

Article

Distribution and Characteristics of Nanotubular Halloysites in the Thach Khoan Area, Phu Tho, Vietnam

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Received: 23 May 2018; Accepted: 25 June 2018; Published: 8 July 2018



Abstract: Two types of halloysite collected from the upper (UPS) and lower (LOS) zones of a weathered pegmatite profile in the Thach Khoan area, Phu Tho were defined by X-ray diffraction (XRD), scanning electron microscopy (SEM), transmission electron microscopy (TEM), Fourier transform infrared spectroscopy (FT-IR), thermal analysis (TG and DTG), and N₂ adsorption-desorption isotherms. XRD analysis showed that halloysite and kaolinite coexist in samples of size fractions <2 µm. Semi-quantitative analysis by XRD after formamide (FA) treatment indicated that the halloysite contents are approximately 81% and 93% in UPS and LOS samples, respectively. The results of SEM and TEM analyses showed that while short halloysite type is mainly distributed in the upper zone, long halloysite type occurs primarily in the lower zone of the weathered pegmatite profile. The length of short halloysite ranging from 250 to 750 nm is most popular, accounting for 47.2% of halloysites in the UPS sample. Meanwhile, long halloysites which have the length of 750–1250 nm are dominant in the LOS sample with 69.9%. In addition, short halloysites with outer diameter of >100 nm constitute 79.1% of halloysites in the UPS sample while long halloysites with outer diameter of 50–100 nm make up 74.2% of halloysites in LOS sample. Specific surface areas are 15.7434 and 22.0211 m²/g and average pore sizes are 18.9837 and 17.0281 nm for the UPS and LOS samples, respectively. The analysis implies that although forming under same natural geographical and climatic conditions, halloysites at different depths in the weathered pegmatite profile may have different morphological and other properties.

Keywords: halloysite; kaolinite; nanotube; pegmatite; Phu Tho

1. Introduction

Halloysite was originally depicted as a 1:1 layered aluminosilicate mineral of the kaolin group by Berthier [1]. The chemical composition structure of halloysite is similar to that of minerals of the kaolin group (kaolinite, nacrite and dickite minerals) but the unit layers of halloysite are isolated by a monolayer

of water molecules [2–4]. Halloysite appears mainly in two different polymorphs: a chemical formula $\text{Al}_2\text{Si}_2\text{O}_5(\text{OH})_4 \cdot 2\text{H}_2\text{O}$ when fully hydrated and $\text{Al}_2\text{Si}_2\text{O}_5(\text{OH})_4$ when dehydrated [5].

Halloysite can be found in a variety of particle morphologies, such as short-tubular, large-tubular and spheroidal, and platy shapes [6]. However, nanotubular morphology is the most common shape of halloysite. The tubular shape can be considered as rolled kaolin sheets with an inner diameter of 1/30 nm, an outer diameter of 30/50 nm and a length of 100–2000 nm [7,8]. The interior surface of halloysite is composed of siloxane (Si–O–Si) groups, while the external is a gibbsite-like array of aluminol (Al–OH) groups [9,10].

Halloysite deposits have been discovered and exploited in different countries such as New Zealand, United States, Australia, China, Brazil, and Turkey [11]. This mineral can be formed both by weathering of igneous rocks and their hydrothermal alteration [12–19]. For instance, the Matauri Bay (New Zealand) halloysite deposit was formed by hydrothermal alteration at low temperature of rhyolite and dacite volcanic rocks [20]. The large mass of halloysite at the Dragon Mine (UT, USA) was formed by irregular replacement of Early Paleozoic dolomite rock in contact with hydrothermal fluids channeled along the Dragon Fissure Zone [21]. Halloysite at TePuke is a weathering product of volcanic rocks of rhyolite and andesite in the Bay of Plenty, New Zealand [22]. The above literature and others have shown that halloysites from different areas also have different morphological and physicochemical properties [7,23–25].

In the recent years, due to its superior properties such as tubular structure, non-toxicity, large surface area, high mechanical strength, lower cost compared to nanotubular carbon, halloysite has attracted considerable attention of scientists and many new possibilities of application [8,26–34]. However, in many cases, differences in morphology, size, as well as other properties of halloysites may have certain impacts on their applicability in practice. For instance, Makaremi et al. [35] used two different types of halloysite nanotubes to improve the properties of apple pectin bionanocomposites as potential films for food packaging applications. Results indicated that the short halloysites with 50–3000 nm length and 50–200 nm outer diameter had better ability for the encapsulation of salicylic acid into their lumen, while the long halloysites with 200–30,000 nm length and 40–55 nm outer diameter made the encapsulation process more difficult. Zheng and Ni [36] prepared an efficient flame-resistant composite using the pentaerythritol-loaded halloysites for the UV-curable epoxy resin. In this study, halloysites have length 300–1000 nm, outer diameter 50–70 nm and BET surface area $36.40 \text{ m}^2/\text{g}$. The obtained composite showed a low moisture absorption and a good stability of the mechanical properties. Pasbakhsh et al. [37] have studied the properties of some halloysites in the world, and have given orientations for their applications. For example, the long-tubular halloysite with 200–5000 nm length and 40–55 nm outer diameter are very suitable for use both as an additive and a carrier. The halloysite tubes showing a wide variation in size may be well suited as microfiber filler. Thus, it can be seen that studying the properties of halloysites from different deposits or even in a deposit is necessary before using them for different applications.

