Ding, Y.; Wen, Z.; Chen, Y.; Cao, Y.; Ren, J. Spatiotemporal change and trend analysis of potential evapotranspiration over the Loess Plateau of China during 2011–2100. *Agric. For. Meteorol.* **2017**, *233*, 183–194. [[CrossRef](#)]
22. Xu, J.; Gao, Y.; Chen, D.; Xiao, L.; Ou, T. Evaluation of global climate models for downscaling applications centred over the Tibetan Plateau. *Int. J. Climatol.* **2017**, *37*, 657–671. [[CrossRef](#)]
23. Dosio, A.; Panitz, H.; Schubert-Frisius, M.; Lüthi, D. Dynamical downscaling of CMIP5 global circulation models over CORDEX-Africa with COSMO-CLM: Evaluation over the present climate and analysis of the added value. *Clim. Dyn.* **2015**, *44*, 2637–2661. [[CrossRef](#)]
24. Wang, L.; Chen, W. A CMIP5 multimodel projection of future temperature, precipitation, and climatological drought in China. *Int. J. Climatol.* **2014**, *34*, 2059–2078. [[CrossRef](#)]
25. Ahmadi, H.; Baaghdeh, M. Assessment of anomalies and effects of climate change on reference evapotranspiration and water requirement in pistachio cultivation areas in Iran. *Arab. J. Geosci.* **2020**, *13*, 332. [[CrossRef](#)]
26. Das, S.; Datta, P.; Sharma, D.; Goswami, K. Trends in Temperature, Precipitation, Potential Evapotranspiration, and Water Availability across the Teesta River Basin under 1.5 and 2 °C Temperature Rise Scenarios of CMIP6. *Atmosphere*. **2022**, *13*, 941. [[CrossRef](#)]
27. Xing, W.; Wang, W.; Shao, Q.; Peng, S.; Yu, Z.; Yong, B.; Taylor, J. Changes of reference evapotranspiration in the Haihe River Basin: Present observations and future projection from climatic variables through multi-model ensemble. *Glob. Planet. Change* **2014**, *115*, 1–15. [[CrossRef](#)]

28. Liu, X.; Li, C.; Zhao, T.; Han, L. Future changes of global potential evapotranspiration simulated from CMIP5 to CMIP6 models. *Atmos. Ocean. Sci. Lett.* **2020**, *13*, 568–575. [[CrossRef](#)]
29. Nistor, M.; Mindrescu, M.; Petrea, D.; Nicula, A.; Rai, P.K.; Benzaghta, M.A.; Dezsi, Ş.; Hognogi, G.; Porumb-Ghiurco, C.G. Climate change impact on crop evapotranspiration in Turkey during the 21st Century. *Meteorol. Appl.* **2019**, *26*, 442–453. [[CrossRef](#)]
30. Zhang, Q.; Zhang, Z.; Shi, P.; Singh, V.P.; Gu, X. Evaluation of ecological instream flow considering hydrological alterations in the Yellow River basin, China. *Glob. Planet. Change* **2018**, *160*, 61–74. [[CrossRef](#)]
31. Ding, Y.; Peng, S. Spatiotemporal change and attribution of potential evapotranspiration over China from 1901 to 2100. *Theor. Appl. Climatol.* **2021**, *145*, 79–94. [[CrossRef](#)]
32. Pan, S.; Xu, Y.; Xuan, W.; Gu, H.; Bai, Z. Appropriateness of Potential Evapotranspiration Models for Climate Change Impact Analysis in Yarlung Zangbo River Basin, China. *Atmosphere* **2019**, *10*, 453. [[CrossRef](#)]
