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Research on Hydrolithospheric Processes Using the Results of Groundwater Inflow Testing

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Abstract: The article considers a mathematical model of the hydrolithospheric process taking into account the skin effect. A methodology for using the results of groundwater inflow testing to determine the parameters of approximating models that take into account skin effects is presented. In addition, the problems of modeling hydrodynamic processes taking into account random factors are considered. A statistical analysis of well monitoring data was carried out and an algorithm for studying processes was developed. Using the obtained approximating models, a procedure for solving the problem of selecting the optimal number of production wells has been developed. Based on the results of the groundwater inflow testing, the prospects for the development and use of new aquifers can be determined.

Keywords: hydrolithospheric processes; approximating models; skin effect; mathematical models; optimization of the number of wells



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1. Introduction

Currently, billions of people directly depend on groundwater resources. At a time when anthropogenic factors exert a significant influence on the environmental aspect of this crucial resource, the issue of the ecological safety of mineral and groundwater is becoming a common challenge for humanity as a whole. When examining the extraction of mineral water on a global scale, several key problems can be identified at present, arising in the development of deposits [1–3].

The first problem, which is at the same time the most important and can also entail very serious consequences, is insufficient study of the area in which the deposit will be developed [4,5]. Mineral water extraction requires a thorough study of the area before development, as well as consideration of many aspects related to the mining conditions of specific deposits. It is also crucial to fully study the area in which mining will be carried out, at all its ecological levels, since the development of a mineral water deposit is a complex process that can affect all natural and environmental aspects of the surrounding area [6–8].

Before starting the development of a mineral water deposit, it is essential to consider all the physical and geological aspects of the deposit itself, its distance from populated areas, the intensity of production, the physical and geological characteristics of the surrounding soils, as well as all the environmental conditions of the area [9–11]. Additionally, it is critical to draw up a plan for the development of the deposit, and not to neglect the recommended conditions for the technological process of developing a mineral water deposit.

Often, neglecting to properly study the area and aspects of mining can lead to dire consequences for the ecology of the surrounding area, as well as for people living in surrounding settlements. For example, insufficient study of aspects of production can lead to multiple violations in field development technology, leading to insufficient or too intensive production, and the consequences of this can include violations of the ecological

balance of the production region, geological disasters, mudflows, landslides, depletion of the region's aquifer, and many others [12–14]. Also, among the global problems, we must highlight the environmental problem of the anthropogenic impact on the quality of underground mineral waters.

In the 19th and 20th centuries, due to globalization and industrial revolutions, the active growth of settlements, factories, and enterprises began. The main substances that these objects release, particularly into the soil, are, for example, petroleum products, zinc or copper compounds, mercury, phenols, and so on—all of these are the main substances that pollute underground water sources. Consequently, the main reason for the decline in water quality and the deterioration of its physical and chemical composition has come to be human activity, that is, our environmentally damaging behavior and environmental pollution.

In general, the state of water resources directly depends on the state of the ecological environment of the region, its biosphere and chemical processes, and hydrogeological conditions. Today, humans' influence on groundwater continues to increase. Our massive development has an active influence on the upper layers of soil, and it is there that the main potable water reserves are located. In addition, they are most susceptible to pollution due to the relatively low distance to the upper soil layer and low filtration. This contributes to the very slow self-purification of such waters. Furthermore, it transpires that after underground waters, pollution travels no further, meaning aquifers become a kind of final reservoir for pollution. It can also be noted that due to the same mass development and anthropogenic influence, the load on the reservoir increases greatly, which leads to the migration of water into deeper or neighboring layers, while continuing to be polluted [15–18].

In discussing mineral water deposits, it is essential to take into account their spatial dimensions. This necessitates conceptualizing the management subject as a system characterized by distributed parameters. Moreover, when examining this system, the impact of random factors on hydrogeological processes should not be overlooked.

Comprehensive details on the construction of mathematical models for hydrological processes are extensively presented in the studies by E.A. Lomakin, I.K. Gavich, P.K. Konosavsky, and V.V. Antonov, among others [18–20]. These researchers have significantly advanced the development of methodologies for creating two-dimensional models. However, their methodologies are not directly applicable to developing three-dimensional models of spatially distributed objects. The analysis of distributed systems, particularly in the frequency domain, was explored in [21–23]. Recent research in control systems with distributed parameters, particularly the works of Director JSC “Narzan” and A.V. Malkov, have been pivotal in refining theoretical approaches to controlling hydrogeological entities [2,24,25]. In the realm of stochastic systems, meanwhile, there are numerous contributions to date [26–28]. However, these studies, whether focusing on concentrated or distributed systems in the control field, do not delve into the analysis of random processes.

Therefore, the primary objectives of this study were enhancing the economic efficiency of mineral water extraction while minimizing environmental threats. With those in mind, we set out to analyze hydrogeological processes under random influences, find methods and algorithms for management, and identify laws for optimal functioning of the technological process [29–31].