This study aims to study the distribution and characteristics of two types of halloysite nanotubes from a weathered pegmatite profile in the Thach Khoan area, Phu Tho Province. Different characteristics of these halloysites were determined using X-ray diffraction (XRD), scanning electron microscopy–energy dispersive X-ray spectroscopy (SEM-EDS), transmission electron microscopy (TEM), Fourier transform infrared spectroscopy (FT-IR), thermal analysis (TG and DTG), and N_2 adsorption-desorption isotherms. The results showed that halloysites from different depths of the weathered pegmatite in the study area have different morphological properties. This information is useful for the understanding of distribution and characteristics of halloysites in the deposit and helping for exploitation and use of these nanotubular minerals effectively.

2. Materials and Methods

2.1. General Geological Setting of Study Area

The study area has many pegmatite bodies with different sizes related to the Late Paleozoic Tan Phuon granite Complex [38]. The surrounding rocks of pegmatite bodies are the metamorphic Thach Khoan formation of Proterozoic age (Figure 1). The composition of this formation consists mainly of mica quartz schist, mica schist, staurolite-bearing quartz, disten, sillimanite, and garnet.

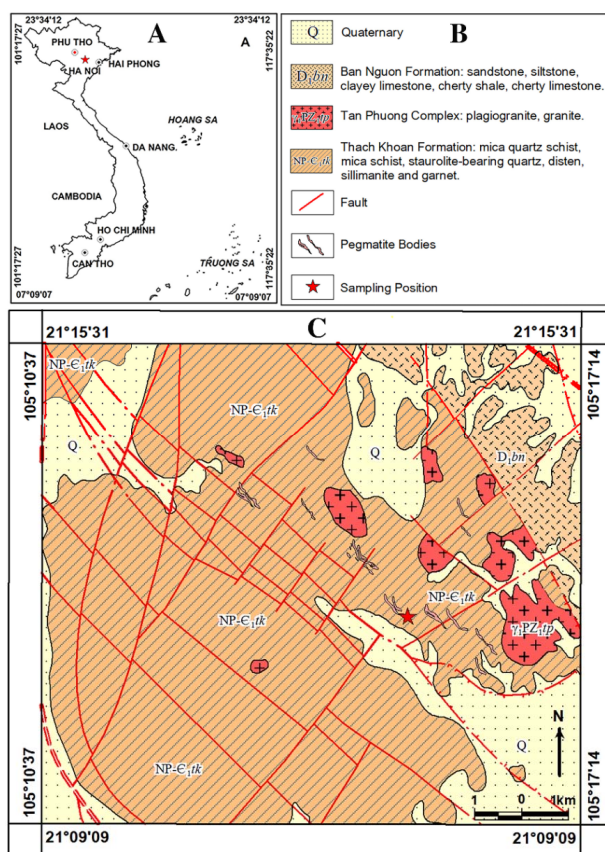


Figure 1. Vietnam map and the site of the study area (A); geological legend (B) and geological map of the study area (C).

The pegmatite bodies have the strike of 60° N–80° W, dipping to the southwest with a slope of 50°–80°. They vary from several hundred up to thousands of meters in length and from tens to hundreds of meters wide. All pegmatite bodies have a similar weathering profile with an upper brown yellow zone (15–20 m), a middle pink zone (5–10 m) and a lower white, light orange zone (5–15 m).

2.2. Samples

A typical outcrop about 40 m high that has the GPS position of 21°11'31" N and 105°15'07" E, was prepared for sampling. For comparison purposes, two samples were collected separately. The first sample, called UPS sample, was taken in upper zone, and the second one, LOS sample, was from the lower zone of the weathered pegmatite profile. The samples were taken from the top down, perpendicular to the weathering layers. Separated samples were mixed homogeneously before using for further steps.

The bulk samples were first dissolved in deionized water by repeated ultrasonic vibration. A portion of the <2 µm clay sample fraction was obtained using the decantation method. The clay fractions were then freeze dried and examined by different analyses.

