

## Article

# Development of Groundwater Flow Models for the Integrated Management of the Alluvial Aquifer Systems of *Dravsko polje* and *Ptujsko polje*, Slovenia

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**Abstract:** With increasing exploitation of groundwater resources and implementation of various activities in their recharge areas, it is vital to conduct a comprehensive assessment of aquifers to ensure their conservation and sustainable management. In the present study, we used a comprehensive approach to conceptualise and identify the functioning of two connected aquifer systems in north-eastern Slovenia: the Quaternary porous aquifers *Dravsko polje* and *Ptujsko polje*. The study presents the conceptual models of both aquifers and their interconnectedness using separate mathematical-numerical models with the aim of ensuring an integrated management of these alluvial aquifer systems. It also highlights the importance of understanding connections between such systems for simulating groundwater flow and transport of different contaminants. To describe the entire aquifer system, the study defines its three essential elements: the geometry of the aquifers, their recharge by precipitation, and other boundary conditions. The geometry of the Quaternary aquifers was defined using Sequential Indicator Simulation (SIS) with the ESRI's ArcMap software. Next, LIDAR was used for determining their surface geometry. The hydrogeologic model was designed using the Groundwater Modelling System (GMS) developed by AQUAVEO. We used the MODFLOW 2000 calculation method based on the finite difference method (FDM). The model was calibrated with the PEST module, which was used to calibrate hydraulic conductivity and hydraulic heads between the measured and modelled data. Finally, the model was validated using the Nash–Sutcliffe (NSE) efficiency coefficient. In addition, the model results estimated using the PEST tool were validated with the hydraulic conductivities determined at the pumping sites (pumping tests), each belonging to water protection zones that define the maximum travel time of the particles. This was performed using the MODPATH method. The paper also presents the possibility of modelling heterogeneous but interdependent aquifers in a groundwater body. Modelling the connection between the two aquifers, which are the most important ones in the region, is essential for a comprehensive management of the entire system of water resources. The models allow for a better understanding of groundwater flow in both aquifers. Moreover, their interconnectedness will be used for further studies in this field, as well as for integrated water management.

**Keywords:** hydrogeological model; Modflow model; aquifer geometry; aquifer recharge; modelling

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## 1. Introduction

Groundwater presents an important source of drinking water, which is why its use and protection has become an important global issue [1]. To achieve comprehensive and integrated management of groundwater resources, various conceptual models were developed worldwide [2–11]. The development of conceptual models relies on frameworks expressing the cause-and-effect structure of groundwater systems. Therefore, different frameworks should be used when dealing with complex natural systems such as aquifers [12]. The main objective of this study is to develop an integrated model of two separate but interconnected aquifers, which will provide a management tool facilitating decision making for the

prevention of groundwater pollution. Regardless of its vital importance to water supply, their recharge area is also a site of diverse and intensive land use, resulting in conflicts of interest. Considering all these aspects, it is imperative to raise the knowledge of the aquifers used in this area. The *Dravska kotlina* groundwater body is located in north-eastern Slovenia and contains two of the largest, interconnected, and most important aquifers used as water resources in Slovenia: the *Dravsko polje* aquifer and the *Ptujsko polje* aquifer. Both are filled with Quaternary sediments of the Drava River and are intergranular aquifers. Their water tables are near the surface and have an average depth of about 8 m. Both aquifers are unconfined and have no overlying low-permeability layers that could protect them from possible pollution. Considering this, they are the most strategic groundwater resources in Slovenia. In recent decades, this area has been affected by high nitrate and pesticide concentrations caused by anthropogenic sources such as animal manure used as fertilizer, synthetic fertilizers, septic systems, and other waste waters. The study area of the two aquifers is intensively used for agriculture. There are many livestock farms in the aquifer recharge area. Contamination of groundwater from diffuse agricultural sources is known to be a major pollution problem [13–17].

The protection of groundwater sources begins with the improvement of groundwater management and the effective implementation of environmental protection measures. The EU Water Framework Directive (2000/60/EC) [18] and its daughter Directive on the Protection of Groundwater Against Pollution (2006/118/EC) [19] established the criteria for defining the qualitative and quantitative status of groundwater. Slovenia recognised this directive in its national regulation: the Water Act [20] and the Regulation on the Protection of Waters Against Pollution Caused by Nitrates from Agricultural Sources [21]. The main objective of this research is to develop a conceptual and hydraulic groundwater flow model on a regional scale, starting from borehole data, grain size distributions and well tests, in order to analyse the overall groundwater flow behaviour. Problems associated with the conceptualisation of groundwater systems have been known for over a decade [22,23]. In recent years, increasing attention has been paid to the development and evaluation of conceptual models [24,25]. The development of an appropriate conceptual model is a critical step in any successful modelling study. A simple model is needed that would be able to capture the behaviour of total groundwater flow in the main aquifers without complex fine-resolution behaviours arising from aquifer heterogeneity. On the other hand, despite their simplicity, such conceptual and numerical models must be accurate and hydrostratigraphically sound. In designing the conceptual models, we defined the study area, created a 3D model of the hydrogeological system, estimated the groundwater recharge from precipitation and surface waters, and defined the boundary conditions (BCs). Previous assessments of groundwater recharge from precipitation and its spatial distribution in the study area were sparse and insufficient. Groundwater recharge in our models was divided into two categories: targeted recharge from surface waters such as rivers, canals, and lakes, and diffuse recharge from infiltration of precipitation through the unsaturated zone into the groundwater. The analysis of groundwater and surface water levels allowed us to define the BCs and the relationship between surface water and groundwater.

This study is the first to evaluate the spatial distribution of long-term average groundwater recharge from precipitation in the *Dravsko polje* and *Ptujsko polje* aquifers as part of the larger *Dravska kotlina* water body. The objective of the research was to develop and present the first hydrogeological model of the interconnected *Dravsko polje* and *Ptujsko polje* aquifers as a basis for integrated management of both aquifers. A detailed overview of the hydrodynamic characteristics of the aquifers was prepared in order to develop a hydrogeological conceptual model that will serve as the foundation for the construction of a numerical groundwater flow and contaminant transport models. The results, along with the geometry of the aquifer and other BCs, can serve as an important input to the development of a numerical model for various contaminants. The model represents a part of the decision support system for groundwater management at national and local levels. Land use in the area under consideration is diverse, and there are considerable conflicts

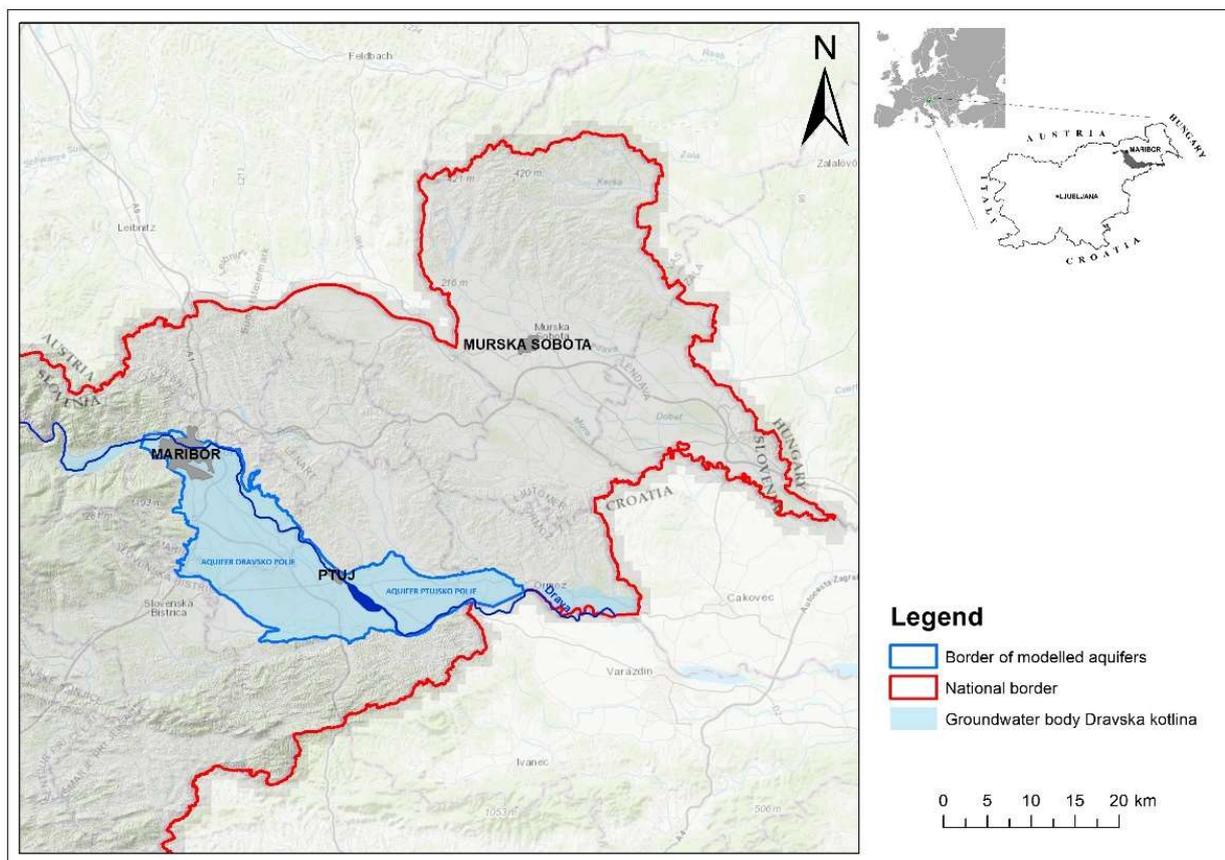
of interest related to various activities. The model provides the basis for an assessment of the qualitative and quantitative state of the water body, as well as for monitoring its changes due to climate change. The model facilitates the assessment of potential sources of pollution and the development of measures to improve groundwater quality. It provides support for groundwater management intended for water supply, irrigation, and other specific uses. The model also serves to verify the proper delineation of drinking water protection areas.