33. Zhu, Y.; Chang, J.; Huang, S.; Huang, Q. Characteristics of integrated droughts based on a nonparametric standardized drought index in the Yellow River Basin, China. *Hydrol. Res.* **2016**, *47*, 454–467. [[CrossRef](#)]
34. Sheng, X.Y.; Wang, J.K.; Cui, Q.T.; Zhang, W.; Zhu, X.F. A feasible biochar derived from biogas residue and its application in the efficient adsorption of tetracycline from an aqueous solution. *Environ. Res.* **2022**, *207*, 112175. [[CrossRef](#)] [[PubMed](#)]
35. Yin, L.; Feng, X.; Fu, B.; Wang, S.; Wang, X.; Chen, Y.; Tao, F.; Hu, J. A coupled human-natural system analysis of water yield in the Yellow River basin, China. *Sci. Total Environ.* **2021**, *762*, 143141. [[CrossRef](#)] [[PubMed](#)]
36. Zhao, F.; Ma, S.; Wu, Y.; Qiu, L.; Wang, W.; Lian, Y.; Chen, J.; Sivakumar, B. The role of climate change and vegetation greening on evapotranspiration variation in the Yellow River Basin, China. *Agric. For. Meteorol.* **2022**, *316*, 108842. [[CrossRef](#)]
37. O'Neill, B.C.; Tebaldi, C.; van Vuuren, D.P.; Eyring, V.; Friedlingstein, P.; Hurtt, G.; Knutti, R.; Kriegler, E.; Lamarque, J.; Lowe, J.; et al. The Scenario Model Intercomparison Project (ScenarioMIP) for CMIP6. *Geosci. Model Dev.* **2016**, *9*, 3461–3482. [[CrossRef](#)]
38. Zhang, Z.H.; Deng, S.F.; Zhao, Q.D.; Zhang, S.Q.; Zhang, X.W. Projected glacier meltwater and river run-off changes in the Upper Reach of the Shule River Basin, north-eastern edge of the Tibetan Plateau. *Hydrol. Process.* **2019**, *33*, 1059–1074. [[CrossRef](#)]
39. Hay, L.E.; Wilby, R.; Leavesley, G.H. A comparison of delta change and downscaled GCM scenarios for three mountainous basins in the United States. *J. Am. Water Resour. Assoc.* **2000**, *36*, 387–397. [[CrossRef](#)]
40. Willmott, C.J.; Matsuura, K. Advantages of the mean absolute error (MAE) over the root mean square error (RMSE) in assessing average model performance. *Clim. Res.* **2005**, *30*, 79–82. [[CrossRef](#)]
41. Taylor, K.E. Summarizing multiple aspects of model performance in a single diagram. *J. Geophys. Res. Atmos.* **2001**, *106*, 7183–7192. [[CrossRef](#)]
42. Jiang, Z.; Song, J.; Li, L.; Chen, W.; Wang, Z.; Wang, J. Extreme climate events in China: IPCC-AR4 model evaluation and projection. *Clim. Change* **2012**, *110*, 385–401. [[CrossRef](#)]
43. Chen, W.; Jiang, Z.; Li, L. Probabilistic Projections of Climate Change over China under the SRES A1B Scenario Using 28 AOGCMs. *J. Clim.* **2011**, *24*, 4741–4756. [[CrossRef](#)]
44. Jiang, D.; Wang, H.; Lang, X. Evaluation of East Asian Climatology as Simulated by Seven Coupled Models. *Adv. Atmos. Sci.* **2005**, *22*, 479–495.