To achieve the above objectives, the study of hydrodynamic processes in aquifers and the construction of mathematical models for the rational use of water resources is a priority, and the specific direction we took with this study was to investigate the control system of hydrolithospheric processes, aimed at preserving the quality and reserves of the mineral water in the deposit by controlling the piezometric level in the exploited aquifer. To that end, it was necessary to conduct a detailed analysis of the hydrogeological object, examining the inherent random processes that affect the hydrodynamic parameters during the exploitation of the deposit [32–35]. With that purpose, systematic analytical modeling of the technical and physical elements of the research object at different management levels was required, followed by the development of a mathematical model that accounts for random disturbances [2,36,37]. The object of study was an exploited site in the Caucasian Mineral Waters region.

2. Materials and Methods

Observational data from monitoring wells indicate the consistency of oscillations in pressure. These fluctuations might stem from various sources, including disruptions in the technological process due to equipment malfunction, or changes in aquifer flow rates influenced by weather and human activities. The human factor is significant; for instance, ongoing urban development over the past two decades in the extraction area under study has impacted soil settlement and altered impermeable layer structures. Moreover, in areas lacking centralized sewage systems, septic tank pollutants seep into the groundwater, exacerbated by rainfall. While such issues might not significantly impact the mineral water quality under high aquifer pressure, unregulated water extraction causes pressure drops, leading to the infiltration of impurities into the exploited layers and aggravating environmental conditions [38–41].

Sometimes, when influenced by a range of factors, the observed pressure variations can only be interpreted as uncertain and random processes. These should be factored into the development of mathematical models for hydrolithospheric extraction processes. Additionally, seismic activities, with their unpredictable or random impacts on the deposit, are also worth considering. The hypothesized random factors are thought to affect water flow speeds in the aquifer and the rate of pressure decrease. When analyzing pressure trends in monitoring wells, it can be inferred that the model's limitations include taking these random fluctuations to have a stationary nature [42–44]. In this study, a statistical analysis of piezometric level data from the monitoring wells in the specific deposit area was conducted to verify this hypothesis [45–48].

2.1. Statistical Analysis

Often, examining the graphical representation of a process is adequate to identify non-stationarity [18–25] in a series. However, to provide statistically robust evidence of stationarity, specific techniques are necessary. Here, methods like constructing the process's correlogram and segmenting it into sequential observation groups were chosen for comparative analysis [49–51].

Our analysis of the piezometric level data series from the deposit's monitoring well, employing the first method and using a correlogram, is depicted in Figure 1. The fact that the first-order autocorrelation coefficient is the highest suggests the presence of a trend in the time series.

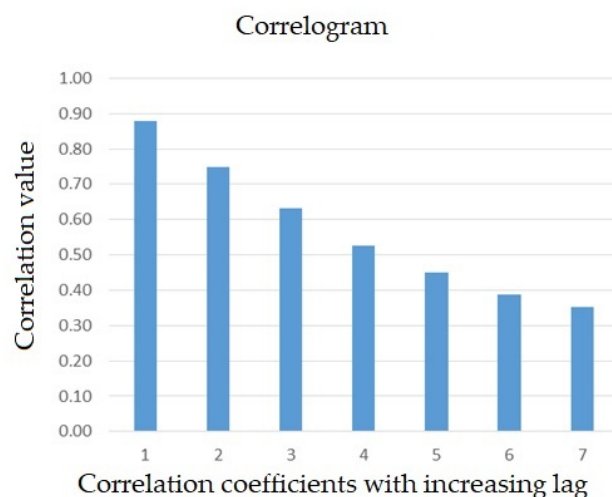


Figure 1. Correlogram showing the correlation coefficients between the original series of observations and the series shifted by time steps (from 1 to 7) [51].

The correlation decreases steadily with increasing lag, pointing to the series' stationarity, indicating that its probabilistic characteristics remain constant over time. To validate these findings, we examined the data using a secondary method for determining stationarity [2,52–54].

The results of previous studies were summarized and systematized in the form of an algorithm that can be used when analyzing monitoring data from sensors at observation wells of a hydrogeological object in order to assess the stationarity of processes. The analysis technique is divided into several stages, some of which can be repeated until the desired results are obtained:

- (1) Monitoring the dynamics of the piezometric level in observation wells. Obtaining time series from sensors;
- (2) Estimating the stationarity of a series using a correlogram:
 - (2.1) Calculating the correlation coefficient for the original series versus series with a lag (1st, 2nd, 3rd order, etc.);
 - (2.2) Constructing a histogram using autocorrelation. Obtaining a correlogram;
 - (2.3) Then, performing correlogram analysis: if the correlogram fades and tends to zero, then we accept the hypothesis that the series is stationary; otherwise, the hypothesis of stationarity is not accepted;
- (3) Estimating the stationarity of a series by using the method of sequential grouping:
 - (3.1) Dividing a series into successive groups with an equal number of observations;
 - (3.2) Calculating the means and variances of the resulting groups;
 - (3.3) Comparing the means using Student's t -test: if $t_{\text{calc}} < t_{\text{table}}$, then the hypothesis of stationarity is accepted; if $t_{\text{calc}} > t_{\text{table}}$, then the stationarity hypothesis is not accepted;
- (4) Removing trends:
 - (4.1) Taking the differences in two consecutive values of the time series Δ ;
 - (4.1) The resulting series of residues is also analyzed using the t -test: if $t_{\text{calc}} < t_{\text{table}}$, then the stationarity hypothesis is accepted; if $t_{\text{calc}} > t_{\text{table}}$, then the hypothesis of stationarity is not accepted—it is necessary to repeat point 4.1 (taking differences in the next order until the series comes to a stationary form).