## 2. Materials and Methods

### 2.1. Study Area

#### 2.1.1. Location and Climate of the Study Area

The study area includes two aquifers—*Dravsko polje* and *Ptujsko polje*—which are both part of the *Dravska kotlina* water body (Figure 1). Both aquifers are hydraulically connected.



**Figure 1.** *Dravska kotlina* groundwater body with marked modelled aquifers.

The *Dravsko polje* aquifer is located on the right bank of the Drava River. It covers an area of 293 km<sup>2</sup> and extends from Maribor in the north to Ptuj in the southeast and Slovenska Bistrica in the southwest. In general, it has the shape of a triangle between Maribor, Ptuj, and Pragersko. It is bordered by the Pohorje Massif to the west, Slovenske Gorice Hills to the east and north, and by Haloze and Dravinjske Gorice to the south. A supply canal for the Zlatoličje Hydro Power Plant runs parallel to the Drava River. In the town of Ptuj, the Drava flows into Lake Ptuj, an artificial lake from which the river continues to flow. The *Dravsko polje* aquifer began to form in the Pleistocene when the Drava eroded tertiary sediments and began filling it with sediments. During this process, the Drava also eroded its own sediments, which can be seen in many terraces [26]. There are four terraces; the highest of which is at an altitude of 270 metres. The two upper terraces are inclined towards Ptuj and have a slope of 1.8% [27]. The *Dravsko polje graben* is filled with

Quaternary sediments [26]. The sedimentology of the aquifer is heterogeneous and changes rapidly in the vertical direction. It is mostly composed of gravel and sand and contains layers and lenses of sand. In some areas, there are also clay deposits. Conglomerate is found mainly on the edges of the terraces.

The *Ptujsko polje* aquifer is located on the left bank of the Drava River and extends between Ptuj and the town of Ormož. It borders the *Dravsko polje* aquifer to the west and the *Ormoško polje* aquifer to the east. It is bordered by the hills of Slovenske Gorice to the north and by the Drava River to the south. It has the shape of an elongated triangle with an area of about 91 km<sup>2</sup>. The surface is slightly inclined to the east. The *Ptujsko polje* aquifer is filled with Quaternary gravel deposits, the thickness of which varies from 4 m up to 22 m. The aquifer is composed of a very thin layer of gravel. The tertiary base of the Quaternary aquifer consists of conglomerate sand, clay and marl with very low permeability. In the vicinity of Lake Ptuj on the western border, the Ljutomer fault runs in the direction of SW to NE. Between Ptuj and the Ljutomer fault, the base of the Quaternary aquifer consists of Pliocene sediments. The rest of the area east of the fault base consists of Lower Pliocene and Miocene sediments. The tertiary base is strongly undulated [26]. The most important watercourse in the *Ptujsko polje* aquifer is the Drava River. It flows along the southern edge of the aquifer to Ormož. After the construction of the Formin Hydro Power Plant between Markovci and the town of Ptuj, the 5 km long and 1.2 km wide Lake Ptuj was created. A flood control dam was built along this artificial lake. When the Formin Hydro Power Plant was built, the *Ptujsko polje* aquifer was subdivided by an 8.5 km long supply canal between Markovci and Formin and an 8 km long guide canal between Formin and the town of Ormož [26].

The intersection between the two models is the area between the Drava River and Lake Ptuj (Figure 1). Based on the flow budget of the *Drava polje* aquifer, the total discharge below Lake Ptuj and the Drava River is the amount of groundwater that is expected to continue to flow from the *Dravsko polje* into the *Ptujsko polje* aquifer.

#### 2.1.2. Geology and Hydrogeology of the Study Area and Geometry of the Aquifers

To describe the geometry of the bedrock of both aquifers (Figures 2 and 3), we collected lithological data from the archives of the Geological Survey of Slovenia (GeoZS). Data were collected from 2510 boreholes drilled from Pleistocene sediments to the bedrock. The boreholes are spatially distributed throughout the study area, which allows a good definition of the geometry of the aquifer. The study area is crossed by a major fault—the Ljutomer fault. It was included in the bedrock model [28,29]. In the south of the model, the Ljutomer fault shows a slight faulting of the bedrock especially near the Drava River and Lake Ptuj. Both aquifers were treated as a single homogeneous layer consisting mainly of gravel and sand [26]. The homogeneity of the aquifer was verified by a lithological model created using the JewelSuite software [30]. We grouped the lithologic description of the sediments from the boreholes into seven categories (gravel; silt and clay with gravel; silt and clay; conglomerate; gravel, clay, silt; humus and bedrock). The subsurface was drawn using the ESRI ArcMap software with lithologic data obtained from the boreholes and expertise on the study area. Sequential Indicator Simulation (SIS) was used to interpolate the defined lithological units in the 3D model. The model extends from the surface (defined by a digital elevation model) to the bedrock of the Quaternary aquifer. Sequential Indicator Simulation is a stochastic interpolation method based on Juang et al. [31]. For the surface layer, the data from LIDAR were used as Slovenia is completely covered by LIDAR. The models are set to cells with a size of 100 × 100 m. In both models, we used a porosity value of 0.15, which is typical for porous aquifers.

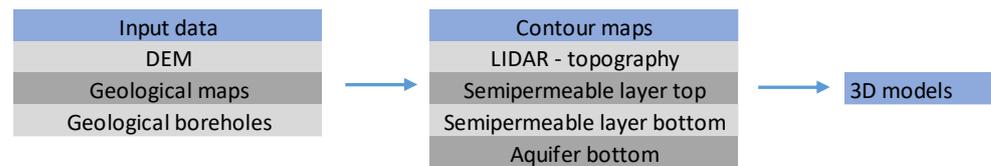


Figure 2. Methodological chart showing the required data and steps for developing 3D models.