45. Ma, Z.; Sun, P.; Zhang, Q.; Zou, Y.; Lv, Y.; Li, H.; Chen, D. The Characteristics and Evaluation of Future Droughts across China through the CMIP6 Multi-Model Ensemble. *Remote Sens.* **2022**, *14*, 1097. [[CrossRef](#)]
46. Pelosi, A.; Medina, H.; Villani, P.; Urso, D.G.; Chirico, G.B. Probabilistic forecasting of reference evapotranspiration with a limited area ensemble prediction system. *Agric. Water Manag.* **2016**, *178*, 106–118. [[CrossRef](#)]
47. Yan, X.; Mohammadian, A. Estimating future daily pan evaporation for Qatar using the Hargreaves model and statistically downscaled global climate model projections under RCP climate change scenarios. *Arab. J. Geosci.* **2020**, *13*, 938. [[CrossRef](#)]
48. Hargreaves, G.H. Moisture Availability and Crop Production. *Trans. ASAE* **1975**, *18*, 980–9984. [[CrossRef](#)]
49. Hargreaves, G.H.; Samani, Z.A. Estimating Potential Evapotranspiration. *J. Irrig. Drain. Div.* **1982**, *108*, 225–230. [[CrossRef](#)]
50. Hargreaves, G.H.; Samani, Z.A. Reference Crop Evapotranspiration from Temperature. *Appl. Eng. Agric.* **1985**, *2*, 96–99. [[CrossRef](#)]
51. Li, H.; Wang, Y.; Jia, L.; Wu, Y.; Xie, M. Runoff Characteristics of the Nen River Basin and its Cause. *J. Mt. Sci.* **2014**, *11*, 110–118. [[CrossRef](#)]
52. Naren, A.; Maity, R. Modeling of local sea level rise and its future projection under climate change using regional information through EOF analysis. *Theor. Appl. Climatol.* **2018**, *134*, 1269–1285. [[CrossRef](#)]
53. North, G.R.; Bell, T.L.; Cahalan, R.F.; Moeng, F.J. Sampling Errors in the Estimation of Empirical Orthogonal Functions. *Am. Meteorol. Soc.* **1982**, *110*, 699–706. [[CrossRef](#)]
54. Singh, V.P.; Xu, C.Y. Evaluation and generalization of 13 mass-transfer equations for determining free water evaporation. *Hydrol. Process.* **1997**, *3*, 311–323. [[CrossRef](#)]
55. Blaney, H.F.; Criddle, W.D. Determining Consumptive Use and Irrigation Water Requirements. *Tech. Bull.* **1962**, *4*, 369–373.
56. Rohwer, C. *Evaporation from a Free Water Surface*; US Department of Agriculture: Washington, DC, USA, 1931.
57. Penman, H.L. Natural Evaporation from Open Water, Bare Soil and Grass. *Proc. R. Soc. Lond. Ser. A Math. Phys. Sci.* **1948**, *193*, 120–145.
58. Priestley, C.H.B.; Taylor, R.J. On the Assessment of Surface Heat Flux and Evaporation Using Large-Scale Parameters. *Mon. Weather Rev.* **1972**, *100*, 81–92. [[CrossRef](#)]
59. Jensen, M.E.; Haise, H.R. Estimating Evapotranspiration from Solar Radiation. *J. Irrig. Drain. Div.* **1963**, *89*, 15–41. [[CrossRef](#)]

60. Peng, S.; Gang, C.; Cao, Y.; Chen, Y. Assessment of climate change trends over the Loess Plateau in China from 1901 to 2100. *Int. J. Climatol.* **2018**, *38*, 2250–2264. [[CrossRef](#)]
61. Yang, X.; Zhou, B.; Xu, Y.; Han, Z. CMIP6 Evaluation and Projection of Temperature and Precipitation over China. *Adv. Atmos. Sci.* **2021**, *38*, 817–830. [[CrossRef](#)]
62. Er-Raki, S.; Chehbouni, A.; Khabba, S.; Simonneau, V.; Jarlan, L.; Ouldbba, A.; Rodriguez, J.C.; Allen, R. Assessment of reference evapotranspiration methods in semi-arid regions: Can weather forecast data be used as alternate of ground meteorological parameters? *J. Arid. Environ.* **2010**, *74*, 1587–1596. [[CrossRef](#)]
63. Zhao, C.; Nan, Z.; Feng, Z. GIS-assisted spatially distributed modeling of the potential evapotranspiration in semi-arid climate of the Chinese Loess Plateau. *J. Arid. Environ.* **2004**, *58*, 387–403.
