

Supplementary Materials for

# Ranking of Empirical Evapotranspiration Models in Different Climate Zones of Pakistan

Mohammed Magdy Hamed<sup>1,2,\*</sup>, Najeebulah Khan<sup>3</sup>, Mohd Khairul Idlan Muhammad<sup>2</sup> and Shamsuddin Shahid<sup>2,\*</sup>

**Table S1.** List of the empirical ET equations and applications.

(a) o	Model (Year)	Equation	Application
(b)	Dalton (1802)	$ET = (0.3648 + 0.07223(u))(e_s - e_a)$	Bormann (2011), Tabari et al (2013), Valipour (2014), Rezaei (2016)
2	Trabert (1896)	$ET = (0.3075)\sqrt{u}(e_s - e_a)$	Bormann (2011), Tabari et al (2013), Valipour (2014), Rezaei (2016)
3	Meyer (1926)	$ET = (0.375 + 0.05026(u))(e_s - e_a)$	Bormann (2011), Tabari et al (2013), Valipour (2014), Rezaei (2016)
4	Rohwer (1931)	$ET = (3.3 + 0.891(u))(e_s - e_a)$	Valipour (2014), Rezaei (2016)
5	Penman (1948)	$ET = (2.625 + 0.000479/u)(e_s - e_a)$	Valipour (2014), Rezaei (2016)
6	Albrecht (1950)	$ET = (0.1005 + 0.297(u))(e_s - e_a)$	Bormann (2011), Tabari et al (2013), Valipour (2014), Rezaei (2016)
7	Makkink (1957)	$ET = 0.61 \left( \frac{\Delta}{\Delta+\gamma} \right) \frac{R_s}{58.5} - 0.12$	Lu et al (2005), Niaghi et al (2013), Rahimikhoob et al (2012)
8	Ivanov (1961)	$ET = 0.00006(25 + T_{mean})^2(100 - RH)$	Valipour (2014), Rezaei (2016)
9	Turc (1961)	$ET = 0.013 \left( \frac{T_{mean}}{T_{mean}+15} \right) (R_s + 50)$	Heydari et al (2013), Lu et al (2005), Xystrakis and Matzarakis (2010)
10	Hamon (1963)	$ET = 0.1651L_d \text{RHOSAT} \times \text{KPEC}$ $\text{RHOSAT} = \frac{216.7 \text{ESAT}}{T_{mean} + 273.3}$ $\text{ESAT} = 6.108 \exp \left( \frac{17.269 \times T_{mean}}{T_{mean} + 273.3} \right)$	Lu et al (2005)
11	Jensen Haise (1963)	$ET = \left( \frac{R_s}{\lambda} \right) (0.025T_{mean} + 0.08)$	Xystrakis and Matzarakis (2010); Heydari et al (2014)
12	Brockamp & Wenner (1963)	$ET = 0.543(u^{0.456})(e_s - e_a)$	Bormann (2011), Tabari et al (2013), Valipour (2014), Rezaei (2016)
13	Papadakis (1966)	$ET = 2.5(e_{ma} - e_a)$	Valipour (2014), Rezaei (2016)
14	WMO (1966)	$ET = (0.1298 + 0.0934(u))(e_s - e_a)$	Bormann (2011), Tabari et al (2013), Valipour (2014), Rezaei (2016)
15	Schendel (1967)	$ET = 16 \left( \frac{T_{mean}}{RH} \right)$	Bormann (2011), Djaman (2015)
16	Mahringer (1970)	$ET = (0.15072)\sqrt{3.6u}(e_s - e_a)$	Bormann (2011), Tabari et al (2013), Valipour (2014), Rezaei (2016)
17	Priestley Taylor (1972)	$ET = \alpha \left( \frac{\Delta}{\Delta+\gamma} \right) \frac{R_n}{\lambda}$	Bormann (2011), Niaghi et al (2013), Rahimikhoob et al (2012)
18	McGuinness & Bordne (1972)	$ET = (0.0082T_{mean} - 0.19) \frac{R_s}{1500} (2.54)$	Tabari et al (2013), Heydari et al (2014)
19	Szasz (1973)	$ET = 0.00536(T_{mean} + 21)^2 \left( 1 - \frac{RH}{100} \right)^{2/3} f(u)$ $f(u) = (0.0519u) + 0.905$	Rácz et al (2013)
20	Caprio (1974)	$ET = \left( \frac{6.1}{10^6} \right) R_s (1.8T_{mean} + 1.0)$	Oudin et al (2005)