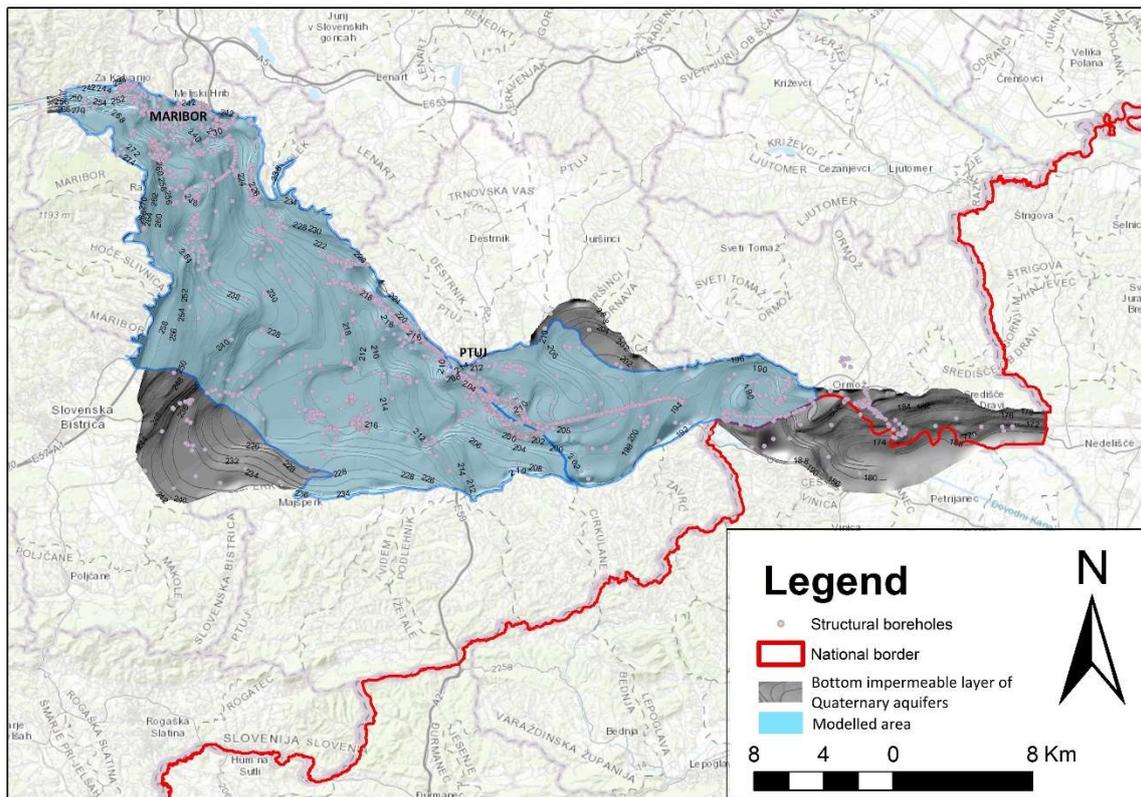
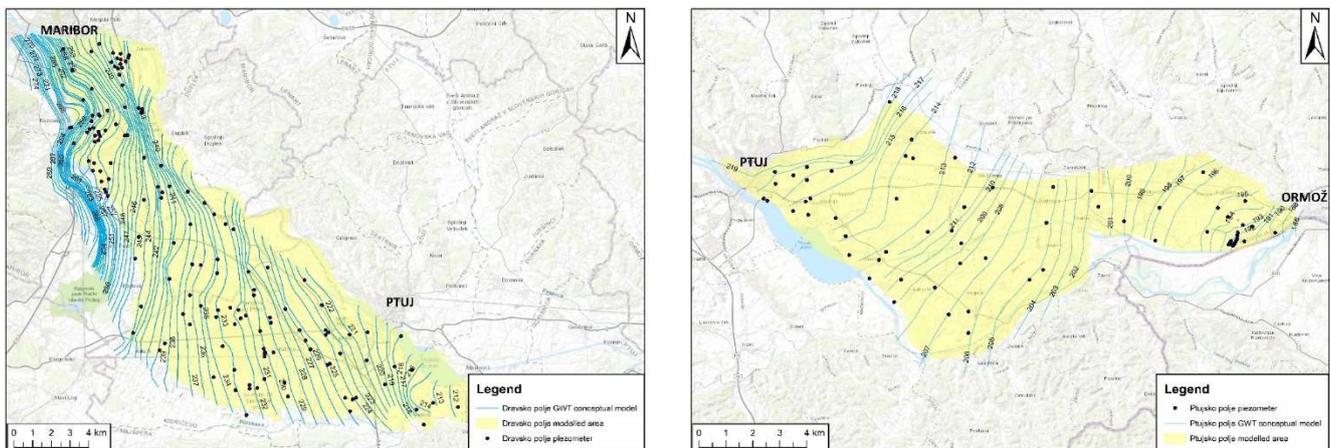


Figure 3. The modelled bottom impermeable layer of Quaternary aquifers.

## 2.2. Groundwater Flow System and Groundwater Level Monitoring

The groundwater level (GWL) and the upper and lower layers of the model were unified for both models. Based on the measured GWL values of 152 wells in the *Dravsko polje* aquifer and 86 wells in the *Ptujsko polje* aquifer, a GWL map was created using ArcGis software. In addition, the modelled GWLs were compared with the measured (observed) values. For the *Dravsko polje* aquifer, most of the data available were collected between 21 and 25 November 2012, so the GWL map was created for November 2012 (Figure 4). For the *Ptujsko polje* aquifer, most of the data available were collected on 20 March 2015 (Figure 4). For these dates, a medium water level is characteristic.



**Figure 4.** Groundwater level map of the Dravsko polje aquifer (**left**) and groundwater level map of the Ptujsko polje aquifer (**right**).

### 2.3. Groundwater Recharge Estimation

The groundwater of the *Dravsko polje* aquifer is recharged by infiltration of streams from the Pohorje Massif and by infiltration of precipitation. The *Dravsko polje* aquifer is not homogeneous; moreover, the differences in the way it is recharged are so vast that it was divided into three hydrogeological units, as described by [26]. Each of these units has a separate precipitation margin and a well-defined regime of groundwater recharge and discharge. The first unit covers the north-western part of the territory, which is mainly urbanised. It includes the foothills of the Pohorje Massif of the *Radvanjski potok* and *Razvanjski potok* creeks. In the south, it is bounded by the surface ridge of the *Razvanjski potok* creek ridgeline and the underground ridge extending from Bohova to Dogoše. In this unit, groundwater flow is quite low due to the small precipitation catchment area and the low thickness of the aquifer. The second hydrogeological unit includes part of the *Dravsko polje* aquifer between Bohova and Dogoše in the north and Hotinja vas, Dravski Dvor and Starše in the south of the *Dravsko polje* aquifer. It also includes the hilly outskirts of the *Pivolski potok*, *Hočki potok*, and *Polanski potok* creeks. Groundwater from the second unit of the *Dravsko polje* aquifer is discharged into the Miklavž Spring and the Drava River.

The third and largest hydrogeological unit includes the hilly outskirts of the *Rački potok* creek and the middle part of the *Dravsko polje* aquifer from the border of the Pohorje Massif in the west to the Drava River in the north-east and the Polskava River in the south. Groundwater is discharged from the Pohorje Massif towards the Zlatoličje Canal and the Hajdina and Pobreš springs [32].

The area is largely covered by agricultural land (44%), followed by forested land (20%) and urban (populated) areas, which represent sources of both urban and industrial pollution (19%), while the rest of the area is covered by meadows. Surface activities are also reflected in the quality of the groundwater.

Most of the recharge volume is provided by infiltration of precipitation. The average annual amount of precipitation is between 800 and 1000 mm/year. The long-term mean annual groundwater recharge is estimated at 300–450 mm [33], including evapotranspiration. To determine the long-term mean precipitation, we selected two monitoring stations: one in the north (Maribor—Airport) and one in the south (Ptuj) of the study area (Figure 4). The year with the highest precipitation was 1972 with 1284 mm/year (Ptuj).

Approximately one-third of the average annual precipitation is infiltrated and discharged into the Drava River. The recharge rate of the Quaternary aquifer is estimated at  $3 \text{ m}^3/\text{s}$  [34]. Thus, groundwater discharge equals  $2 \text{ m}^3/\text{s}$  and direct recharge equals  $0.8 \text{ m}^3/\text{s}$ . Surface runoff accounts for the smallest share of the average annual precipitation volume, and about 41% is infiltrated by precipitation (27–47%) [34]. In general, the *Dravsko polje* aquifer has good hydraulic conductivity, ranging from about 68 to 790 m/day [26].