This technique may be used to model random impacts on a hydrogeological object.

2.2. Mathematical Models of Aquifers

Based on the findings of our studies of the hydrogeological object, a mathematical model of the deposit was compiled, as shown in Figure 2. Extraction of hydromineral raw materials is carried out using one production well. To monitor the state of hydrolithospheric processes, two monitoring wells were installed.

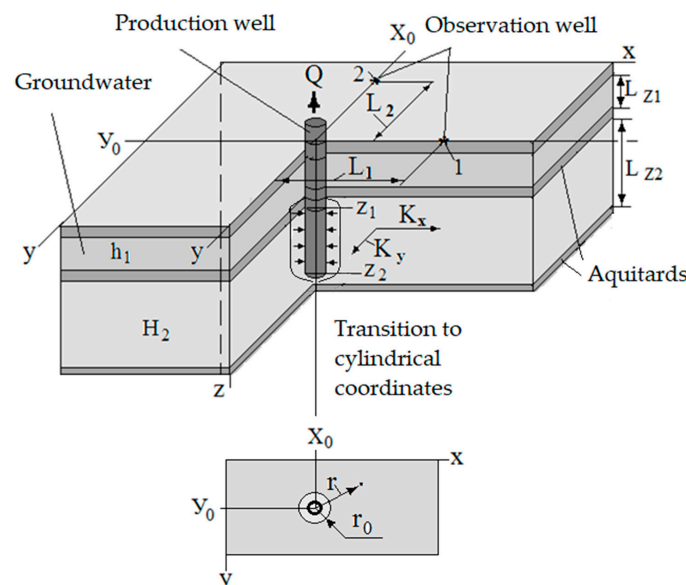


Figure 2. Object of study.

The groundwater inflow testing carried out showed that a skin effect occurs around the production wells (the fine fraction of the rock is carried out along with the extracted raw materials). The hydraulic radius of the well with a negative skin effect increases, and with a positive one, it decreases. Figure 3 shows the formation of the hydraulic radius of the well ("well") and depression funnel.

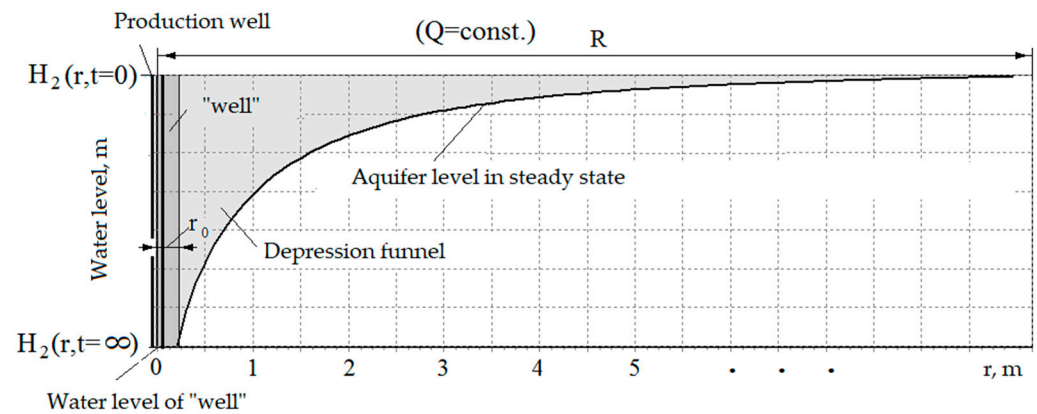


Figure 3. Formation of a depression funnel.

Taking into account the formation of the hydraulic radius of the well and funnel, the mathematical model of hydrodynamic processes should be written in Cartesian and radial coordinates [55–58].