21	Blaney Criddle (1977)	$ET = k \times p(0.46T_{mean} + 8.13)$	Niaghi et al (2013, Heydari et al (2014), Josilva et al (2016), Poyen et al (2018))
22	Linacre (1977)	$ET = \frac{700(T_{mean} \pm 0.006z)}{100 - L} + 15(T_{mean} - T_d)$	Josilva (2016)
23	Kharuffa (1985)	$ET = 0.34pT_{mean}^{1.30}$	Heydari et al (2014)
24	Hargreaves Sa- mani (1985)	$ET = \left(0.0023 \frac{R_a}{2.45}\right) TD^{0.5} (T_{mean} + 17.8)$	Heydari et al (2013), Oudin et al (2005), Rahimikhoob et al (2012), Xu and Singh (2002), Xystrakis and Matzarakis (2010), Josilva et al (2016)
25	Ritchie (1990)	$ET = \alpha_1 (3.87 \times 10^{-3}) (R_s (0.6T_{max} + 0.4T_{min} + 29))$	Tabari et al (2013)
26	Abtew (1996)	$ET = 0.53 \left(\frac{R_s}{\lambda}\right)$	Abtew, 1996
27	FAO56 PM (1998a)	$ET = \frac{0.408(R_n - G) + \gamma \frac{900}{T_{mean} + 273} u(e_s - e_a)}{\Delta + \gamma(1 + 0.34u)}$	Heydari et al (2013), Niaghi et al (2013), Rahimikhoob et al (2012), Tabari et al (2013), Valipour (2014), Rezaei (2016)
28	Irmak-Rs (2003)	$ET = -0.611 + 0.149R_s + 0.079T_{mean}$	Heydari et al (2013), Niaghi et al (2013), Tabari et al (2013), Heydari et al (2014)
29	Irmak-Rn (2003)	$ET = 0.489 + 0.289R_n + 0.023T_{mean}$	Heydari et al (2013), Niaghi et al (2013), Tabari et al (2013), Heydari et al (2014)
30	Trajkovic (2007)	$ET = (0.0023R_a)TD^{0.424}(T_{mean} + 17.8)$	Djaman (2015)
31	Ravazzani (2012)	$ET = (0.817 + 0.00022z)(0.0023R_a)(TD^{0.5})(T_{mean} + 17.8)$	Djaman (2015)

ET is the potential evapotranspiration ( $\text{mm day}^{-1}$ ).  $ET$  is in  $\text{mm day}^{-1}$  in all equations except Ritchie and McGuinness & Bordne models, ET is in  $\text{cm day}^{-1}$ .  $R_n$  is the net radiation ( $\text{MJ m}^{-2} \text{day}^{-1}$ ).  $G$  is the soil heat flux ( $\text{MJ m}^{-2} \text{day}^{-1}$ ).  $R_a$  is the extraterrestrial radiation ( $\text{MJ m}^{-2} \text{day}^{-1}$ ).  $\gamma$  is the psychrometric constant ( $\text{kPa}^{\circ}\text{C}$ ).  $e_s$  is the saturation vapour pressure ( $\text{hPa}$ ).  $e_a$  is the actual vapour pressure ( $\text{hPa}$ ).  $e_s$  and  $e_a$  are in  $\text{hPa}$  in all equations except Papadakis, Rohwer, Penman and FAO56 Penman-Monteith models,  $e_s$  and  $e_a$  are in  $\text{kPa}$ .  $\Delta$  is the slope of the saturation vapour pressure–temperature curve ( $\text{kPa}/^{\circ}\text{C}$ ).  $\lambda$  is latent heat of evaporation ( $\text{MJ/kg}$ ).  $T_{mean}$  is the average daily air temperature ( $^{\circ}\text{C}$ ).  $T_{mean}$  is in  $^{\circ}\text{C}$  in all equations except McGuinness & Bordne model,  $T_{mean}$  is in degree Fahrenheit.  $U$  is the mean daily wind speed at 2 m ( $\text{m s}^{-1}$ ).  $f(u)$  is function of wind speed.  $Z$  is the elevation (m).  $L$  is local latitude (degrees).  $T_d$  is dew point temperature ( $^{\circ}\text{C}$ ).  $T_{min}$  is the minimum air temperature ( $^{\circ}\text{C}$ ).  $T_{max}$  is the maximum air temperature ( $^{\circ}\text{C}$ ). TD is maximum and minimum temperature difference ( $^{\circ}\text{C}$ ). RH is the average relative humidity (%).  $R_s$  is the solar radiation.  $R_s$  is in  $\text{MJ m}^{-2} \text{day}^{-1}$  in all equations except Turc, Makkink, Ritchie and McGuinness & Bordne models,  $R_s$  is in  $\text{Cal/m}^2 \text{day}$  and Caprio model,  $R_s$  is in  $\text{kJ/m}^2 \text{day}$ .  $e_{ma}$  is the saturation vapour pressure at the monthly mean daily maximum temperature (kPa).  $k$  is monthly consumptive use coefficient which depends on location, season and vegetation type. The average value of  $k$  is 0.8 which also requires local calibration (Poyen et al, 2018).  $p$  is the mean annual percentage of daytime hours for different latitudes that can be obtained from Doorenbos and Pruitt (1977).  $p$  is expressed as constant (0.274) in Josilva et al (2016).  $L_d$  is daytime length in multiples of 12 h. RHOSAT is saturated vapor density ( $\text{g m}^{-3}$ ). ESAT is saturated vapor pressure (mbar). KPEC is calibration coefficient (1.2).  $\alpha$  is constant (1.26).  $\alpha_1$  is constant (1.1).