

Article



Microwave-Assisted Acid Activation of Clays Composed of 2:1 Clay Minerals: A Comparative Study

Juan Antonio Cecilia ¹^(D), Laura Pardo ¹, Manuel Pozo ²^(D), Eva Bellido ³^(D) and Francisco Franco ^{1,*}

- ¹ Department of Inorganic Chemistry, Crystallography and Mineralogy, Facultad de Ciencias, Campus de Teatinos s/n, Universidad de Málaga, 29071 Málaga, Spain; jacecilia@uma.es (J.A.C.); lapardo@uma.es (L.P.)
- ² Department of Geology and Geochemistry, Facultad de Ciencias, Universidad Autónoma de Madrid, Campus de Cantoblanco, 28049 Madrid, Spain; manuel.pozo@uam.es
- ³ Instituto Geológico y Minero de España (IGME), 28760 Madrid, Spain; e.bellido@igme.es
- * Correspondence: ffranco@uma.es; Tel.: +34-952-136-644

Received: 7 July 2018; Accepted: 27 August 2018; Published: 30 August 2018



Abstract: The effect of the microwave-assisted acid treatment (MAT) on the structure and texture of dioctahedral (montmorillonite) and trioctahedral (saponite) smectites, kerolitic clays and sepiolites were studied by scanning electron microscopy, nitrogen adsorption, X-ray fluorescence, X-ray diffraction and Fourier transform infrared spectroscopy. This study shows that in the smectite group of minerals the effectiveness of MAT is notably influenced by the chemical composition of the octahedral sheet, whereas in kerolitic clays it is influenced by the proportion of the expandable phase in the mixed layer sequence. On the other hand, the reactivity of sepiolites depends on the width of the natural fibres. With these treatments, extremely high specific surface area (SSA) increments are achieved in just 16 min. The SSA increments reach values up to $231 \text{ m}^2/\text{g}$ (% $\Delta S_{BET} = 80$) in sepiolites, $198 \text{ m}^2/\text{g}$ (% $\Delta S_{BET} = 155$) in Mg-smectites, and $161 \text{ m}^2/\text{g}$ (% $\Delta S_{BET} = 61$) in kerolitic clays.

Keywords: montmorillonite; saponite; kerolite/Mg-smectite mixed layers; sepiolite; microwave-assisted acid treatment

1. Introduction

Some clay materials named after special clays are composed of mostly one clay mineral that has important physical and chemical properties for the industrial and technological applications. Composed of 2:1 clay minerals, the main special clays include bentonite, sepiolite and kerolitic clay.

Sepiolite $(Mg_8Si_{12}O_{30}(OH)_4(OH_2)_4 \cdot nH_2O)$ is a clay mineral, the structure of which is similar to other 2:1 trioctahedral phyllosilicates (e.g., talc). Sepiolite has a continuous two-dimensional tetrahedral sheet, but it lacks continuous octahedral sheets. Indeed, discontinuities in the octahedral sheets and inversions of the silica sheets give rise to structural tunnels and ribbons expanded along the *a*-axis. Sepiolite is included in the class of modulated non-planar hydrous phyllosilicates [1,2].

The structure of the smectite group of clay minerals is based on the stacking of negatively charged 2:1 layers. For this reason, the presence of hydrated exchangeable cations, which are located in interlayer positions (mostly Ca^{2+} , Mg^{2+} , Na^+) [3] is necessary. Moreover, depending on whether the octahedral sheet is filled with trivalent or divalent cations, smectite minerals can be either dioctahedral (e.g., montmorillonite) or trioctahedral (e.g., saponite, stevensite).

Kerolitic clays are composed of kerolite/Mg–smectite mixed layers. Random interstratifications (mixed layers) of kerolite and Mg-smectite (stevensite, saponite) layers can occur [4–8]. The structure of kerolite is believed to be similar to talc (and thus, a variety of talc), but with the presence of H_2O ,

either on the edges associated with broken bonds or in the interlayer [9]. Brindley et al. reported the composition as $Mg_3Si_4O_{10}(OH)_2 \cdot nH_2O$ where n = 0.8-1.2 [10].

The physical and chemical properties of these special clays, and therefore their industrial applications [11], are influenced by several features such as: high cation exchange capacity, high surface area, laminar to fibrous shape, small particle size, and layer charge. The activation is a common process to enhance the properties of these clay materials.

Various surface modification methods (activation), such as heat and acid treatment, have been studied by many researchers [12–31] in order to improve the adsorption capacity of sepiolites. Acid treatments in particular, have classically been used to increase the surface area and to obtain solids with high porosity and a high number of acidic centres from this clay mineral [17]. With these acid treatments, variable amounts of structural Mg^{2+} ions are removed from the sepiolite structure, whereas the tetrahedral sheets are released forming an amorphous silica gel. The presence of this phase of amorphous silica increases the surface area [18,19]. Several studies have shown the ability for adsorption on acid activated sepiolite [20–22].

Actually, many chemical processes use microwave techniques for heating [32–37]. They offer advantages over conventional methods including a higher heating rate and provide homogeneous heating, notably reducing the treatment time required to reach high surface area values [38–40]. The aim of this work was to compare the effect of microwave-assisted acid treatment on three special clays composed of 2:1 clay minerals. For this study, representative samples of bentonites, sepiolites and kerolitic clays were chosen.