64. Bourletsikas, A.; Argyrokastritis, I.; Proutsos, N. Comparative evaluation of 24 reference evapotranspiration equations applied on an evergreen-broadleaved forest. *Hydrol. Res.* **2018**, *49*, 1028–1041. [[CrossRef](#)]
65. Lang, D.; Zheng, J.; Shi, J.; Liao, F.; Ma, X.; Wang, W.; Chen, X.; Zhang, M. A Comparative Study of Potential Evapotranspiration Estimation by Eight Methods with FAO Penman–Monteith Method in Southwestern China. *Water* **2017**, *9*, 734. [[CrossRef](#)]
66. Yan, X.; Mohammadian, A. Forecasting daily reference evapotranspiration for Canada using the Penman–Monteith model and statistically downscaled global climate model projections. *Alex. Eng. J.* **2020**, *59*, 883–891. [[CrossRef](#)]
67. Omer, A.; Elagib, N.A.; Ma, Z.; Saleem, F.; Mohammed, A. Water scarcity in the Yellow River Basin under future climate change and human activities. *Sci. Total Environ.* **2020**, *749*, 141446. [[CrossRef](#)]
68. Jiang, Z.; Yang, Z.; Zhang, S.; Liao, C.; Hu, Z.; Cao, R.; Wu, H. Revealing the spatio-temporal variability of evapotranspiration and its components based on an improved Shuttleworth–Wallace model in the Yellow River Basin. *J. Environ. Manag.* **2020**, *262*, 110310. [[CrossRef](#)]
69. Roderick, M.L.; Farquhar, G.D. The cause of decreased pan evaporation over the past 50 years. *Science* **2002**, *298*, 1410–1411. [[CrossRef](#)]
70. Li, Z.; Zheng, F.; Liu, W. Spatiotemporal characteristics of reference evapotranspiration during 1961–2009 and its projected changes during 2011–2099 on the Loess Plateau of China. *Agric. For. Meteorol.* **2012**, *154–155*, 147–155. [[CrossRef](#)]
71. Wang, W.; Shao, Q.; Peng, S.; Xing, W.; Yang, T.; Luo, Y.; Yong, B.; Xu, J. Reference evapotranspiration change and the causes across the Yellow River Basin during 1957–2008 and their spatial and seasonal differences. *Water Resour. Res.* **2012**, *48*, 113–122. [[CrossRef](#)]
72. Cook, P.A.; Black, E.C.L.; Verhoef, A.; Macdonald, D.M.J.; Sorensen, J.P.R. Projected increases in potential groundwater recharge and reduced evapotranspiration under future climate conditions in West Africa. *J. Hydrol. Reg. Stud.* **2022**, *41*, 101076. [[CrossRef](#)]
73. Gharbia, S.S.; Smullen, T.; Gill, L.; Johnston, P.; Pilla, F. Spatially distributed potential evapotranspiration modeling and climate projections. *Sci. Total Environ.* **2018**, *633*, 571–592. [[CrossRef](#)] [[PubMed](#)]
74. Dallaire, G.; Poulin, A.; Arsenaault, R.; Brissette, F. Uncertainty of potential evapotranspiration modelling in climate change impact studies on low flows in North America. *Hydrol. Sci. J.* **2021**, *66*, 689–702. [[CrossRef](#)]
75. Liu, J.; Lu, Y. How Well Do CMIP6 Models Simulate the Greening of the Tibetan Plateau? *Remote Sens.* **2022**, *14*, 4633. [[CrossRef](#)]
76. Stouffer, R.J.; Eyring, V.; Meehl, G.A.; Bony, S.; Senior, C.; Stevens, B.; Taylor, K.E. CMIP5 Scientific Gaps and Recommendations for CMIP6. *Bull. Amer. Meteorol. Soc.* **2017**, *98*, 95–105. [[CrossRef](#)]
77. Le, T.; Bae, D. Response of global evaporation to major climate modes in historical and future Coupled Model Intercomparison Project Phase 5 simulations. *Hydrol. Earth Syst. Sci.* **2020**, *24*, 1131–1143. [[CrossRef](#)]
78. Gao, C.; Booij, M.J.; Xu, Y. Assessment of extreme flows and uncertainty under climate change: Disentangling the uncertainty contribution of representative concentration pathways, global climate models and internal climate variability. *Hydrol. Earth Syst. Sci.* **2020**, *24*, 3251–3269. [[CrossRef](#)]
79. Meinshausen, M.; Lewis, J.; McGlade, C.; Gütschow, J.; Nicholls, Z.; Burdon, R.; Cozzi, L.; Hackmann, B. Realization of Paris Agreement pledges may limit warming just below 2 °C. *Nature* **2022**, *604*, 304–309. [[CrossRef](#)]