The *Ptujsko polje* aquifer is recharged mainly by precipitation and partly by the Drava River. Groundwater discharges into numerous springs called Zvirenčine at the foot of the high Quaternary gravel terrace. The town of Ormož is supplied with drinking water from the wells in the eastern part of the *Ptujsko polje* aquifer. The supply canal, which belongs to the Formin Hydro Power Plant, affects the groundwater only to a small extent, as all necessary measures were taken to prevent the canal from polluting the groundwater. This canal is sealed only on the eastern side of the *Ptujsko polje* aquifer at a distance of 6 km from Formin in the direction of the town of Ormož to the tertiary base. However, the last 2 km are permeable, allowing water to drain into the Drava River. GWL fluctuations are caused mainly by precipitation and partly by the natural flow of the Drava. Groundwater discharges into the Drava and into springs located along the Drava. The discharge of groundwater into the Drava River is significant for the part after the confluence of the Dravinja River with the Drava River. Before the Dravinja flows into the Drava, an equilibrium is reached between the river and the groundwater, so that it is neither drained nor dammed [35].

In the eastern part between the village of Formin and the town of Ormož, the groundwater no longer flows evenly, as the Formin supply canal is sealed up by the non-permeable tertiary base. Therefore, the groundwater flow is divided into the southern and northern parts of the Formin supply canal. In the northern part, groundwater flows in a W-E direction and discharges into the Drava River, with the last 2 km of the canal being permeable again. Only precipitation has an influence on GWL fluctuation. In this part, water is tapped for the fresh water supply of the town of Ormož. In the southern part of the canal, GWL fluctuation is influenced both by precipitation and the Drava River.

#### 2.4. Conceptual Models of the Study Area

The modelled area of the *Dravsko polje* aquifer is bounded along the western edge by the 'Specified head' BC, which enables a constant flow from the Pohorje Massif to the *Dravsko polje* aquifer. The constant head corresponds to the GWL and is divided into three sections (Figure 5). From the north to the edge of the south-east, the groundwater flow is determined by two BCs: 'Drain', where the groundwater flows further down or into Lake Ptuj, and 'River', where the Drava and Dravinja rivers are located. The 'Drain' stage corresponds to the lower layer of the lake, where the groundwater is higher than the bottom layer of the lake. This is an indication that the groundwater flow exists between the tertiary bottom and the lake bottom. The bottom layer of the lake is shown in Figure 6. It was adapted from [36], which contains an image of the bottom layer, and redrawn using ArcMap.

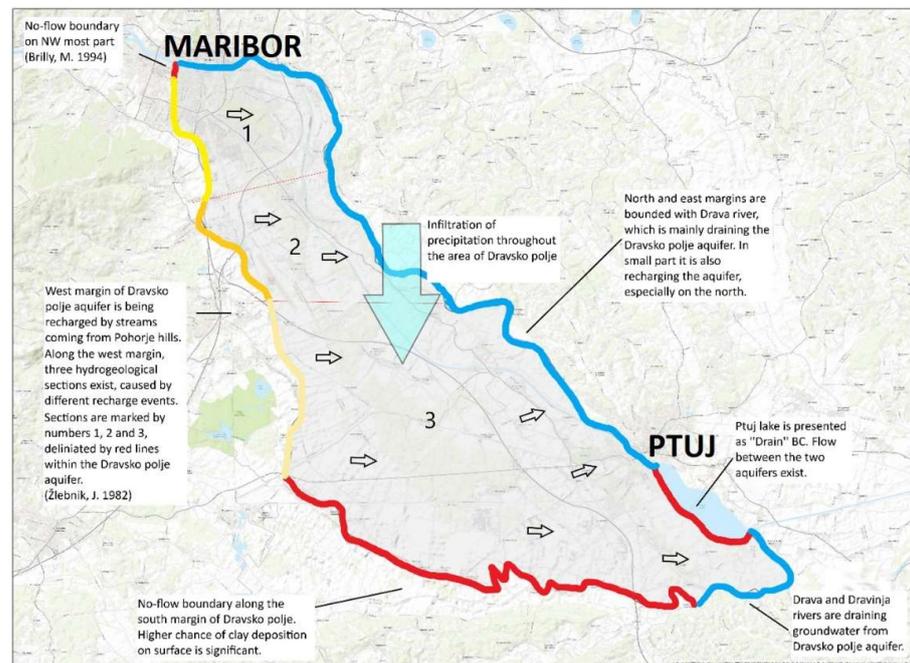


Figure 5. Conceptual model of the *Dravsko polje* aquifer, Žlebnik, L. [26], Brilly, M. [37].

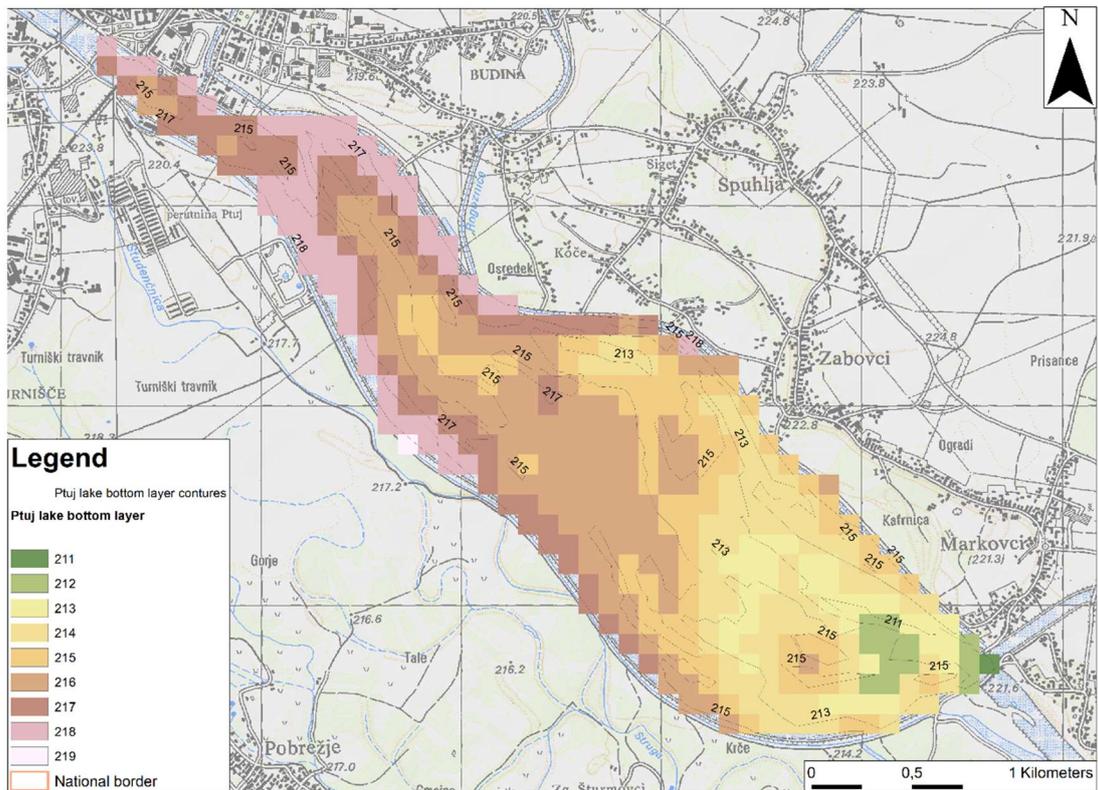


Figure 6. Bottom layer of Lake Ptuj (adapted after Veršič, 2007).