A description of the process in groundwater is as follows:

$$\frac{\partial h_1(r, z, \tau)}{\partial \tau} = k_1 \frac{\partial^2 h_1(r, z, \tau)}{\partial r^2} + \frac{k_1}{r} \cdot \frac{\partial h_1(r, z, \tau)}{\partial r} + k_{1,z} \frac{\partial^2 h_1(r, z, \tau)}{\partial z_1^2}; \quad (1)$$

$$0 < r < R; 0 < z < L_{z1}.$$

A description of the process in the aquifer is as follows:

$$\frac{\partial H_{2k}(\tau)}{\partial \tau} = 1/\mu \cdot (-Q + \Delta V_k) / (\pi \cdot r_0^2),$$

$$\Delta V_k = 2\pi \cdot r_0 \cdot \int_{L_{z1}}^{L_{z2}} (k_{2r} \cdot (H_2(r_0, z, \tau) - H_{2k})) \cdot dz, \quad (2)$$

$$0 < r < r_0; L_{z1} < z < L_{z2};$$

$$\frac{\partial H_2(r, z, \tau)}{\partial \tau} = \frac{1}{\eta_2} \left(k_2 \frac{\partial^2 H_2(r, z, \tau)}{\partial r^2} + \frac{k_2}{r} \cdot \frac{\partial H_2(r, z, \tau)}{\partial r} + k_{2,z} \frac{\partial^2 H_2(r, z, \tau)}{\partial z_2^2} \right), \quad (3)$$

$$r_0 < r < R; L_{z1} < z < L_{z2}.$$

Boundary conditions are described in the form of Darcy conditions. The boundary between the groundwater layer and the exploited aquifer is as follows:

$$h_1(r, L_{z1}, \tau) = h_1(r, L_{z1}, \tau) + b_1 \cdot (H_{2k}(r, 0, \tau) - h_1(r, L_{z1}, \tau)) \cdot \partial \tau,$$

$$H_{2k}(r, 0, \tau) = H_{2k}(r, 0, \tau) - b_1 \cdot (H_{2k}(r, 0, \tau) - h_1(r, L_{z1}, \tau)) \cdot \partial \tau, \quad (4)$$

$$0 < r < r_0;$$

$$H_{2,k}(r_0, z, \tau) = H_{2k}(r_0, z, \tau) + k_{2z} \cdot (H_2(r_0, z, \tau) - H_{2k}(r_0, z, \tau)) \cdot \partial \tau,$$

$$H_2(r_0, z, \tau) = H_2(r_0, z, \tau) - k_{2z} \cdot (H_2(r_0, z, \tau) - H_{2k}(r_0, z, \tau)) \cdot \partial \tau, \quad (5)$$

$$\partial H_{2k}(0, z, \tau) / \partial r = 0,$$

$$0 < z < L_{z2};$$

$$\partial H_{2k}(r, L_{z2}, \tau) / \partial z = 0, \quad 0 < r < r_0; \quad (6)$$

$$h_1(r, L_{z1}, \tau) = h_1(r, L_{z1}, \tau) + b_1 \cdot (H_2(r, 0, \tau) - h_1(r, L_{z1}, \tau)) \cdot \partial \tau,$$

$$H_2(r, 0, \tau) = H_2(r, 0, \tau) - b_1 \cdot (H_2(r, 0, \tau) - h_1(r, L_{z1}, \tau)) \cdot \partial \tau, \quad (7)$$

$$r_0 < r < R;$$

$$\partial H_2(r, L_{z2}, \tau) / \partial z = 0, \quad r_0 < r < R; \quad (8)$$

$$h_1(R, z, \tau) = h_{1,0}; \quad 0 < z < L_{z1}, \quad H_2(R, z, \tau) = H_{2,0}; \quad 0 < z < L_{z2}. \quad (9)$$

Here, b_1 —flow coefficient 1/day; h_1 —pressure in the groundwater, m; H_2 , H_{2k} —pressure in the exploited aquifer, m; k_1 , $k_{1,z}$ —parameters of groundwater conductivity level along coordinate axes, m^2/day ; k_2 , $k_{2,z}$ —filtration parameters of the exploited formation along the coordinate axes, m/day ; $k_{2,r}$ —flow parameter on the side surface of the well, m/day ; η_2 —elastic capacity of the aquifer, 1/m; μ —water loss coefficient, 1/m; r —radial coordinate, m (radius of the cylindrical coordinate system); z —Cartesian coordinate, m (height of the cylindrical coordinate system); L_{z1} , L_{z2} , R —values of the height of the layers and the radius of the depression funnel, m; τ —simulation time; $h_{1,0}$, $H_{2,0}$ —initial pressure levels in the aquifers, m; and Q —flow rate, m^3/day .

The developed mathematical model (1–9) cannot be solved by means of analytical methods. The studies given in [1] show that for a given flow rate of a production well, with an increase in the well radius r_0 (Figure 2), the influence of flow rate on the level of the aquifer decreases.

The mathematical model was verified using field experiments at the mineral water deposit. When simulating the discrete model, it was found that with a flow rate $Q = 100 m^3/day$ (in Equation (2)), the decrease in water level at the location of the production well will be 2.55 m. Then, during the field experiments for 10 days with an average flow rate $Q = 100 m^3/day$ on the real deposit, the decrease in water level was 2.57 m. The model values and natural values of the water level decrease were close, which indicated the adequacy of the use of the mathematical model.