2. Materials and Methods

2.1. Starting Materials

Six special clays (bentonites, kerolitic clays and sepiolites) were tested in this study. The Bentonites were comprised of samples of Bm, composed of montmorillonite (SWy-1, from Crook County, OR, USA), and Bs, composed of saponite (SapCa-2, from Ballarat, CA, USA), both supplied by the Clay Minerals Repository.

The kerolitic clays (samples K1 and K2) were collected from the Esquivias deposit (Madrid Basin, Spain). These "bleaching earths" occur in quarries and outcrops near the Esquivias village, 40 km south of Madrid. In this zone, the Magnesian Unit [6] is about 15 m thick, including up to four beds of kerolite-stevensite mixed layers. The thickness of the single Mg–clay lutite beds ranges from 1 to 4 m (average: 1.5 m) depending on the location in the deposit. The lutites are characterized by their pinkish to pale brown colour and granular to massive lithofacies. Earlier mineralogical research [5,41] showed that they are composed mostly of mixed layered kerolite-stevensite, although the presence of sepiolite and saponite was also reported [6,8,42,43].

Two samples of sepiolite from the Vicálvaro (Sp1) and Batallones (Sp2) deposits, both in the Madrid Basin, were selected. These materials are well known and have been the subject of many studies in recent years [43,44].

2.2. Microwave-Assisted Acid Treatment (MAT)

Five grams of each starting sample were treated with 50 mL of 0.2 N HNO₃ for 4, 8, 12 and 16 min under 800 W of microwave radiation, in an open glass cylinder reactor [39,40]. A commercially available microwave system EMS20100OX (Electrolux, Stockholm, Sweden) operating at 800 W and 2.45 GHz was used for the treatments. The microwave irradiation was applied in a discontinuous form [38,39] to avoid exceeding 100 °C in the system. After 1 min of microwave irradiation, the suspensions were cooled for 5 min at room temperature. Thus, the treatment time is that of irradiation plus the cooling time after every minute of irradiation. After the microwave-assisted acid treatment, the samples were centrifuged and washed with deionized water until the NO₃⁻ ions were completely eliminated (confirmed by a Griess test), and finally, the samples were air-dried. The acid

for the experiments was selected according to Franco et al. [39]. These authors proved that the effect of nitric acid in the structure of several sepiolites was stronger than that of hydrochloric acid.

2.3. Analytical Methodology

Powder patterns were collected on an X'Pert Pro MPD automated diffractometer (PANalytical B.V., Almelo, The Netherlands) equipped with a Ge(111) primary monochromator (strictly monochromatic CuK α_1 radiation) and an X'Celerator detector. The overall measurement time was 33 min per pattern in order to obtain very good statistics over the 2 θ range of 2–65° with 0.017° step size. Clay fraction samples (<2 μ m) of starting samples were prepared from suspensions oriented on glass slides. The clay fraction minerals were identified on oriented air-dried samples, with ethylene glycol solvation, and also after heating at 550 °C.

The percentage of expandable layers in the kerolite/Mg-smectite mixed layers was calculated using the NEWMOD program [5]. The parameters used to model kerolite/Mg-smectite mixed layers were d(001) Mg–Sm = 17.1 Å, d(001), Ke = 9.65 Å, defect free distance = 2, "high N" = 3–7 and reichweite R = 0 (random interstratification).

DRIFT spectra were collected on a Varian 3100 FT-IR spectrophotometer (Varian, Hiroshima, Japan) with a Harrick HVC-DRP cell. We used 200 scans to obtain the interferograms using the spectrum of KBr to calculate the background. About 30 mg of finely ground clay-based materials were placed in the sample holder.

X-ray fluorescence analysis (XRF) (major elements) of untreated and activated samples was performed using the MagiX X-ray fluorescence spectrometer of PANanytical (PANalytical B.V., Almelo, The Netherlands). Sodium was determined with an atomic absorption (AA) spectrophotometer (Varian 220-FS QU-106). Loss on ignition (LOI) was calculated after heating to 950 °C.

 N_2 adsorption-desorption isotherms of the starting and treated samples were measured at -196 °C using a Micromeritic ASAP 2020 apparatus (static volumetric technique) (Micromeritic, Norcross, GA, USA). The total specific surface areas (S_{BET}) were determined through the Brunauer-Emmett-Teller (BET) equation [45]. The total specific surfaces areas (S_{ET}) were determined through the BET equation [46]. The external specific surfaces areas (S_{text}) and micropore areas (S_{tmicro}) were estimated using the t-plot method [45]. One aliquot (0.15 g) of each sample was outgassed for 16 h at 200 °C under vacuum. Nevertheless, taking into account that the thermal treatment affects the sepiolite structure and consequently its S_{BET} , for sepiolite samples two aliquots of each were degassed for 16 h at 120 and 200 °C, respectively. With this methodology we can determine the influence of the thermal treatment on the modification of the real S_{BET} values of starting sepiolites.

Gold-sputtered (10 nm thick) samples were studied by scanning electron microscopy (SEM) using a JEOL SM-6490 LV (Peabody, MA, USA) combined with X-ray energy dispersive spectroscope (EDX).