2.3. Characterization

X-ray diffraction (XRD) patterns of the samples were collected by using a D8-Advance Bruker diffraction (Bruker Corporation, Billerica, MA, USA) with radiation of $\text{CuK}\alpha$ ($\lambda = 1.5406 \text{ nm}$) generated at 40 kV and 40 mA. The data were archived in the Bragg angle (2θ) range of 3° – 70° with scanning speed of 2° min^{-1} . Minerals were defined by using the software of Evaluation 10.0 with database (PDF-2 2004) provided by the International Centre for Diffraction Data. Formamide (FA) treatment was used to estimate the content of halloysite and kaolinite in the samples [22].

The Fourier transform infrared (FT-IR) spectra for each sample were achieved in transmission mode on pellets containing a pressed mixture of approximately 1.0 mg of the sample and 100 mg of KBr. The IR spectra were recorded in the range from 4000 to 400 cm^{-1} with a resolution of 2 cm^{-1} (Shimadzu IR Prestige-21 spectro-meter instrument, Kyoto, Japan).

Scanning electron microscope (SEM) coupled with energy dispersive X-ray spectroscopy (EDS) (Quanta 450, FEI Company, Hillsboro, OR, USA) were initially used to analyze the morphology of minerals and elements present in the samples. Transmission electron microscopy (TEM) images were obtained by a JEM 1010 operated at an accelerating voltage of 200 kV. The samples were suspended by using a drop-wise of ethanol and evaporated on 200 mesh copper grids covered with amorphous Formvar carbon.

Thermogravimetric analyses (TG) were carried out on a SETARAM Instrument (Caluire-et-Cuire, France). Approximately 2–3 mg of the samples were heated from 50 to 1050°C in a platinum crucible with a heating rate of $10^\circ\text{C min}^{-1}$, under an atmosphere of high purity N_2 .

The specific surface area of the samples was measured from N_2 gas adsorption at 77 K by using a TriStar 3000 (Micromeritics Corp., Norcross, GA, USA). Surface areas were calculated from the linear part of the (Brunauer-Emmett-Teller) BET plot. The N_2 isotherms and the Barret-Joyner-Halenda (BJH) method were used to calculate pore size distributions of halloysites.

3. Results and Discussion

3.1. XRD Analysis

XRD patterns of the UPS and LOS samples with size fractions $<2 \mu\text{m}$ in natural condition are presented in Figure 2. The results indicated that minerals of kaolin group coexisted in the samples. The basal reflections of 10 \AA halloysite were recorded at peaks of 10.0° , 4.44° , 3.36° and 2.56° ($\text{Al}_2\text{Si}_2\text{O}_5(\text{OH})_4 \cdot 2\text{H}_2\text{O}$, hexagonal structure, PDF No. 00-29-1489). The peaks at 7.38° and 3.60° refer to kaolinite ($\text{Al}_2\text{Si}_2\text{O}_5(\text{OH})_4$, with a triclinic structure, PDF No. 01-089-6538).

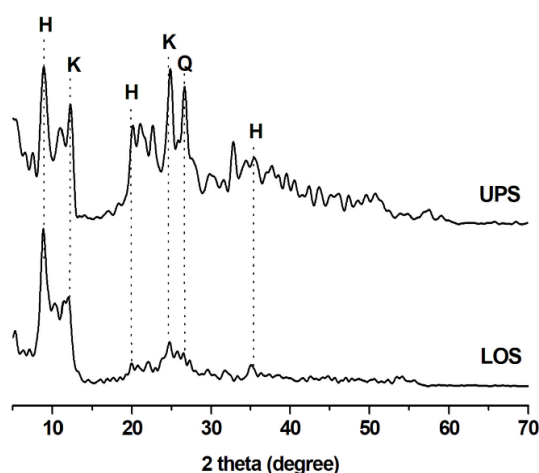


Figure 2. XRD results of the UPS and LOS samples. H—halloysite; K—kaolinite; Q—quartz.

For the measurement of the content of halloysite and kaolinite, the formamide (FA) treatment was applied to the samples following Churchman et al. [22]. The percentage of halloysite a in a sample was defined by the equation:

$$a = \frac{I_{10}}{(I_7 + I_{10})} \cdot 100\%, \quad (1)$$

where I_7 and I_{10} denote the height of the peaks near 7 and 10 Å of XRD patterns, respectively. Figure 3 shows the XRD results of UPS and LOS samples with size fractions <2 µm after formamide treatment. It can be seen that the intensities of the 10 Å peak were significantly higher than those of the 7 Å peak in both the UPS and LOS samples. The estimated percentages of halloysite in the samples from Figure 3 were approximately 81% and 93% for UPS and LOS samples, respectively.