2.3. Problems of Modeling the Hydrolithospheric Processes under Consideration

The mathematical models discussed above do not have an analytical solution. In practice, discrete analogues of the equations under consideration are used to obtain numerical solutions. In this case, the following tasks are solved:

1. Assessing the possibility of using the deposit under consideration in practical activities (determining the optimal location and number of production wells, maximum flow rates, etc.);
2. Designing closed-loop control systems for production wells;
3. Forecasting the development of hydrolithospheric processes in the near and long term.

Real hydrolithospheric processes certainly differ from the processes described by mathematical models. When describing mathematical models, it is assuming the following:

1. Physical parameters do not depend on spatial coordinates. In reality, the parameters for different points in the spatial domain will differ;
2. The geometric parameters of the areas in which the processes under consideration occur change according to nonlinear laws (for example, the thickness of aquitards nonlinearly depends on spatial coordinates; therefore, the flow coefficients for different points will differ);
3. The geometric dimensions of the “wells” formed around production wells change during the operation of the wells.

In practice, solving problems of studying hydrolithospheric processes requires finite-dimensional approximation methods. These approximation methods should be used very carefully when studying systems with distributed parameters, because discretization of partial differential equations significantly [2] changes the properties of the model (a discontinuity in phase spaces occurs). That is, by developing a discrete model instead of a mathematical one, we produce a completely new object that differs in properties from the originally specified one. Research shows that when the discretization step size along Cartesian and radial coordinates changes, the characteristics of the discrete model also change.

In order to adjust the parameters of the developed field model, the results of groundwater inflow testing are used. When carrying out groundwater inflow testing, the physical

parameters of the hydrolithospheric process can be calculated [3], and the dynamic and static characteristics of the process can be determined.

Solutions to practical problems using only the results of the groundwater inflow testing need to be considered, along with the determination of the parameters of the link that approximates the statics and dynamics of the process under consideration.

The research results given above show that in order to determine the parameters of distributed approximating links [4], it is preferable to use the results of phase analysis carried out at the deposit under study. If such results are absent, then verified discrete models should be used that describe the hydrolithospheric processes of the aquifer in question.

Through groundwater inflow testing, for the deposit under consideration, it was found that at a production well flow rate of $100 \text{ m}^3/\text{day}$ [2,49,59], the value of the production well influence coefficient $K_1 = 0.0257$. Furthermore, the values of the coefficients of influence of the production well on the control wells are as follows: when the control well is removed from the production well at a distance $L_1 = L_x = 50 \text{ m}$, $K_2 = 0.00413$; when the control well is removed from the production well at a distance $L_2 = L_y = 60 \text{ m}$, $K_3 = 0.00237$.

To determine the dynamic characteristics of the process under consideration, a harmonic input action of the following form was applied to the input [60–64]:

$$Q(\tau) = Q + 0.3 \cdot Q \cdot \sin(\omega \cdot \tau), \quad \omega = 0.0005 \text{ 1/s}, \quad (10)$$

$$Q = 100 \text{ m}^3/\text{day} \text{ (or } Q = 100/(3600 \cdot 24) \text{ m}^3/\text{s}).$$

A graph of level changes in a production well (at the point where the sensor was installed at a depth of 75 m), with a harmonic input effect and over a time interval when the output function was close to a quasi-stationary mode, is shown in Figure 4. The value of the phase shift in this case indicates that at a high flow rate, the water level in the aquifer falls, and at a low flow rate, the water level in the reservoir increases [65–68]. Through calculations, the phase shift of the output signal relative to the input influence was obtained:

$$\Delta\varphi = -2\pi \cdot 0.738/3.625 = -1.279. \quad (11)$$

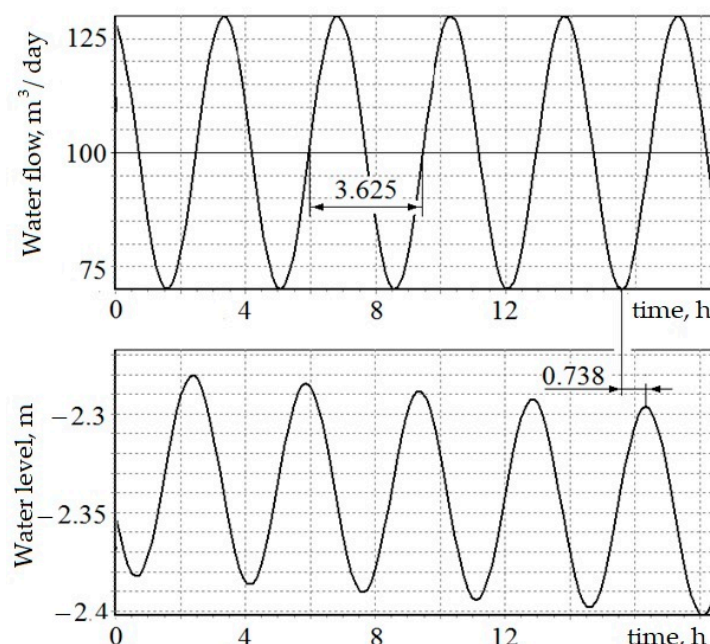


Figure 4. Change in level in a production well.