3. Results and Discussion

3.1. X-Ray Diffraction of the Starting Minerals

3.1.1. Bentonite

The XRD powder patterns of the starting saponite (Bs) and montmorillonite (Bm) are shown in Figure 1a,b. The 001 reflection of the Bs sample shows two components that may be related to the great variability in the hydration degree of the interlayer space. Moreover, the intensity of this doublet is notably higher than the 001 peak of the Bm sample, suggesting a higher ordering along the c^* axis in the Bs sample [47]. The position of 060 reflection agree with the trioctahedral character (d(060) = 1.52 Å) of the 2:1 layers of the Bs sample and the dioctahedral character (d(060) = 1.49 Å) of the Bm sample [47]. On the other hand, the Bm sample has a minor amount of quartz, and the Bs sample has minor amounts of quartz, feldspars and tremolite (Figure 1).



Figure 1. X-ray diffractograms of Bs (**a**), Bm (**b**), K1 (**c**), K2 (**d**), Sp1 (**e**) and Sp2 (**f**). Qtz: Quartz; Fds: Feldspars; Tr: Tremolite; Cal: Calcite.

3.1.2. Kerolitic Clay

Figure 1c,d shows the XRD powder patterns of the starting kerolitic clays (K1 and K2). The XRD profiles show a diffraction profile similar to those previously reported for kerolite-stevensite mixed layers [5]. The starting kerolitic samples show a broad 001 basal reflection towards low angles, as well as symmetrical third-order basal reflections at 3.19 Å. 02,11 and 06,33 reflections at 4.54 and 1.525 Å, respectively, are sharp and strong, whereas 13,20, and 15,24,31 at 2.56 and 1.72 Å, respectively, are very broad. The line broadening effect is consistent with a very small particle size and the turbostatic nature of the layer stacking [8]. Moreover, the b parameter (9.15 Å) taken from a 06 diffraction reflection indicates a trioctahedral structure. Comparison of the XRD patterns of kerolitic clay samples (K1 and K2) with that of saponite (Bs) shows weaker and wider reflections corroborating their higher structural disorder. Following the procedures used by de Vidales et al. and Pozo et al. [5,6], it was identified that K1 has similar proportions of kerolite and stevensite (50%) (type II), while sample K2 has 80% kerolite in the mixed layer (type I).

3.1.3. Sepiolite

The XRD powder patterns of the two starting sepiolites (Sp1 and Sp2) are displayed in Figure 1e,f. Sample Sp1 is almost a pure sepiolite, whereas an almost negligible amount of calcite was found in Sp2. Some differences in the profiles suggest differences in crystallinity and chemical compositions [44]. Small differences in the d-spacing of 110 reflections have been related to slight differences in the crystalline cell, since isomorphic substitution can produce compositional variations [48]. Therefore, the slightly higher d-spacing values of the 110 reflection of Sp2 suggest that this sepiolite sample could have small compositional differences from Sp1.

Reflections at 5.02, 3.00, 2.82, 2.67, 2.61, 2.58, 2.56, and 2.40 Å in both Sp1 and Sp2 are well defined. This suggests that both starting samples are well ordered sepiolites (HCS) [49]. Nevertheless, the lowest FWHM of the 110 reflection corresponding to Sp1 indicates that it has a slightly higher "crystallinity" than Sp2.

3.2. X-Ray Diffraction of the Microwave-Assisted Acid-Treated Samples

3.2.1. Bentonite

Figure 2a,b show the XRD diffraction patterns of the natural smectites (Bs and Bm) and the treated samples in order to identify the structural transformations occurring during the microwave-assisted acid treatment. Two different trends in the behaviour of smectite samples are observed.

The trioctahedral smectite (Bs) (Figure 2a) shows intense changes in the diffraction profile throughout the treatment. These modifications can be detected from the first minutes of the experiment, indicating that saponite has a great reactivity to the conditions in which the acid treatment is performed [40].

The major modifications are observed in the reflections located between 19–22 and $61-63^{\circ}$ 20 (Figure 2a). Reflections located between 19 and 22° 20 of this trioctahedral smectite decrease in intensity at the same time that they are replaced by a broad band between 19 and 30° 20. In addition, reflection 060 widens considerably and decreases until it almost completely disappears after the first 12 min of treatment.

The loss of the ordering along the c^* axis can be observed with the 001 reflection (Figure 3) [47]. The starting Bs sample displays two 001 reflections at 14.8 and 12.6 Å corresponding to the same layered mineral with two different hydration degrees in the interlayer space. After 4 min of MAT treatment, the hydration degree, at least in the interlayer space, is homogenous showing a unique basal spacing at 14.8 Å. Figure 3 shows a dramatic loss of the 001 reflection after 8 min of treatment and the almost complete disappearance after 16 min of treatment. This result seems to indicate that two processes are occurring simultaneously: (1) A progressive loss of the 060 reflection [29,40]. The appearance of a broad band between 19 and 30° 20 suggests the precipitation of a solid residue of amorphous silica [41,50].

On the other hand, the X-ray diffraction patterns of the treated dioctahedral smectite (Bm) (Figure 2b) shows that the reflections located between $19-22^{\circ} 2\theta$ appear almost unchanged even after 16 min of treatment compared with that of the untreated sample. Figure 3 shows that at least 12 min of treatment are necessary to slightly decrease the 001 diffraction peak of Bm (12.4 Å). This result suggests that the structure of the Bm sample shows a great resistance to the MAT treatment, at least in the conditions made for this work [40]. The 001 reflection of Bm increases in the first 4 min of MAT treatment. This suggests an increase of the ordering along the *c** axis, possibly caused by a greater homogeneous hydration of the interlayer space of the sample [47] that occurs when it is submerged in the acid solution.