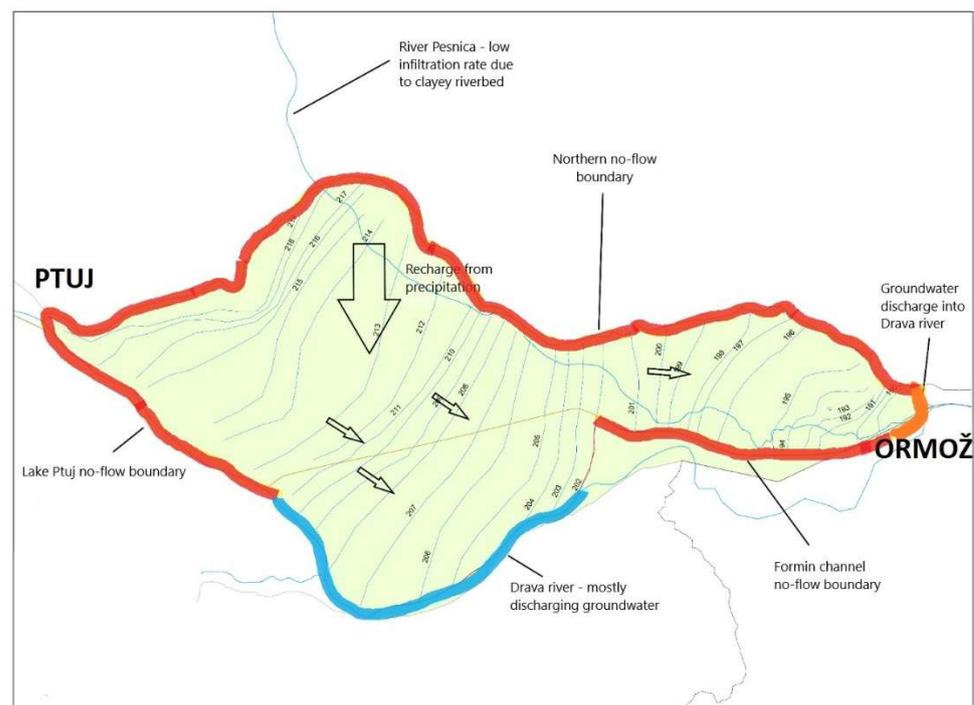
The approximate groundwater flow beneath the lake was calculated based on Darcy's flow for unconfined aquifers. Most of the groundwater is discharged into streams. However, groundwater discharge is substantial in the northern part of the river. The 'River' BC was based on the GWL map; the southern and south-western parts have a 'No-flow' BC. Infiltration of precipitation into the *Dravsko polje* aquifer is 300 mm/yr as 'Recharge' BC [38].

Near the eastern edge of the aquifer and in the south-east, there are two drainages that drain a small amount of groundwater. The model includes 7 pumping fields with a total of 10 wells (Betnava: 40 m<sup>3</sup>/day, Dobrovce: 60 m<sup>3</sup>/day, Šikole: 44.5 m<sup>3</sup>/day, Skorba: 135.5 m<sup>3</sup>/day, Kidričevo: 183 m<sup>3</sup>/day, Lancova vas: 15.5 m<sup>3</sup>/day, Bohova: 50 m<sup>3</sup>/day). The total pumping capacity is 528.5 m<sup>3</sup>/day. The porosity of the entire model was set to 0.15, as is typical for porous aquifers. The approximate values of water balance were estimated from data in the literature (Table 1).

**Table 1.** Estimation of the *Dravsko polje* aquifer groundwater flow.

	Wells	River	Recharge from Precipitation	Recharge from the Western Boundary (Pohorje)	Lake Ptuj
Flow in [m <sup>3</sup> /s]	0	50.000 [34]	106.272 [33]	259.200 [34]	0
Flow out [m <sup>3</sup> /s]	528.5	app. 400.000 (groundwater discharges mainly into the river)	0	0	50.000 (calculated based on Darcy's flow)
TOTAL in-out [m <sup>3</sup> /s]	415.472				

The conceptual model for the *Ptujsko polje* aquifer is shown in Figure 7. Infiltration by precipitation is described as 'Recharge' BC. A negligible amount of surface water from streams infiltrates from the hills of Slovenske Gorice along the northern edge of the aquifer. Therefore, to simplify the model, the northern edge of the aquifer was defined as a basin with a no-flow BC.



**Figure 7.** Conceptual model for the *Ptujsko polje* aquifer.

In the eastern part of the aquifer between the village of Formin and the town of Ormož, the supply canal leading to the Formin Hydro Power Plant represents an impermeable boundary, as it is sealed by the impermeable tertiary base of the Quaternary aquifer. However, the last 2 km of this canal are permeable and here, the groundwater is discharged into the Drava River. The eastern part of the *Ptujsko polje* aquifer is also permeable and allows groundwater to discharge into the *Ormož polje* aquifer. Therefore, the canal between the village of Formin and the town of Ormož is without discharge except for the last two kilometres. In the western part of the aquifer, groundwater is discharged mainly into the Drava River. In our model, the Drava was described as ‘River’ BC. The head stage of the River BC was based on the GWL map. On the west side between Ptuj and Markovci, artificial Lake Ptuj was represented as ‘Specified Flow’ BC. The recharge rate below the lake was taken from the model of the *Dravsko polje* aquifer, as we assumed that the same amount of groundwater flows from the *Dravsko polje* aquifer into the *Ptujsko polje* aquifer.

The Pesnica River flows through the *Ptujsko polje* aquifer towards *Ormoško polje*. It was neglected, as the amount of water infiltrating into the groundwater is low due to clay deposits at the bottom of the riverbed. In the eastern part of the *Ptujsko polje* aquifer, the *Sejanica pri Mihovcih* pumping station supplies drinking water to the town of Ormož. The water is pumped from 16 pumping wells. However, to avoid depletion of the aquifer, the groundwater is recharged by 5 infiltration fields, where surface water from the canal is infiltrated into the groundwater. The recharge rates are shown in Tables 2 and 3. Table 4 shows the groundwater balance of the *Ptujsko polje* aquifer.

**Table 2.** Recharge rates from infiltration fields at the *Ptujsko polje* aquifer for the time used in the model.

Infiltration Field	Recharge Rate (m <sup>3</sup> /day)
1	347.87
2	581.73
3	0
4	233.37
5	65.03

**Table 3.** Pumping rates on the east side of the *Ptujsko polje* aquifer for the time used in the model.

Pumping Well	Pumping Rate (m <sup>3</sup> /day)
1	255.07
2	0
3	158.97
4	246.73
5	58.83
6	218.37
7	117.50
8	262.57
9	303.77
10	271.37
11	308.87
12	140.67
13	321.10
15	167.33
16	46.57
17	73.43

**Table 4.** Estimated groundwater flow values for the *Ptujsko polje* aquifer.

	Wells	Infiltration Fields (m <sup>3</sup> /day)	Recharge from Precipitation (m <sup>3</sup> /day)	Recharge from the Western Boundary Lake Ptuj
Flow in	0	1228	77.433	15.000 (approximated based on the result from the <i>Dravsko polje</i> aquifer numerical model)
Flow out	2951.15	0	0	0
TOTAL in-out	93.661			

### 2.5. Numerical Models and Boundary Conditions (BCs)

Numerical groundwater modelling was performed using the Groundwater Modelling System (GMS 10.1.3) based on MODFLOW 2000, a comprehensive graphical user environment developed by Aquaveo. GMS is compatible with GIS-based graphical pre-processing tools to automate and streamline the modelling process for the conceptualisation and implementation of groundwater flow simulations. The GMS system consists of a graphical user interface (the GMS software) and a number of modelling codes (MODFLOW, MT3DMS, etc.) that provide tools for each phase of groundwater simulation, including site characterisation, model development, calibration, post-processing, and visualisation. GMS supports the appropriate ArcMap version as well as finite difference and finite element models in 2D and 3D, including MOD-FLOW 2000, MODPATH, MT3DMS/RT3D, SEAM3D, FEMWATER, PEST, UTEXAS, MODAEM, and SEEP2D [39].

The *Dravsko polje* and *Ptujsko polje* aquifer models were designed separately. First, conceptual models were created for each area, and then, two numerical models were designed using Modflow 2000. Groundwater flow was modelled with Modflow 2000 using the finite difference method, which uses cell-centred saturated flow. The finite difference method (FDM) is a method used to solve differential equations that are difficult or impossible to solve analytically. In both cases, steady-state models were created. The units describing the flow between the cells are L for unit length and T for unit time. The groundwater flow in the model was calculated using Darcy's law.

The three-dimensional movement of groundwater with constant density through porous earth material may be described by the partial-differential equation:

$$\frac{\partial}{\partial x} \left( K_{xx} \frac{\partial h}{\partial x} \right) + \frac{\partial}{\partial y} \left( K_{yy} \frac{\partial h}{\partial y} \right) + \frac{\partial}{\partial z} \left( K_{zz} \frac{\partial h}{\partial z} \right) + q = S_s \frac{\partial h}{\partial t} \quad (1)$$

where:  $K_{xx}$ ,  $K_{yy}$ ,  $K_{zz}$ —values of hydraulic conductivity along x, y, and z coordinate axes, which are assumed to be parallel to the main axes of hydraulic conductivity,  $h$ —hydraulic head [L],  $q$ —volumetric flux per unit volume representing sources and sinks of water, with  $q < 0$  for the flow out of the groundwater system, and  $q > 0$  for the flow into the system [T<sup>-1</sup>],  $S_s$ —is the specific storage of the porous media [L<sup>-1</sup>],  $t$ —time [T].