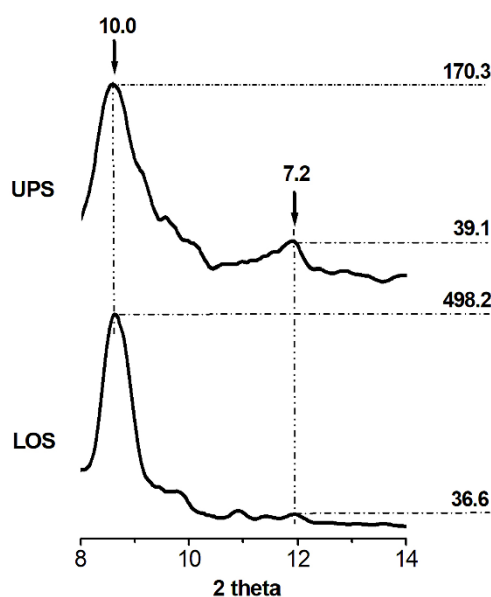


Figure 3. XRD results of the UPS and LOS samples after formamide treatment.

One of the models concerning the transformation from halloysite to kaolinite during weathering of crystalline rocks was based on the dissolution and recrystallization mechanism [39,40]. According to this model, the thermodynamics of halloysite is less stable than that of kaolinite under the prevailing weathering conditions. Thus, halloysite is formed early in the weathering profile, then, dissolved and kaolinite is eventually crystallized under suitable weathering conditions such as time, the activity of water table in each weathering zones. In this study, that the percentage content of halloysites in the lower zone (93%) is higher than that in the upper zone (81%) of the weathered pegmatite profile is consistent with the previously reported results of Inoue et al. [41].

3.2. FT-IR Analysis

Figure 4 shows the FT-IR spectra of the two samples (UPS and LOS). It can be seen that the IR spectra of the samples were quite similar and all present the existence of kaolin minerals [42]. The absorptions bands at 3696 and 3620 cm^{−1} in the FTIR spectra are assigned to the stretching vibration due to the inner-surface of O–H groups. The absorption at 1640 cm^{−1} is assigned to the interlayer water [43]. The intensity of this absorption band increases as the interlayer water content increases. It may come from the significant content of halloysite in the LOS sample.

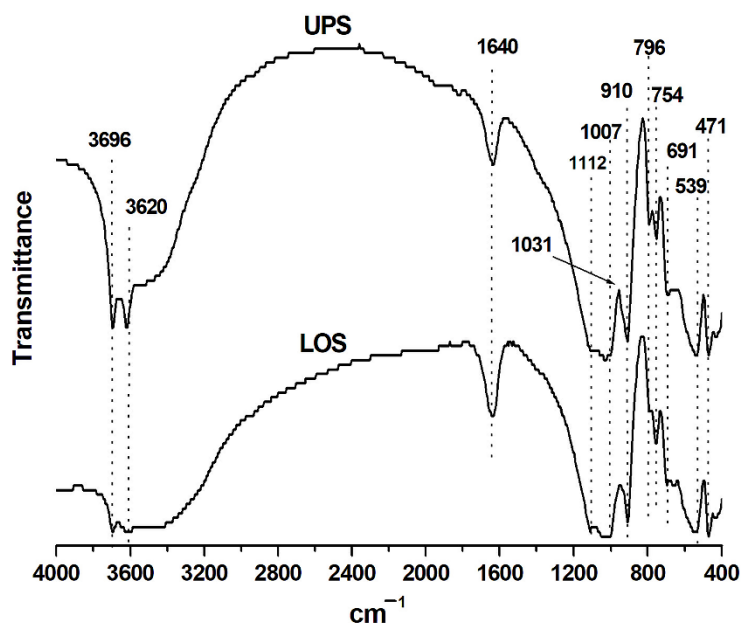


Figure 4. FT-IR graphs of the UPS and LOS samples.

The Si–O stretching region comprises three absorption bands at 1112, 1031, and 1007 cm^{-1} . The 1031 cm^{-1} peak of LOS disappears. The reason for the disappearance is due to a decrease of kaolinite content and increase of halloysite content in those samples. The band at 910 cm^{-1} is assigned to the bending vibration of Al–OH. The bands observed at 796, 754 cm^{-1} of all samples are assigned to the stretching mode of Al–O–OH. The low stretching band Si–O defined at 691 cm^{-1} confirms the existence of halloysite in the sample LOS. The bands at 539 and 471 cm^{-1} are due to the vibration of Al–O–Si and Si–O–Si, respectively [44].