001

Tr

Qtz

Fds





₀₆₀ a

Bs-16 min

001

020

Qtz

Figure 2. X-ray diffraction patterns of Bs (**a**), Bm (**b**), K1 (**c**), K2 (**d**), Sp1 (**e**) and Sp2 (**f**) and acid-treated samples. Qtz: Quartz; Fds: Feldspars; Tr: Tremolite.



Figure 3. X-ray diffraction patterns of Bs, Bm, K1, K2, Sp1, Sp2 and acid-treated samples.

3.2.2. Kerolitic Clay

Figure 2c,d shows the XRD patterns of kerolite/stevensite mixed layer samples (K1 and K2) and the obtained materials after 4, 8, 12 and 16 min of microwave-assisted acid treatments. In the K1 sample, the acid treatments cause a general decrease of the diffraction reflections, especially the basal one (near 15 Å), a small reflection remaining at 10 Å (Figure 3). On the other hand, the K2 sample (Figure 2d), mainly formed by kerolite (80% in the mixed layer) [5] presents a resistance to acid treatments notably higher than those observed in the K1 sample. In this case, it was observed that in general, most of the diffraction reflections maintain their intensity in the first 12 min of treatment but a slight decrease at 16 min is observed. These results suggest that MAT treatment causes the loss of the stevensite layers (15 Å) [40] in the stacking sequence of the mixed layers, whereas the kerolite layers remain unchanged [5].

3.2.3. Sepiolite

Different behaviour may also be observed in the case of the sepiolite samples Sp1 and Sp2 (Figure 2e,f). The comparison of the X-ray diffraction profiles of the Sp1 sample and the acid-activated materials shows that the microwave-assisted acid treatment causes a progressive decrease of the original diffraction reflections, which is more evident after 8 min of treatment. From this time of treatment, the 110 reflection of Sp1 decreases considerably, remaining a negligible shoulder at 16 min. Figure 2e also shows that reflections located between 20 and 40° 20 broaden and decrease in intensity with the treatment time, while a broad band appears instead between 20 and 30° 20. This result suggests the appearance of an amorphous phase whose proportion increases with the progressive disappearance of sepiolite [39].

On the contrary, in the case of the Sp2 sample, 110 reflection intensity retains almost 50% of its initial intensity even after 16 min of treatment, suggesting that the Sp2 sample appears to be more resistant than Sp1 (Figure 2f) [39]. Considering that both starting sepiolites have similar structural ordering and similar chemical compositions, other factors, such as differences in length and thickness of the fibres can be responsible for this different resistance.

3.3. Fourier Transform Infrared Spectroscopy Study of the Starting Samples and the Microwave-Assisted Acid-Treated Samples

3.3.1. Bentonite

Figure 4a,b shows the FTIR spectra of the starting smectite samples (Bs and Bm) and the material obtained after 16 min of MAT. The OH stretching region (between 3800 and 2800 cm⁻¹) of the Bs sample displays an intense peak at 3676 cm⁻¹ accompanied by a weak shoulder at 3716 cm⁻¹, and a broad band centred at 3616 cm⁻¹. The band at 3676 cm⁻¹ is related to the OH stretching mode of structural Mg(OH)₂ groups in the octahedral sheet of saponite and stevensite structure [50,51]. The band at 3716 cm⁻¹ is assigned to the vOH of isolated Si–OH groups located on the external surface of laminar particles [52], whereas the band at 3616 cm⁻¹ is related to the OH stretching modes of Al(OH)Si groups formed by isomorphic substitutions of Si(IV) by Al(III) in the tetrahedral sheet of saponite [53,54].

In the case of the Bm sample (dioctahedral smectite), Figure 4b shows that the OH stretching band appears at lower wavenumbers compared to those of the Bs sample (trioctahedral smectites). The OH region of the FTIR spectrum (3800–2800 cm⁻¹) is dominated by an intense band at 3646 cm⁻¹ assigned to the Al(OH)Al-stretching vibrations [55]. A small shoulder can be observed at 3695 cm⁻¹, related to the presence of pyrophyllite-like local structural fragments [55] or to a small amount of Mg in the octahedral sheet of montmorillonite [40].

Figure 4a,b also shows the Si–O stretching region of the FTIR spectra of the Bs and Bm samples. In both cases, the wide bands centred at 1268 (for the Bs sample) and 1275 cm⁻¹ (for the Bm sample) are related to the Si–O stretching band [56] and the band at 1627 cm⁻¹ (for both samples) are related to the water bending mode [57]. The FTIR profile of Bm sample suggest the existence of a negative band centred at 1070 cm⁻¹. This anomalous feature is related to the absorption and specular reflectance that currently give the DRIFT technique [58]. This can be minimized by a mechanical treatment of the sample.

The OH stretching region $(3800-2800 \text{ cm}^{-1})$ of Figure 4a shows that, in the case of saponite, the 3676 cm⁻¹ band disappears completely within the MAT treatment and is replaced by a band at 3739 cm⁻¹. This is commonly observed in the FTIR spectra of the acid activated smectites and is usually related to the OH stretching mode of Si–OH groups [25] while bands at 3633 cm⁻¹ broaden progressively and shift to 3680 cm⁻¹. After the acid treatment, three changes in the FTIR spectra are observed: (1) The water bending band of the Bs sample shifts from 1627 to 1653 cm⁻¹, suggesting a change in the local environment of the adsorbed water molecules or the presence of acid water (H₃O⁺) whose bending band appears near 1700 cm⁻¹, (2) A decrease in intensity of the wide band at 1268 cm⁻¹ (3) The appearance of a new set of bands in the range 1516–1300 cm⁻¹, which are classically related to the Si–O–Si stretching bands of low "crystallinity" phases, mainly amorphous silica, originated in the treatment [59,60].