We experimented with a technique for determining the parameters of links that approximate the static and dynamic characteristics of the objects under consideration. In this case, we assumed that the filtering coefficients for spatial coordinates $\{x, y\}$ were

different. The structure of the link that approximates the static and dynamic parameters of the object was written in the following form:

$$W_a = \frac{K}{\beta} \cdot \exp(-\beta \cdot (x^2 + (k_n \cdot y)^2)^{0.5}); \beta = (s/a + D)^{\frac{1}{2}}. \quad (12)$$

Here, K , K_n , D , a —determined coefficients; s —Laplace operator; and x , y —spatial coordinates.

As is known, around the production well, as a result of the skin effect, a “well” is formed. A mathematical description of the processes involved is shown in (2). Let us assume that its radius is r_0 (Figure 2). Relation (12) is written in cylindrical coordinates:

$$\begin{aligned} r > r_0 &\rightarrow W_a = \frac{K}{\beta} \cdot \exp(-\beta \cdot r), \\ r &= ((x - X_0)^2 + K_n^2 \cdot (y - Y_0)^2)^{0.5}; \\ \beta &= (s/a + D)^{\frac{1}{2}}. \end{aligned} \quad (13)$$

Taking into account the “well”, (13) can be written:

$$\begin{aligned} r > r_0 &\rightarrow W_a = \frac{K}{\beta} \cdot \exp(-\beta \cdot r), \\ r &= ((x - X_0)^2 + K_n^2 \cdot (y - Y_0)^2)^{0.5}; \\ r \leq r_0 &\rightarrow W_a = \frac{K}{\beta} \cdot \exp(-\beta \cdot r_0); \\ \beta &= (s/a + D)^{\frac{1}{2}}. \end{aligned} \quad (14)$$

The method for determining the static parameters of the approximating link (12) ($a = j\omega = 0$) takes the following stages.

When equating the static gains of the approximating link (14) to the values of K_1 and K_2 , and assuming that $L_1 - r_0 \approx L_1$ (Figure 2) and $L_2 - r_0 \approx L_2$, we obtain a system of equations:

$$\begin{cases} K_1 = \frac{K}{\beta} \cdot \exp(-\beta \cdot r_0), \\ K_2 = \frac{K}{\beta} \cdot \exp(-\beta \cdot L_1), \\ K_3 = \frac{K}{\beta} \cdot \exp(-\beta \cdot K_n \cdot L_2), \beta = (D)^{\frac{1}{2}}. \end{cases} \quad (15)$$

When substituting in the calculated values $K_1 = 0.0257$, $K_2 = 0.00413$, $r_0 = 0.2$ m, $L_1 = L_x = 50$ m, $L_2 = L_y = 60$ m, and $K_3 = 0.00237$, the following is produced:

$$\begin{cases} 0.0257 = \frac{K}{\beta} \cdot \exp(-\beta \cdot 0.2), \\ 0.00413 = \frac{K}{\beta} \cdot \exp(-\beta \cdot 50), \\ 0.00237 = \frac{K}{\beta} \cdot \exp(-\beta \cdot K_n \cdot 60), \beta = (D)^{\frac{1}{2}}. \end{cases} \quad (16)$$

When solving the resulting system, the following result will be obtained: $D = 0.001449$, $K = 0.0010549$, and $K_n = 1.07647$.

The mathematical model of the object, reflecting the static interaction, is written as follows:

$$\begin{aligned} r > r_0 &\rightarrow W_a = \frac{0.0010549}{\beta} \cdot \exp(-\beta \cdot r), \\ r &= ((x - X_0)^2 + (1.07647)^2 \cdot (y - Y_0)^2)^{0.5}; \\ r \leq r_0 &\rightarrow W_a = \frac{0.0010549}{\beta} \cdot \exp(-\beta \cdot r_0); \\ \beta &= (0.001449)^{\frac{1}{2}}. \end{aligned} \quad (17)$$

During groundwater inflow testing, the dynamic characteristics of the control object were also studied. Since the sensor was located in the well area, the relationship for determining the phase of the approximating link could be written as follows:

$$\Delta\Phi_1 = -\text{Im}(\beta) \cdot r_0 - \arctan(\text{Im}(\beta)/\text{Re}(\beta)), \beta = (j\omega/a + D)^{\frac{1}{2}}. \quad (18)$$

By substituting in the original data $D = 0.001449$, $K = 0.0010549$, $\omega = 0.0005$, and $\Delta\varphi_1 = -1.279$, and solving equation (18) numerically, we obtained $a = 0.00004103$. Since the dynamic characteristics (Figure 3) were measured inside the “well”, the dynamics and statics of the process under consideration can be described by the following equations:

$$\begin{aligned} r > r_0 &\rightarrow W_a = \frac{0.0010549}{\beta} \cdot \exp(-\beta_1 \cdot r), \\ r &= ((x - X_0)^2 + (1.07647)^2 \cdot (y - Y_0)^2)^{0.5}; \\ \beta_1 &= (0.001449)^{\frac{1}{2}}. \\ r \leq r_0 &\rightarrow W_a = \frac{0.0010549}{\beta} \cdot \exp(-\beta \cdot r_0); \\ \beta &= (s/0.00004103 + 0.001449)^{\frac{1}{2}}. \end{aligned} \quad (19)$$

2.4. Determining the Optimal Number of Wells

When developing a deposit, the prospects for the volume of hydromineral raw materials produced are determined. The possibility of constructing a “cluster” of production wells is considered. The area in which to locate the wells in question is determined based on landscape conditions.