In general,  $S_s$ ,  $K_{xx}$ ,  $K_{yy}$  and  $K_{zz}$  are functions of space, and  $q$  is a function of space and time. Equation (1) describes groundwater flow under non-equilibrium conditions in a heterogeneous and anisotropic medium. The main axes of hydraulic conductivity are aligned with the coordinate directions. Unless the system is very simple, an analytical solution is possible, but usually numerical methods must be used to solve this equation. One approach is the finite difference method, in which the space is described with elements having predetermined head values at the centre of each element. The solution provides the head values at specific points and times [40].

In our model, we used the PEST tool to calibrate hydraulic conductivity and hydraulic heads between the measured and modelled data (hydraulic conductivity with a range from  $k = 0.01$  m/s to  $k = 1 \times 10^{-9}$  m/s and recharge from precipitation). The purpose of PEST (parameter estimation) is to assist in data interpretation, model calibration, and predictive analysis.

## 2.6. Validation of the Modelled Data

The Nash–Sutcliffe [41] efficiency coefficient was used to validate the modelled data. It indicates how well the measured and the simulated data match. The range of the efficiency coefficient is between minus infinity, which corresponds to a very poor model, and 1. The closer the coefficient is to 1, the better the modelled results. The following equation is used:

$$CE = 1 - \frac{\sum_{i=0}^n (y_i - y_{SIM,i})^2}{\sum_{i=0}^n (y_i - \bar{y})^2} \quad (2)$$

where:

$CE$ —coefficient of efficiency [/];

$y_i$ —measured parameter [ $m^3/s$ ];

$y_{SIM,i}$ —modelled parameter [ $m^3/s$ ];

$\bar{y}$ —mean measured parameter [ $m^3/s$ ].

The calibration is considered successful if the error between the modelled and the measured data is minimal.

MODPATH was used as a model validation method of hydraulic conductivity with particle tracking simulated using MODPATH [39]. After successfully calibrating the model using PEST, we performed validation of travel times within groundwater protection areas (VVOs). If particles travel from the source to the edge of the protected area within the time frame specified in the regulation, the model has been successfully calibrated. In Slovenia, VVOs are subdivided according to the type of aquifer: a karst aquifer or a porous aquifer. The travel times of pollutants are shorter in the case of a karst aquifer. In our study, we performed validation for porous aquifers. Our aim was to ensure sustainable use of fresh drinking water in the protected areas of the *Dravsko polje* and *Ptujsko polje* groundwater pumping stations, and prevent pollution of this valuable source. In addition to being the most vulnerable direct freshwater source, and; as such; subject to the strictest restrictions, the area around the source is divided into sections that are protected according to their respective distance from the source. These areas are divided into the narrowest inner area (VVO I), the narrowest area (VVO II) and the widest area (VVO III). The strictest measures apply to the inner area, followed by strict measures and finally milder measures for the outer area. The measures for each protected area are based on the travel time (up to 50 days in VVO I and up to 400 days in VVO II) of the pollutant from the intake to the source.

## 2.7. Uncertainties and Limitations

Numerical groundwater simulations play an important role in the description of dynamic groundwater processes and quantitative assessments of groundwater resources [42]. However, the reliability of numerical groundwater simulations is confounded by uncertainties associated with hydrogeologic structure, spatial local-scale heterogeneity, and parameter values of modelled geological materials. Uncertainties related to hydrogeological structure are often of greatest concern [43]. Hydrogeological structure mainly refers to the spatial distribution of the spatial characteristics of aquifers and aquitards [44]. A proper description or model of their hydrogeological structure is required to accurately resolve uncertainties and effectively evaluate groundwater resources.

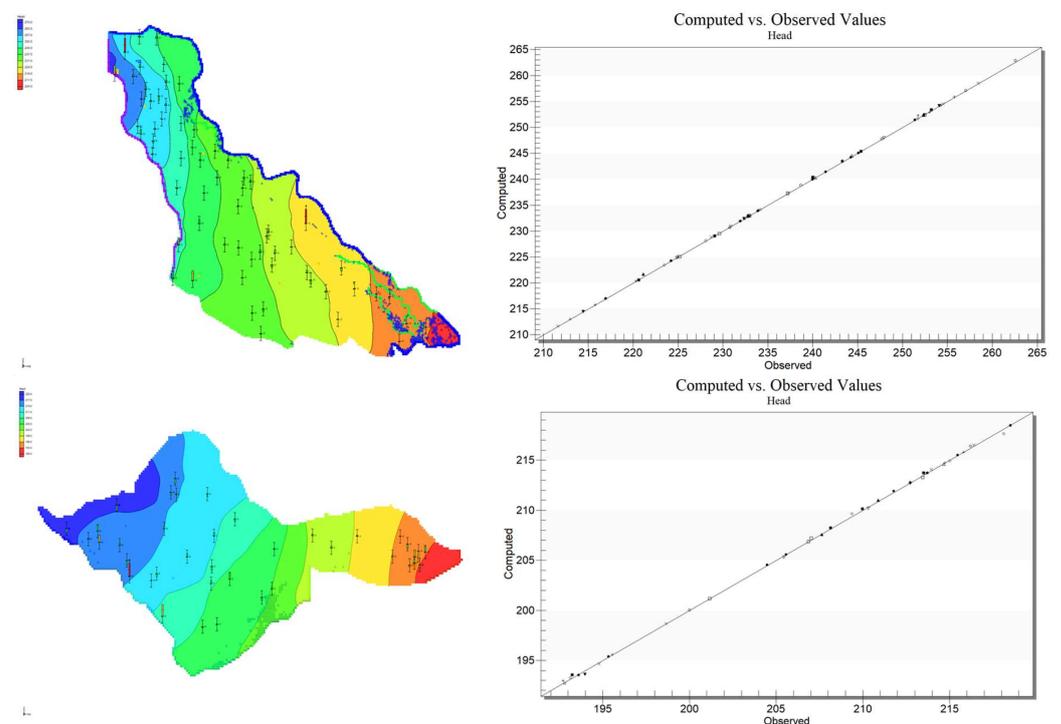
While the water balance of the presented porous aquifer system is well understood and previously well researched, a window of data is unpredictable and, therefore, remains unknown. For the purposes of the targeted model, area simplifications had to be applied, which proved sufficient to create a steady-state model of the system. For instance, the amount of precipitation is influenced by numerous factors, such as altitude, outskirts, natural barriers, local climate changes, etc., which have a major impact on groundwater recharge. Another variable is the recharge of groundwater from the outskirts, which is not regularly measured and needed to be inferred from data estimates from more than a decade ago. The same was true for GWLs, which had been measured in different seasons over the

years, and each measurement also required taking into account the weather conditions to achieve the best accuracy. Another factor impossible to measure is whether the riverbed is recharged by groundwater or the aquifer is recharged by a riverbed leakage. Generally, both scenarios are known to be valid, but it is impossible to determine them locally on such a scale. In addition, groundwater flow is generally known, but local changes in flow remain unknown and need to be explored locally for the needs of targeted studies. Therefore, this portion of the data must also be viewed with caution. However, to create a genuine and precise simultaneous model of a large spatial reach, it seems more appropriate to model smaller parts targeting local problems and gradually combine them to form the entire aquifer system.