On the contrary, no significant changes could be observed in the FTIR spectra of the Bm samples after the microwave-assisted acid treatment (Figure 4b) corroborating that the structural transformations observed in saponite do not take place in montmorillonite.

Although the best technique to study the OH group of these clay minerals is through the deuteration of these species, it is very important to highlight the complementarity of the actual FTIR spectra and XRD profile of the minerals and the obtained materials. There is a direct relationship between the modifications observed in the X-ray diffraction patterns and the FTIR spectra. This complementarity suggests the loss of the layered structure of the Bs and Bm samples simultaneously with the dissolution of the OH groups of the octahedral layers and the consequent formation of an amorphous silica phase [40].



Figure 4. Comparison of the FTIR spectra of saponite Bs (**a**), montmorillonite Bm (**b**), kerolite K1 (**c**), kerolite K2 (**d**), sepiolite Sp1 (**e**) and sepiolite Sp2 (**f**) and the material obtained after 16 min of MAT.

3.3.2. Kerolitic Clay

Figure 4c,d represent the FTIR spectra of starting kerolitic clay samples. The $3800-2800 \text{ cm}^{-1}$ region of the starting samples shows a narrow band centred at 3677 cm^{-1} and a small shoulder at 3618 cm^{-1} . They are classically ascribed to the OH stretching bands of Mg(OH)₂ groups located in the octahedral sheet of talc structure [50,51], whereas the small shoulder centred at 3618 cm^{-1} is related to the presence of a small proportion of Al in the octahedral sheets [55]. Two wider bands, related to the OH-stretching modes of water molecules [58] located in the interlayer space of the stevensite phase and at the outer surface of the particles, are centred at $3438 \text{ and } 3244 \text{ cm}^{-1}$.

Figure 4c,d also show, between 1800 and 800 cm⁻¹, the Si–O stretching region of the FTIR spectra of the starting samples. In this region the more intense band appears centred at 1015 cm⁻¹. This is ascribed to the Si–O stretching modes of trioctahedral phyllosilicates [57].

A comparison of the FTIR spectra of starting K1 and K2 with the 16-min treated samples shows that the acid treatments cause a progressive decrease of the band centred at 3672 cm^{-1} , assigned to stretching modes of Mg–OH groups, as well as a generalized decrease of the stretching modes of the water molecules (3438 and 3244 cm⁻¹ bands). These can be related to the loss of the OH groups during the dissolution of the Mg²⁺ cations of the octahedral sheets but also to the loss of water molecules [40]. Moreover, simultaneously to the loss of OH groups within the octahedral sheet, a strong decrease in the intensity of the Si–O stretching band (1018 cm⁻¹) is observed. The occurrence of a new band between 1304 and 1082 cm⁻¹, is related to amorphous silica segregation [60,61].

3.3.3. Sepiolite

The FTIR spectra of the starting sepiolite samples are shown in Figure 4e,f. According to Frost [61], between 3800 and 2800 cm⁻¹, sepiolite shows two bands at 3685, and 3625 cm⁻¹ associated with the stretching vibration of the Mg–OH groups and a band at 3561 cm⁻¹ assigned to the stretching modes of water molecules coordinated with magnesium. Moreover, two bands at 3355 and 3233 cm⁻¹ are attributed to zeolitic and coordinated water [62].

Previous studies show that several types of water are present in sepiolite [50]. The best spectroscopic method for studying water is to examine the H–O–H bending mode band centred at 1650 cm⁻¹ since this band does not overlap other bands of sepiolite. In this region of the FTIR spectra, two major bands are observed at 1660 and 1629 cm⁻¹ in the starting sepiolites. According to Farmer [60], these bands can be attributed to the zeolitic and adsorbed water, respectively, whereas bands between 1225 and 1000 cm⁻¹ are assigned to Si–O stretching modes [61]. A band at 1470 cm⁻¹ in the FTIR spectrum of Sp2 (Figure 4f) is attributed to the (CO₃)^{2–} antisymmetric and symmetric stretching modes, which indicates the presence of carbonate.

Figure 4e,f also shows the FTIR spectra of the 16 min treated sepiolite samples. These FTIR spectra reveal the almost disappearance of the OH stretching bands at 3685, 3625 and 3561 cm⁻¹ in Sp1, whereas in Sp2 the FTIR spectra show that these bands decrease in intensity but remain clearly visible in the profile. This suggests the complete loss of the OH groups along the MAT treatment of Sp1 while this loss is only partial in the case of Sp2. The FTIR spectra of acid-treated sepiolite also show the appearance of a small band at 3733 cm⁻¹, which is commonly observed in acid-treated minerals and is related to the Si–OH groups [25].

Figure 4e,f also shows that the band of the zeolitic water (1660 cm^{-1}) decreases in intensity along the MAT treatment while the band related to adsorbed water (1629 cm^{-1}) increases, suggesting the evolution of water molecules from sepiolite channels that are lost in the structural changes caused by the dissolution of octahedral sheets.

Figure 4e,f shows strong changes in the bands related to the bands of the Si–O stretching modes. These bands are partially replaced with wider bands located between 1319 and 1085 cm⁻¹. These new bands are classically related to amorphous silica [59,60].