3.3. Electron Microscopy Analysis

The scanning electron micrograph (SEM) images and EDS data of the samples with the size fraction $<2 \mu\text{m}$ are shown in Figure 5. The rod-shaped minerals were interwoven and overlapped each other as matrices. From these images, it can be seen that there may be two types of halloysite available in the samples: short halloysites in the upper zone (UPS sample) and long halloysites in the lower zone (LOS sample) of the weathered pegmatite profile in the study area. EDS spectra shows the main elements of Al, Si, and O, which are relative to halloysite chemical formula $(\text{Al}_2\text{Si}_2\text{O}_5(\text{OH})_4 \cdot 2\text{H}_2\text{O})$. The transmission electron micrographs in Figure 6 also display tubular morphology of these minerals clearly. Under the same magnification, the lengths of halloysite in the UPS sample are generally shorter than the lengths of halloysite in the LOS sample (Figure 6A,B). Closed view of these minerals are presented in Figure 6A₁,B₁. The distributions of the lengths and outer diameters of these halloysites using TEM images are presented in Figures 7 and 8, respectively. Results showed that short halloysites in UPS sample are distributed mainly in the length range from 250 to 750 nm, accounting for 47.2% of halloysites in the sample. Meanwhile, long halloysites are dominant in the LOS sample with 69.9% of a length range from 750 to 1250 nm (Figure 7). In addition, short halloysites with an outer diameter of $>100 \text{ nm}$ constitute 79.1% of halloysites in the UPS sample, while long halloysites with an outer diameter of 50–100 nm make up 74.2% of halloysites in the LOS sample (Figure 8). This difference in size of halloysites between weathering zones may be due to the structure of early formed halloysites in the upper zone partially replaced by new small kaolinites crystals [41].

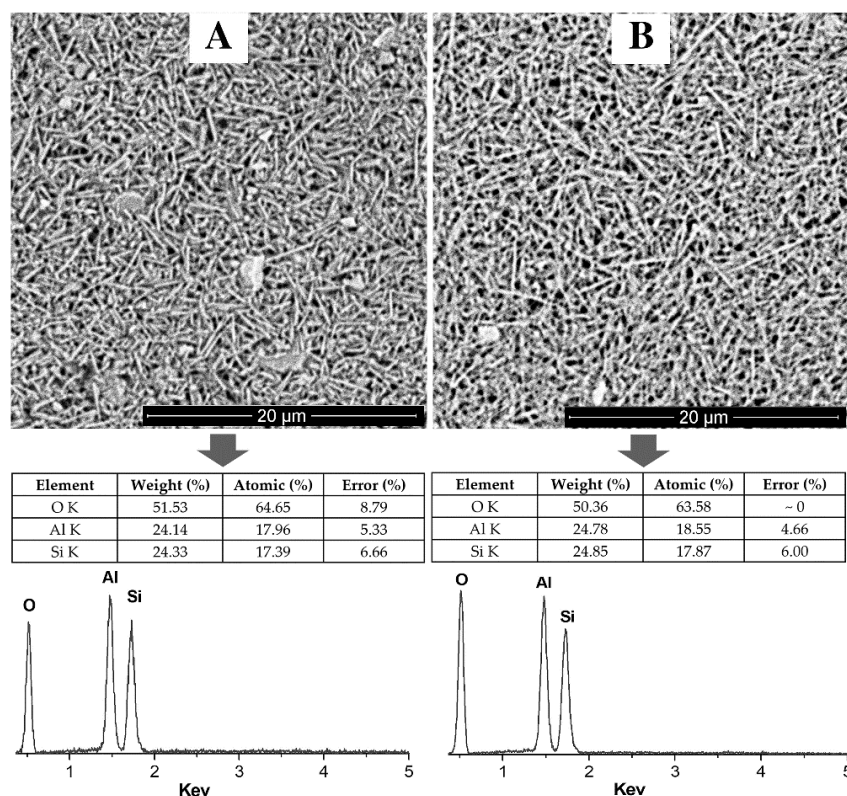


Figure 5. SEM-EDS images of the UPS and LOS samples with the size fraction <2 µm. LOS sample (A) and UPS sample (B). O—oxygen; Al—aluminum; Si—silicon.