A popular region for the extraction of hydromineral raw materials is the Caucasian Mineral Waters region, which has a mountainous landscape. Taking this as the setting, let the production well be positioned in the gorge shown in Figure 5.

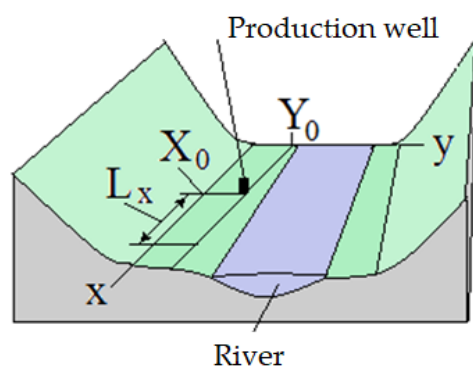


Figure 5. Production well location.

2.5. Statement of the Problem of Optimizing the Number of Wells in a Selected Area

For the deposit under consideration, we sought to determine the number of production wells (n) that provides the maximum income (MI) over ten years of operation. In doing so, we assumed that the extraction of hydromineral raw materials was carried out within 3650 days; $r_{0,i} = 0.2$ m (radii of the “wells”); N —the cost of 1 m^3 of hydromineral raw materials was 500/1,000,000 million RUB; C_p —the cost of developing and maintaining one well for 10 years was 30 million RUB; the subsoil use tax was 7.5%; and P_z —the fixed costs over 10 years amounted to 100 million RUB. At the same time, the value of level decreases (ΔH_i) in the zones where production wells were located was 10 m. The static transmission coefficient was described by relation (17).

The number of parameters that are taken into account when calculating profit can be expanded in accordance with the legislative framework. Increasing the number of parameters will lead to adjustments of the calculating formula.

3. Results

The procedure for solving the problem is divided into the following stages.

The influence of the j -th production well on the decrease in the level in the μ -th production well is described by the following ratio:

$$\Delta H_\mu = \frac{K_\mu \cdot Q_\mu}{\beta} \cdot \exp(-\beta \cdot (r_{0,\mu})) + \sum_{j=1, j \neq \mu}^n \frac{K_j \cdot Q_j}{\beta} \cdot \exp(-\beta \cdot (r_{\mu j})), \beta = (D) \quad (20)$$

where ΔH_μ —level decrease in the considered μ -th well ($\mu = 1, \dots, n$); $r_{0,j}$ —given radius; $r_{\mu,j}$ is the distance from the μ -th to the j -th interacting wells (if the coordinates of producing wells change in x and y , then the calculation of $r_{\mu,j}$ is carried out using (14)); n —number of wells; Q_μ —flow rate of μ -th well; and Q_j —flow rate of the j -th well.

By taking

$$C_\mu = \frac{K_\mu}{\beta} \cdot \exp(-\beta \cdot (r_{0,\mu})), \quad C_{\mu,j} = \frac{K_j}{\beta} \cdot \exp(-\beta \cdot (r_{\mu,j})), \quad \beta = (D)^{\frac{1}{2}}, \quad (21)$$

and transforming it, we obtain a matrix equation for determining the flow rates of production wells.

$$\begin{bmatrix} Q_1 \\ Q_2 \\ \dots \\ Q_n \end{bmatrix} = \begin{bmatrix} C_1, C_{1,2}, \dots, C_{1,n} \\ C_{2,1}, C_2, \dots, C_{2,n} \\ \dots \\ C_{n,1}, C_{n,2}, \dots, C_n \end{bmatrix}^{-1} \cdot \begin{bmatrix} \Delta H_1 \\ \Delta H_2 \\ \dots \\ \Delta H_n \end{bmatrix} \quad (22)$$

The total water flow rate (Q) can be determined from the ratio:

$$Q = \sum_{\mu=1}^n Q_\mu. \quad (23)$$

The profit over ten years of field operation is determined by the following equation:

$$MI = (Q \cdot N - Q \cdot N \cdot 0.075) \cdot 3650 - C_p \cdot n - Pz. \quad (24)$$

A program was developed to calculate the maximum income when the number of production wells was changed. Based on the calculation results, we plotted the graphs shown in Figure 6.