Moreover, the FTIR of Sp2 at 2 min of MAT treatments (not shown in Figure 4), shows the complete disappearance of the band at 1470 cm^{-1} , suggesting that the dissolution of carbonates occurs at the beginning of the treatment.

3.4. Evolution of the Chemical Composition

3.4.1. Bentonite

Table 1 shows the chemical analysis of starting samples and obtained materials after 16 min of MAT, whereas the variation of MgO/SiO₂, Al₂O₃/SiO₂ and Fe₂O₃/SiO₂ weight ratios of chemical composition of the materials obtained through the microwave-assisted acid treatment are shown in Figure 5. This graph corroborates the different behaviour of smectites with the microwave-assisted acid treatment according to the composition of the octahedral sheet. An intense dissolution of Mg²⁺ ions of the octahedral sheet in the trioctahedral smectite (Bs) is observed. In this case, the dissolution of 32% of Mg²⁺ of octahedral sheet occurs in the first 8 min of treatment and an additional 45.2% is lost at 16 min of treatment. In the case of dioctahedral smectite (Bm), the amount of this element in the octahedral sheet is notably smaller compared with that of the Bs sample (Figure 5a), and the loss of Mg²⁺ is notably smaller; 12% in the first 8 min and an additional 2% at 16 min.

Table 1. Chemical analysis of starting samples and obtained materials after 16 min of MAT.

Sample	% SiO ₂	$% Al_2O_3$	% MgO	% Fe ₂ O ₃	$\% Na_2O$	% K ₂ O	% CaO	% MnO	% P ₂ O ₅	% TiO ₂
Bs	53.73	4.82	30.08	1.03	2.76	0.55	5.34	0.03	0.11	0.53
Bs-16 min	77.11	2.51	16.49	0.98	0.00	0.71	5.59	0.02	0.03	0.83
Bm	65.75	22.61	3.21	3.96	1.67	0.60	1.65	0.02	0.09	0.15
Bm-16 min	68.52	23.30	2.85	3.97	0.30	0.47	0.20	0.01	0.04	0.15
K1	53.94	2.66	25.26	0.77	0.09	0.59	0.32	0.01	0.05	0.15
K1-16 min	67.75	2.22	16.24	0.88	0.07	0.64	0.12	0.02	0.04	0.19
K2	55.00	0.70	27.70	0.20	0.03	0.15	0.20	0.01	0.04	0.10
K2-16 min	66.49	0.45	17.84	0.23	0.03	0.13	0.23	0.01	0.03	0.09
Sp1	56.89	0.69	23.97	0.02	0.01	0.20	0.59	0.05	0.04	0.03
Sp1-16 min	81.77	0.59	2.40	0.03	0.02	0.23	0.60	0.04	0.04	0.08
Sp2	56.41	1.56	22.39	0.44	0.03	0.22	0.82	0.05	0.04	0.22
Sp2-16 min	70.48	1.18	11.57	0.36	0.04	0.21	0.06	0.05	0.04	0.07



Figure 5. Variation of MgO/SiO₂ (**a**), Al_2O_3/SiO_2 (**b**), and Fe_2O_3/SiO_2 (**c**) along the experiments: Bm (black triangle), Bs (black square), K1 (red square), K1 (red triangle), Sp1 (blue square), Sp2 (blue triangle).

The Al^{3+} cations of the octahedral sheet of Bs and Bm appear to be more resistant to the dissolution with the microwave-assisted acid treatment. The amount of Al^{3+} in the Bs sample represents 26% of the Al content in the Bm sample. In this case, the proportion of Al^{3+} in Bm remains almost unchanged

along the entire treatment, whereas in Bs, the dissolution of 24% of Al^{3+} of the octahedral sheet occurs in the first 8 min and an additional loss of 47.93% at 16 min of treatment possibly due to the crumbling of the octahedral sheets of 2:1 layers in saponite. On the contrary, the Fe³⁺ amounts remain almost unchanged along the treatment in both di- and trioctahedral smectite samples.

3.4.2. Kerolitic Clay

The chemical composition of the kerolitic clay samples indicates that starting samples are formed mostly by trioctahedral clay minerals with Mg as predominant cation in the octahedral layer (MgO, 25–28 wt %). The detection of a small amount of Al, Fe and K should be mostly related to detrital minerals (feldspar, illite-mica) identified in the mineralogical characterization as well as to minor discrete saponite commonly associated with kerolite-stevensite mixed-layer deposits. The low content of detrital minerals or other Mg-smectites, other than stevensite in the mixed layer, explains the low content of Al, Fe and K in the kerolite-rich sample (K2).

In the case of kerolitic clay samples, a large decrease in the Mg content relative to silica is shown. This can be explained by the dissolution of part of the Mg^{2+} located in the octahedral sheet of the kerolitic clay. This process is simultaneous accompanied by the formation of an amorphous silica phase. The samples show a similar Mg^{2+} loss rate up to 8 min of treatment but show a slightly different behaviour above this time of treatment (Figure 5). In the first 8 min, K1 sample lost 28% of the initial amount of Mg^{2+} and K2 sample lost 29% of this cation. From 8 to 16 min K1 sample lost 21% of its Mg^{2+} cations whereas K2 sample lost 19%. On the contrary, Al^{3+} and the scarce amounts of Fe^{3+} are released at lower rates (Figure 5). Mg^{2+} loss rates are significantly lower than those observed in the case of saponite. This result suggests that the presence of kerolite layers in the mixed layers decreases the reactivity of these materials against the acid treatments. In these samples the small amounts of Al^{3+} and Fe^{3+} remain almost constant along the entire treatment.