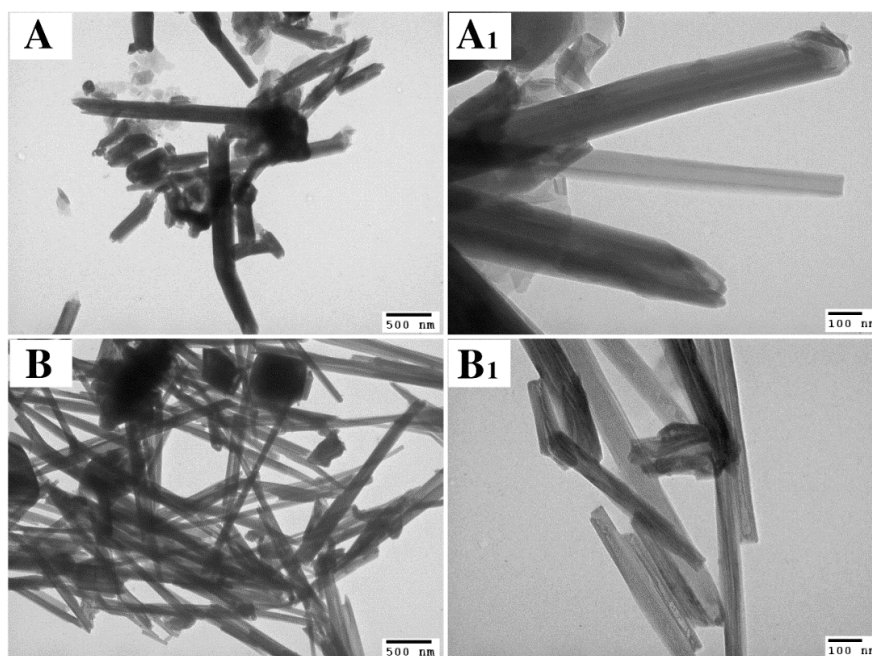


Figure 6. TEM images of the UPS (A,A₁) and LOS (B,B₁) sample. Scale bar in (A,B) images represents the length of 500 nm and in (A₁,B₁) images is 100 nm.

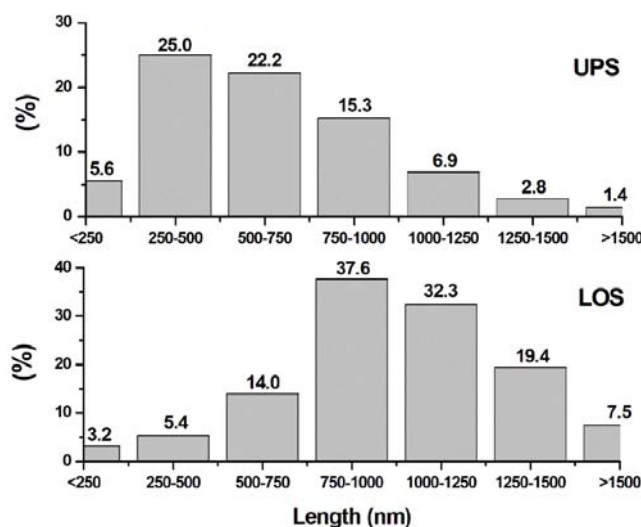


Figure 7. The distributions of the lengths of halloysite from TEM for the UPS and LOS samples. Unit is nanometer (nm).

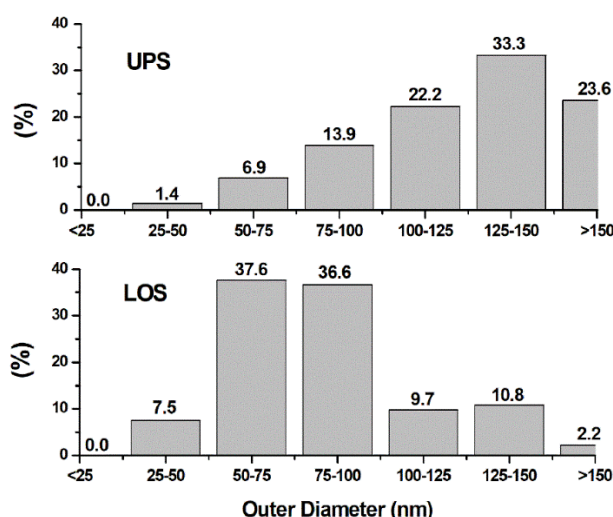


Figure 8. The distributions of the outer diameters of halloysite from TEM for the UPS and LOS samples. Unit is nanometer (nm).

3.4. Thermal Analysis

Figure 9 presents the weight loss traces, thermogravimetry (TG) and derivative thermogravimetry (DTG) curves of the UPS and LOS samples. As can be seen, two main mass loss steps were determined in the TG curves. At first, the endothermic peaks at 91.4 °C with a mass loss of 0.6% and at 87.5 °C with a mass loss of 9.5% are ascribed to the removal of physisorbed water in the UPS and LOS samples, respectively. In the second endothermic peak at 510.4 °C (mass loss of 12.2%) for UPS and at 516.0 °C (mass loss of 13.0%) for LOS are due to the dehydroxylation of the structural aluminol groups in halloysite. The TG (DTG) curves of these thermal analysis were in agreement with previous literature [45–47]. The difference of halloysite percentages in samples (81% and 93% halloysites for UPS and LOS samples, respectively) may be one of the reasons for differences in temperature of the endothermic peaks and their lost weights [19].