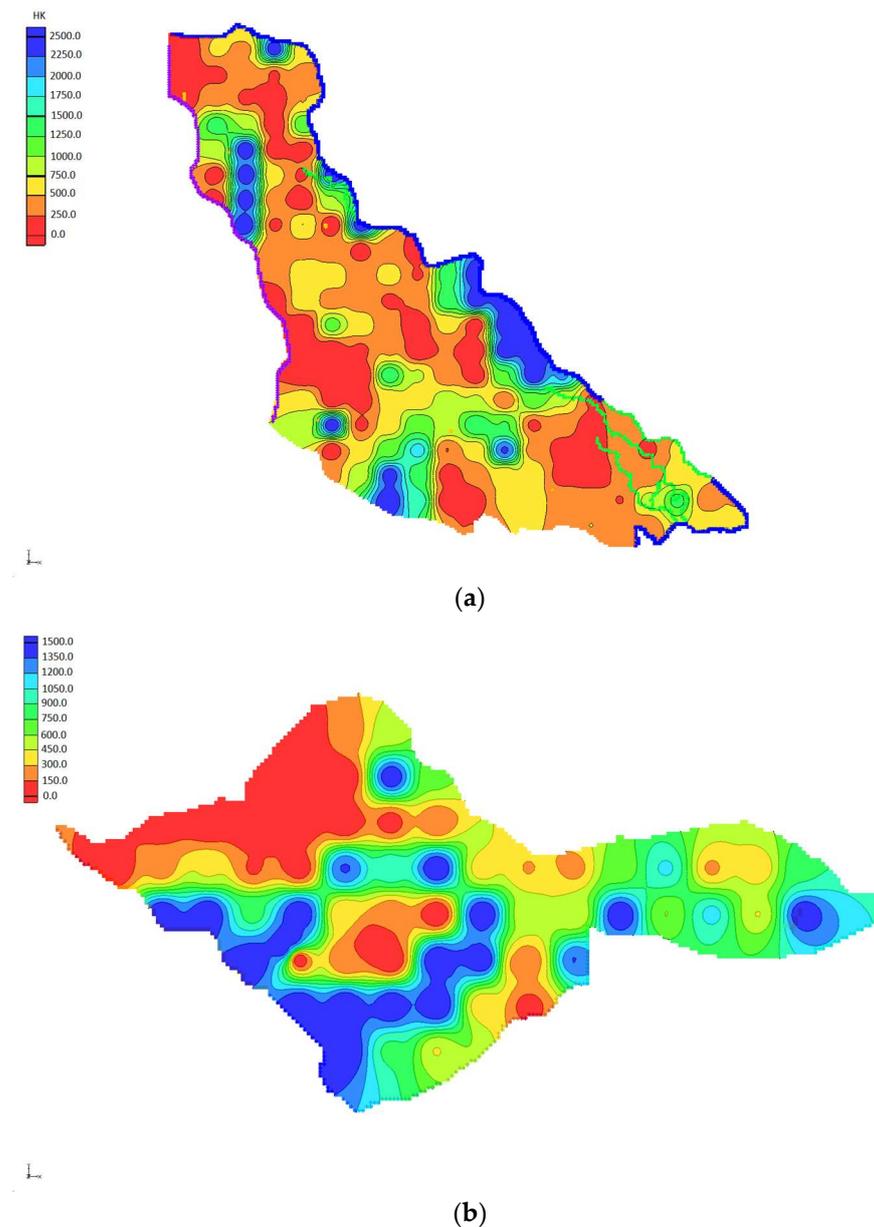
### 3. Results

#### 3.1. Groundwater Hydraulic Heads

The steady-state model of the *Dravsko polje* aquifer is presented in Figure 8. A high water level is characteristic for this time of the year. The steady-state model of the *Ptujsko polje* aquifer is presented in Figure 8. Intermediate water level is characteristic for this time of the year. As expected, the calculated flow direction was W-E in the northern part and slightly SE in the southern part. The modelled hydraulic heads were consistent with the expectations of the observed hydraulic heads from the conceptual model. The model of the *Ptujsko polje* aquifer showed the expected development of the calculated hydraulic heads in the flow direction (Figure 9); in the western part, the groundwater flow direction was NW to SE and in the eastern part, parallel to the Drava River, the groundwater flow direction was W-E. Hydraulic head alignment between the computed and the observed values is shown in Figure 8 for *Dravsko polje* and Figure 8 for *Ptujsko polje*.



**Figure 8.** Modelled hydraulic heads in the *Dravsko polje* aquifer; left—model results with observation points in relation to modelled data; right—Comparison between computed and observed hydraulic heads on the *Dravsko polje* aquifer (above). Modelled hydraulic head in the *Ptujsko polje* aquifer; left—model results with observation points in relation to modelled data, where the deviation is less than 0.25 m, marked with a green line, right—deviation between computed vs. observed values in the *Ptujsko polje* aquifer (below).



**Figure 9.** (a) Hydraulic conductivity in the *Dravsko polje* aquifer; (b) hydraulic conductivity in the *Ptujsko polje* aquifer.

### 3.2. Hydraulic Conductivity of Aquifers

The modelled hydraulic coefficient of the *Dravsko polje* aquifer ranged from 0.1 m/day to 2500 m/day. Values between 500 and 250 m/day predominate, which is a very good result based on the measured values of hydraulic conductivity [26]. The modelled hydraulic conductivity in the *Ptujsko polje* aquifer ranged from 0.8 to 1500 m/day. Hydraulic conductivity was calculated on the basis of the measured GWL data. The distribution of hydraulic conductivity is shown in Figure 9.

### 3.3. Flow Budget

As shown in Table 5, the results of the calculated flow budget were well estimated. The total inflow and outflow of the *Dravsko polje* aquifer is higher than predicted, but still sufficient. Most of the groundwater flows into the Drava River and the rest into drains. The calculated discharge into Lake Ptuj was lower than the estimated discharge, which was 50.000 m<sup>3</sup>/s; the calculated discharge was 13.632 m<sup>3</sup>/s. The results of the modelled flow

budget for the *Ptujsko polje* aquifer were very close to the estimated discharge. The result was considered to be very good. Most of the groundwater flows into the Drava River, and the rest migrates towards the drains.

**Table 5.** Model flow budget results for the *Dravsko polje* and *Ptujsko polje* aquifers.

	Dravsko Polje		Ptujsko Polje	
	Flow In (m <sup>3</sup> /day)	Flow Out (m <sup>3</sup> /day)	Flow In (m <sup>3</sup> /day)	Flow Out (m <sup>3</sup> /day)
Wells		528.5		2.951
Infiltration filed			1228	
Drains		42.984		40.065
Lake Ptuj		13.632	9.765	
River leakage	105.479	482.762	7.457	51.619
Constant head (Pohorje)	299.119	27.617		
Recharge (Precipitation)	162.997		76.205	
Total in-out	567.595	567.539	94.656	94.635

Along the western edge of the *Dravsko polje* aquifer, the model of the general head boundary is divided into three hydrogeological sections: 1, 2, and 3, as described by Žlebnič [26]. Table 6 shows the flow rates for all three sections.

**Table 6.** Flow rate for the western margin of the *Dravsko polje* aquifer.

Constant Head Section	Flow Rate (m <sup>3</sup> /day)
1 section	101.631
2 section	87.947
3 section	81.922

### 3.4. Interaction between the Aquifers

In the place where the *Dravsko polje* aquifer is connected to the *Ptujsko polje* aquifer, the discharge from Lake Ptuj should be unified. Therefore, it is important to obtain similar values for the two areas. The values for the river should vary, as the conditions on the left and right banks of the river are not necessarily the same. From the *Dravsko polje* aquifer, the lake discharges approximately 14.000 m<sup>3</sup>/day. The river discharges 18.000 m<sup>3</sup>/day. The inflow from Lake Ptuj into the *Ptujsko polje* aquifer is calculated to be about 10.000 m<sup>3</sup>/day, and so, some discrepancy between the inflow and outflow budgets was acknowledged, but it was still acceptable. However, the river leakage was neglected due to the unknown natural interaction with the groundwater and the dimensions of the river bank.

The flow budget with the intersecting part between the *Ptujsko polje* aquifer and the *Dravsko polje* aquifer is presented in Table 7.

**Table 7.** Flow budget result for the intersecting part *Ptujsko–Dravsko polje* aquifer.

	Ptujsko Polje		Dravsko Polje	
	Flow In	Flow Out	Flow In	Flow Out
Lake (m <sup>3</sup> /day)	9.765	0		13.632
River (m <sup>3</sup> /day)	4.095	14.525	2.982	17.492
Total in-out	13.860	14.525	2.982	31.124

### 3.5. Validation of Models

Finally, the data were estimated using the efficiency coefficient by Nash & Sutcliffe (NSE). According to Nash–Sutcliffe’s efficiency coefficient, both models were well calibrated, with a result NSE of 1 (Table 8).