3.4.3. Sepiolite

The variation of the octahedral cation percentages in sepiolite samples are also shown in Figure 5. This figure reveals that MAT causes the release of Mg^{2+} ions of the octahedral sheet, whereas the amount of Si remains in the solid. Mg^{2+} ion dissolution rates are similar in the two sepiolites in the first 8 min of treatment (31% for Sp1 and 26% for Sp2), whereas it is higher in Sp1 (62%) than Sp2 (33%) between 8 and 12 min. In the Sp1 sample, Mg^{2+} cations disappear almost completely after 16 min of treatment. According to the observed decrease of the 110 reflection in these sepiolite samples (Figure 1c), this result suggests a direct relation between the dissolution of Mg^{2+} cations and the loss of the sepiolite structure. Similar to that observed in kerolitic samples, the small amounts of Al^{3+} and Fe^{3+} observed in sepiolite remains almost constant along the entire treatment.

3.5. SEM Examination

SEM micrographs of starting clays and the obtained materials after 16 min of microwave-assisted acid treatment are shown in Figures 6-11.

3.5.1. Bentonite

The individual particles of starting montmorillonite (Bm) display laminar morphologies with particle sizes smaller than 5 μ m, which appears to be associated in soft aggregates of irregular morphologies (Figure 6a). After 16 min of microwave-assisted acid treatment the individual particles show no significant modification in the shape and distribution of sizes compared to those of the natural montmorillonite (Figure 6b).



Figure 6. SEM images of Bm (a), Bm after 16 min of microwave-assisted acid treatment (b).



Figure 7. Cont.



Figure 7. SEM images of Bs (a) and Bs after 16 min of microwave-assisted acid treatment (b).

The Bs sample shows aggregates similar dimensions than those of the Bm sample. These aggregates are formed by extremely thin lamellar particles with larger dimensions along the (001) faces and with very jagged contours (Figure 7a). Contrary to that observed in Bm sample, the microwave-assisted acid treatment of Bs sample (Figure 7b) strongly affects the morphological features of the saponite particles and its aggregates. This treatment causes the size of the aggregates and their individual particles to decrease.



Figure 8. Cont.



Figure 8. SEM images of the kerolite K1 (**a**) and kerolite K1 after 16 min of microwave-assisted acid treatment (**b**).

3.5.2. Kerolitic Clay

Regarding kerolitic clays, the individual particles of the starting K1 sample exhibit extremely thin lamellar morphologies that appear to be, in general, smaller than those observed in the saponite sample (S) (Figure 8a). The sizes of the higher proportion of the individual particles are smaller than 10 μ m. These individual particles form highly porous aggregates with sizes ranging between 20 and 50 μ m. Similar observations can be made in the SEM micrograph of K2 particles shown in Figure 9a.



Figure 9. Cont.



Figure 9. SEM images of the kerolite K2 (**a**) and kerolite K2 after 16 min of microwave-assisted acid treatment (**b**).

Figures 8 and 9 also show SEM micrographs of the 16-min acid-activated K1 and K2 samples. A detailed comparison of Figure 8a,b suggests that no significant decrease in the size of the porous aggregates occurs after 16 min of MAT. However, some textural changes in the external surface of aggregates suggest changes in their porosity. On the contrary, the comparison of Figure 9a,b suggests a slight decrease in the size of the porous aggregates of K2 sample. Nevertheless, the proportion of clay-sized particles (<2 μ m) appears to be unchanged during the treatment in both K1 and K2 samples.



Figure 10. Cont.



Figure 10. SEM images of the sepiolite Sp1 (**a**), sepiolite Sp1 after 16 min of microwave-assisted acid treatment (**b**).

3.5.3. Sepiolite

The starting sepiolites (Sp1 and Sp2) show their typical fibrous habit (Figures 10a and 11a). In the Sp1 sample the individual fibres are commonly longer than 6 μ m (Figure 10a), whereas in Sp2, the main population consists of fibres shorter than 2 μ m (Figure 11a). Moreover, the diameter of the fibres seems remarkably wider in Sp2. These fibres commonly form bundles consisting of parallel or tangled fibres [59]. The two starting sepiolites also differ regarding size and morphology of their aggregates. In general, the largest aggregates are found in the Sp1 sample (Figure 10a) and the smallest in Sp2 (Figure 11a).



Figure 11. Cont.



Figure 11. SEM images of the sepiolite Sp2 (**a**) and sepiolite Sp2 after 16 min of microwave-assisted acid treatment (**b**).

The SEM images of the acid activated sepiolites show a slight disaggregation of the blocky aggregates (bundles) in the first minutes (not shown), followed by a sharp decrease in the length of aggregates as well as in the length of the individual fibres. Finally, after 16 min of MAT, the fibres show a very poorly defined morphology (Figures 10b and 11b). In the Sp1 sample, 16 min of MAT causes the almost disappearance of fibrous particles (Figure 10b), whereas some small fibres, smaller than 1 μ m, remain in Sp2, coexisting with the new irregular particles (Figure 11b).

3.6. Textural Properties Evolution

The textural properties of the starting clays and the obtained materials after 16 min of microwave-assisted acid treatment were determined by N₂ adsorption-desorption at -196 °C. The variation of the specific surface area (SSA), determined by the BET equation [45] of the samples throughout the treatment time is shown in Figure 12. The SSA values of the bentonites are notably lower than those of kerolitic clay and sepiolite samples, suggesting that the initial SSA values are closely related to the observed particle sizes in the untreated samples.