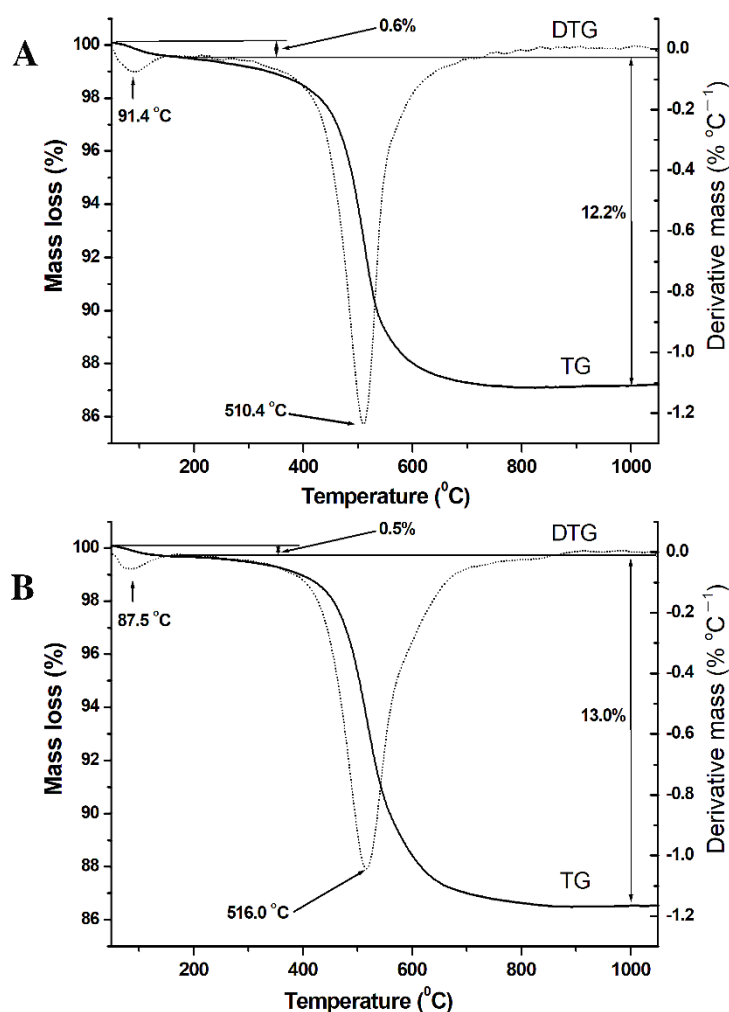


Figure 9. Thermogravimetry (TG) and derivative thermogravimetry (DTG) graphs of the UPS (A) and LOS (B) samples.

3.5. Surface Area and Pore Size

The nitrogen adsorption-desorption isotherms of the UPS and LOS samples are displayed in Figure 10. These isotherms exhibit type II with an H3 hysteresis loop in the relative pressure (P/P_0) and this type of isotherm is a typical characteristic of mesoporous structures [48]. The values for specific surface areas (S_{BET}), the cumulative specific surfaces obtained from isotherms for both adsorption (S_{ads}) and desorption (S_{des}) of nitrogen on each sample are summarized in Table 1. The surface areas were determined for the UPS and LOS samples with S_{BET} of 15.7434 and 22.0211 m^2/g , respectively. Halloysites in the LOS sample have higher value of S_{BET} than that of the halloysites in the UPS sample because halloysites in the LOS sample have longer and thinner cylindrical structure [37,49]. Average pore sizes of for the UPS and LOS samples are 18.9837 and 17.0281 nm, respectively.

Table 1. Surface area and pore size data of the UPS and LOS samples.

Sample	S_{BET} (m^2/g)	S_{ads} of Pores (1.70–300.0 nm Diameter) (m^2/g)	S_{des} of Pores (1.70–300.0 nm Diameter) (m^2/g)	V_{ads} of Pores (1.70–300.0 nm Diameter) (cm^3/g)	Average Pore Size (nm)	$S_{\text{BET}} - S_{\text{ads}}$ (m^2/g)	$S_{\text{des}} - S_{\text{BET}}$ (m^2/g)
UPS	15.7434	15.361	19.2166	0.072904	18.9837	0.3824	3.4732
LOS	22.0211	22.350	27.9054	0.095144	17.0281	−0.3289	5.8843