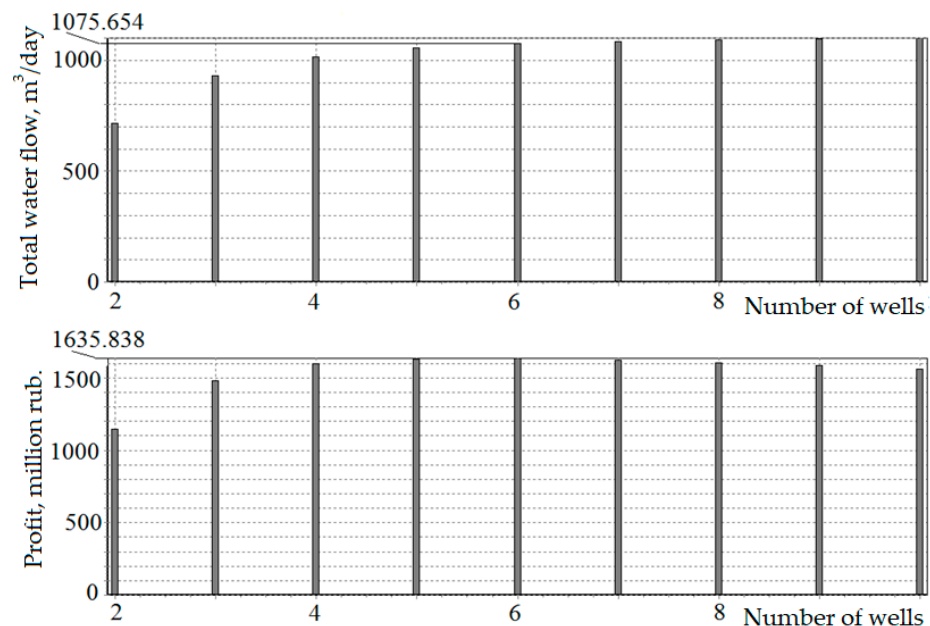


Figure 6. Graphs of changes in water flow rate and profit with an increase in the number of wells.

The optimal number of production wells was found to be six. With six wells, the maximum income over 10 years of field operation will be 1635.838 million RUB and the total daily flow rate will be 1075.654 m³ of mineral water. So, based on the results obtained, there is no need to construct more wells.

4. Discussion

Excessive extraction of water results in a decline in the quality of the extracted mineral water and alters the chemical composition of the product. The expansion of a depression funnel can cause disturbances in the structure of the aquifer system and the disappearance of mineral springs. This scenario is applicable to all water sources situated in the Caucasian Mineral Waters region. Consequently, there is a potential risk of environmental deterioration of mineral water deposits.

Employing mathematical modeling allows for efficient control of the mineral water extraction process. Through the analysis of random factors, it becomes feasible to anticipate a decline in the piezometric level within aquifers. Furthermore, the optimization of the number of production wells can be achieved by developing mathematical models for hydrogeological entities.

Upon commencing production in the field, a limited number of wells is typically installed. With a surge in consumer demand for the supplied products (hydro-mineral raw materials), the number of wells can be increased. However, the construction of wells is very expensive. Furthermore, each well requires constant maintenance, which leads to a significant increase in labor and financial costs. To ensure it is worthwhile, a technique is used to determine the optimal number of wells, which allows the production company to maintain high profitability [1,2,65].

The maximum number of wells (for the aforementioned case) to be outfitted in segment L_x (Figure 6) should be capped at six. As has been shown, with a further increase in the number of wells, the profit does not increase, but even decreases. This suggests that there is no need to build even more wells. This optimal well count ensures not only high profits but also adherence to production standards since it allows for the preservation of mineral water quality and prevents the destruction of aquifers.

Techniques that have been applied to develop specific design solutions have a certain value to science. Furthermore, they can be used in engineering practice for studying hydrolithospheric processes in other fields. For instance, this technique for determining the optimal number of wells can be used not only for models with the skin effect but also for other distributed models [65–68]. Similarly, the outcomes derived from computer modeling methods [29,69,70] can find practical applications in related domains [71,72], particularly in distributed systems with hydrodynamic processes.

5. Conclusions

In this study, we have explored methods for investigating hydrogeological entities and methods for researching distributed control systems using the results of test filtration works. The existing technological process for extracting mineral water from deposits has also been analyzed. Among the main tasks accomplished within the study are the following:

1. Statistical processing of operating mode data for the deposit revealed the stationarity of hydrolithospheric processes, enabling the modeling of stationary random influences. These influences may lead to a decrease in the piezometric level below the permissible norm, disruption of the structure of the hydrogeological entity, and a loss of the mineral water source.
2. Through the results of groundwater inflow testing, the prospects for developing and utilizing new deposits were determined. Additionally, possibilities for upgrading the existing mineral water deposits were identified through the selection of the optimal number of extraction wells.

In the future, we plan to develop a control system using the resulting model. The synthesis of controllers will make it possible to automatic control the process of mineral water extraction in real-time, further increasing environmental safety in the region.

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