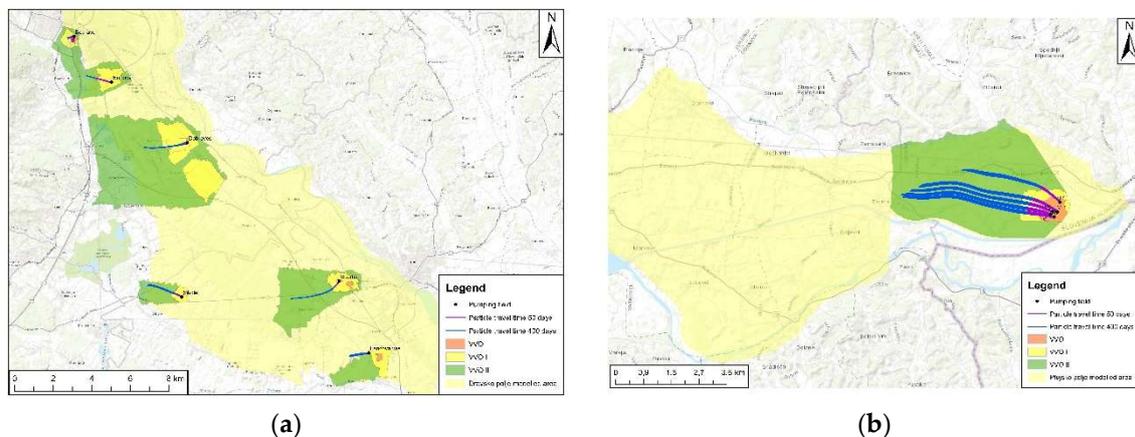
**Table 8.** Nash–Sutcliffe coefficient of efficiency.

	Dravsko Polje	Ptujsko Polje
$(y_i - y_{SIM,i})^2$	1.18	1.32
$(y_i - \bar{y})^2$	9991.77	2806.61
$\bar{y}$	237.53	206.91
NSE	0.99	0.99

### 3.6. Modelled Data Compared to Travel Times after PEST Calibration

The model results of the *Dravsko polje* and *Ptujsko polje* aquifers were also estimated using the MODPATH model, whereby the validation of hydraulic conductivity was performed. We simulated the particle pathways from the pumping source. The following pumping sites are located along the *Dravsko polje* aquifer sources: Lancova vas, Skorba, Šikole, Dravski dvor, Dobrovce, Bohova, and Betnava. In *Ptujsko polje*, there are pumping sites in Sejanica pri Mihovcih with several pumping wells. For a better visual representation, we used five wells out of seventeen.

We simulated two travel times, which also coincided with our water protection zones (VVO). The first was up to 50 days and the second was up to 400 days. Both results are shown in Figure 10. Based on the results of both travel times, both models were shown to be well calibrated. The travel time frames for 50 (purple line) and 400 (blue line) days both achieve the results within the sections of the VVO. Based on the PEST calibration, the particle travel times were shown to be well correlated. This indicates that the hydraulic conductivity was well calibrated and the effective porosity of the model was realistic.



**Figure 10.** (a) Modelled travel times for the *Dravsko polje* aquifer; (b) Modelled travel times for the *Ptujsko polje* aquifer.

## 4. Discussion and Conclusions

Sustainable management of hydrological resources is regulated by the European Water Directive (60/2000/EC), which requires each water authority in the EU to prepare a watershed management plan. It is noted that to fulfil these obligations, it is necessary to use integrated management tools and methods, such as computer models, while also taking into account the complexity of water bodies monitoring and uncertain future scenarios brought about by climate change.

Modelling is important to improve our understanding of hydrogeological systems and to assess future conditions in aquifers. Integrated hydrogeological models represent a valuable tool not only in the planning, but also in the implementation and operation of activities related to groundwater management, land use, and water protection action plans. A model provides a description of a specific hydrogeological system. Data availability is extremely important when developing models. For better estimates of the key parameters,

field measurements should be conducted. To assess the degree of validity of the model outputs, uncertainty analysis should also be carried out.

This article presents the first results of two separate hydrogeological numerical models for the *Dravsko polje* and *Ptujsko polje* aquifers, which were combined into an integrated groundwater resource management system. The combination of the two models allowed us to follow and optimise the same approach in groundwater management in this area. It is a tool that helps solve potential problems related to the quantity and quality of groundwater. The models were constructed using different grid-based maps: GWL, lower and upper aquifer layers. Both models were at steady state.

Based on the results, we successfully calibrated both models. From the calibrated model for the *Dravsko polje* and *Ptujsko polje* aquifers, it appeared that the hydraulic heads for both aquifers were well coordinated, and most were within a 25% discrepancy. There are larger discrepancies in hydraulic heads at the western margins of the aquifers, which is to be expected since the aquifers are highly variable. The western margin of the *Dravsko polje* aquifer presented a major modelling challenge. Due to the very high elevation of the tertiary layer and the resulting low thickness of the aquifer (less than 0.5 m), the high GWL, the constant pumping rate, and the unknown recharge rate from the surrounding area, conditions were not ideal. In addition, under such unpredictable conditions, it was very challenging for PEST to calibrate the type of groundwater flow and meet all the given conditions: hydraulic conductivity, GWL within the aquifer thickness, recharge rate, and constant pumping rate.

To achieve the best possible approximation of the model to the natural conditions, it is imperative to continuously measure GWL, infiltration by precipitation, and recharge from the surrounding area, and to update the model regularly. Only then can the model represent a good approximation of the natural conditions. The comparison between the measured and the modelled data shows a very good alignment for both models and indicates that both numerical models provide realistic results of hydraulic pressure distribution and groundwater flow direction.

In addition, the overall water budget of the *Dravsko polje* aquifer and the *Ptujsko polje* aquifer showed only minimal deviation between inflow and outflow, indicating that both models were stable. The modelled water balance data and the estimated water balance data derived from the literature showed that the water balance estimate was consistent with the modelled data. This indicates that the understanding of the groundwater system was correct from the beginning, which facilitated the attainment of correct model results.

Finally, both models were compared to particle travel times from the source of the groundwater protection areas, using the MODPATH procedure for validation. Based on the data presented, the particle travel times were very well aligned. This indicates that the hydraulic conductivity was well calibrated, and the effective porosity was realistic.

However, the models were created separately as this made them easier to present, and the modelled areas were smaller, which gave more control to the BCs. Therefore, we had to unify the intersecting part of the two models. They were connected to the Drava River and Lake Ptuj. Based on the flow budget of the *Drava polje* aquifer, the total discharge below Lake Ptuj and the Drava River was  $31.124 \text{ m}^3/\text{s}$ , and this amount of groundwater was expected to continue to flow into the *Ptujsko polje* aquifer. The modelled flow budget of the *Ptujsko polje* aquifer showed that the inflow at its western edge totalled  $13.860 \text{ m}^3/\text{s}$ , which was less than the amount from the *Dravsko polje* aquifer. Based on these volumes, more than half of the groundwater volume from the *Dravsko polje* aquifer disappeared. However, since the Drava River flows downstream, it is possible that part of the groundwater does not infiltrate into the *Ptujsko polje* aquifer, but enriches the Drava River. It should be kept in mind that Lake Ptuj is artificially regulated, which is another factor for possible discrepancies and a higher probability of errors.

Practice has shown that it is difficult to predict all input values to perfectly match the natural system. However, with such models based on the mathematical background of

cell-to-cell flow rate calculations, we can predict behaviour similar to the natural system, which is exactly the case in our example.

In summary, both models were successfully modelled and validated. Not only did these two models provide a better understanding of groundwater flow in the Drava River water body, but they can also be used to model pollution distribution and infiltration through the saturated zone of the aquifer. The hydraulic model will facilitate the planning of measures for chemical and quantitative improvement of the aquifer. According to the European Water Framework Directive, the groundwater body of the *Dravska kotlina* is classified as a groundwater body with poor chemical and quantitative status. For this reason, measures must be taken to restore its chemical and quantitative status. The hydraulic model of the aquifer is, therefore, essential both for a more detailed definition of the current situation and for predicting the impact of future measures aimed at improving the situation. In addition, these models allow for more accurate simulations of groundwater dynamics and pollutant migration in aquifers, which, in turn, can contribute to more efficient management of groundwater resources. Last but not least, both models together form the basis for a complete decision support system for the management of the entire groundwater body.

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**Data Availability Statement:** The data presented in this study are available on request from Geological Survey of Slovenia research databases.

**Conflicts of Interest:** The authors declare no conflict of interest.

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