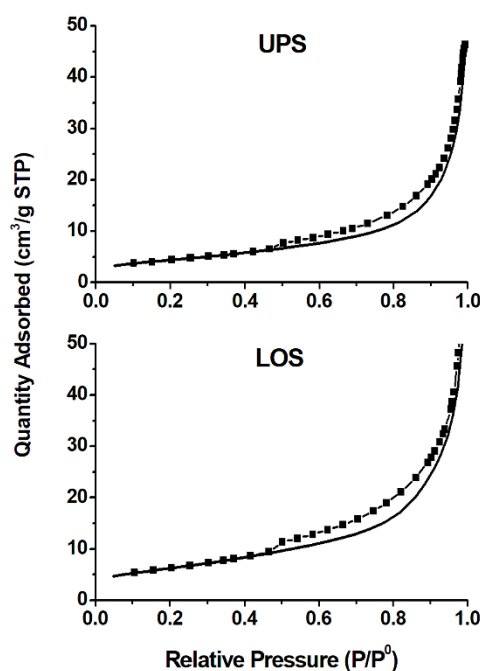


Figure 10. Nitrogen gas adsorption-desorption isotherms of the UPS and LOS samples.

The low discrepancy between S_{ads} and S_{BET} indicates that these samples likely contain mainly cylindrical pores of varying radius and slit-shaped pores are the dominant shape in both samples [39,48,49].

The pore size distributions using the Barrett-Joyner-Halenda (BJH) theory for the UPS and LOS samples are presented in Figure 11. From Figure 11, the narrow peaks centered at 4.3 nm are signed to internal/surface pores, including spaces between the overlaps of folded halloysite sheets in the samples [39]. The stronger intensity of this peak in the LOS sample indicates that halloysite formed a more concentrated and uniform pore size distribution. The peaks at 9.2, 10.7, and 13.4 nm are identified as the lumens of halloysites in the samples and are agreeable with measurements on TEM images.

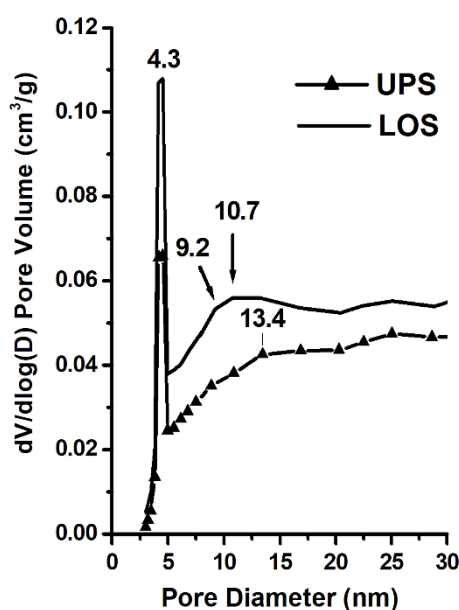


Figure 11. BJH pore size distribution of the halloysite in the UPS and LOS samples.

4. Conclusions

In conclusion, two main types of halloysite were formed in the weathered pegmatite profile in the Thach Khoan area, Phu Tho, northern part of Vietnam. Analysis methods of XRD, SEM-EDS, TEM, FT-IR, TG and N₂ adsorption-desorption isotherms were used to characterize these halloysites. The results showed that the short halloysite type is mainly distributed in the upper zone and long halloysite type can be found in the lower zone of the weathered pegmatite profile. The short halloysites have the length ranging mainly from 250 to 750 nm, the outer diameter of >100 nm (79.1%), the specific surface areas of 15.7434 m²/g and the average pore sizes of 18.9837 nm. Meanwhile, the length ranging mainly from 750 to 1250 nm (69.9%), the outer diameter of 50–100 nm (74.2%), the specific surface areas of 22.0211 m²/g, and the average pore sizes of 17.0281 nm are properties of the long halloysites. XRD after formamide (FA) treatment indicated that the halloysite contents are approximately 81% and 93% for the upper zone and the lower zone of the weathered pegmatite profile, respectively. The results provided useful information for the understanding of distribution and characteristics of different halloysites in the deposit and for exploiting and using these nanotubular minerals effectively.

Author Contributions: Conceptualization, B.H.B.; Formal analysis, K.T.H., D.M.A., T.T.V.A. and D.V.C.; Funding acquisition, B.H.B.; Investigation, K.T.H., T.T.V.A., D.V.C. and B.T.T.; Methodology, N.T.D., L.Q.K. and N.V.L.; Project administration, B.H.B.; Resources, N.T.D.; Validation, L.Q.K.; Visualization, D.M.A. and P.V.S.; Writing—review & editing, B.H.B.

Funding: This research was supported by the Ministry of Education and Training of Vietnam (MOET) under grant number B2017-MDA-17TD and by the Vietnam National Foundation for Science and Technology Development (NAFOSTED) under grant No. 105.99-2017.317.

Conflicts of Interest: The authors declare no conflicts of interest.

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