3.6.1. Bentonite

Saponite (Bs) and montmorillonite (Bm) samples have SSA values of 128 and 49 m²/g, respectively (Table 2). In the Bm sample the SSA values increase very little ($\Delta S_{BET} = 13 \text{ m}^2/\text{g}$; $\%\Delta S_{BET} = 27$) during the treatment time. Nevertheless, the Bs samples show a noticeable increase in their SSA values compared to montmorillonite, at 8 min reaching an SSA value of 324 m²/g, while from the treatment time the SSA appears to be invariable, suggesting that it has achieved the maximum activation ($\Delta S_{BET} = 198 \text{ m}^2/\text{g}$; $\%\Delta S_{BET} = 155$). Contrary to that observed in Bm sample, the great particle size reductions observed during microwave-assisted acid treatment of saponite and an increase in the porosity produce a noticeable increase in their SSA values.



Figure 12. SSA evolution of the 2:1 clay samples studied vs. treatment time: Bm (black triangle), Bs (black square), K1 (red square), K1 (red triangle), Sp1 (blue square), Sp2 (blue triangle).

	SBET	St _{ext}	St _{micro}		SBET	St _{ext}	St _{micro}
Sample	$(m^2 \cdot g^{-1})$	$(m^2 \cdot g^{-1})$	$(m^2 \cdot g^{-1})$	Sample	$(m^2 \cdot g^{-1})$	$(m^2 \cdot g^{-1})$	$(m^2 \cdot g^{-1})$
Degasification Temperature		200 °C		Degasification Temperature		200 °C	
Μ	49	25	24	K1	265	101	163
M-4 min	49	22	27	K1-4 min	337	139	198
M-8 min	56	28	28	K1-8 min	363	155	209
M12 min	61	32	29	K1-12 min	393	159	234
M-16 min	62	33	29	K1-16 min	426	195	230
S	128	76	52	K2	279	167	112
S-4 min	257	128	129	K2-4 min	339	236	104
S-8 min	324	138	186	K2-8 min	367	241	126
S-12 min	331	114	217	K2-12 min	369	248	121
S-16 min	326	106	221	K2-16 min	435	266	170
	S _{BET}	St _{ext}	St _{micro}	S _{BET}	Stext	St _{micro}	
Sample	$(m^2 \cdot g^{-1})$	$(m^2 \cdot g^{-1})$	$(m^2 \cdot g^{-1})$	$(m^2 \cdot g^{-1})$	$(m^2 \cdot g^{-1})$	$(m^2 \cdot g^{-1})$	Best S_{BET}
Degasification Temperature	200 °C				(m-·g -)		
Sp1	117	102	15	288	117	171	288
Sp1-2 min	175	132	43	327	161	166	327
Sp1-4 min	244	150	94	358	173	185	358
Sp1-8 min	378	191	187	406	222	184	406
Sp1-16 min	519	437	82	437	348	89	519
Sp2	133	111	22	320	127	193	320
Sp2-2 min	154	133	21	328	170	158	328
Sp2-4 min	323	187	136	280	174	106	323
Sp2-8 min	371	174	197	371	176	141	371
Sp2-16 min	438	197	241	396	205	191	438

Table 2. Specific surface areas obtained after outgassing at 200 $^{\circ}$ C or 120 $^{\circ}$ C of starting 2:1 clay minerals and obtained materials after acid treatment.

3.6.2. Kerolitic Clay

The SSA values of the starting kerolitic clays are significantly higher than other magnesian clay minerals such as saponite or sepiolite [39,40]. The SSA values of the starting kerolitic clays (K1 and K2) are very similar (265, and 279 m²/g, respectively) values in agreement with those reported by other authors [5]. Both samples have a similar trend during the entire treatment showing the higher SSA values in the K2 sample. According to Table 2 the greatest contribution to the SSA comes from the micropores in the K1 sample, but from the external surface of the particles in K2. This result agrees with the observed modifications in the SEM images (Figures 8 and 9). In both cases, the MAT treatment causes similar increments of SSA. In the case of K1, SSA increases from 265 to 426 m²/g ($\Delta S_{BET} = 161 \text{ m}^2/\text{g}$; $\%\Delta S_{BET} = 61$), and in the case of K2 from 279 to 435 m²/g ($\Delta S_{BET} = 156 \text{ m}^2/\text{g}$; $\%\Delta S_{BET} = 56$).

3.6.3. Sepiolite

In order to obtain SSA values close to reality, samples must be properly degassed without affect the structure of sepiolite [59]. For this reason, to find the appropriate conditions for the analysis of the samples, starting and treated sepiolite samples were kept for 16 h at two different temperatures (120 and 200 °C). Table 2 shows that SSA values are notably influenced by the temperature used to degas the sample. When samples are degassed at 200 °C, the two starting sepiolites exhibit a similar external surface area (St_{ext}), a similar and very small microporous area (St_{micro}), and, therefore, a similar BET surface area (S_{BET}). Thus, S_{BET} values are notably smaller than those reported in the literature [49,51]. Nevertheless, the obtained St_{micro} and S_{BET} values of natural sepiolites are notably higher when samples are degassed at 120 °C. These results suggest that the surface area from samples degassed at 200 °C is influenced by some textural or structural modifications [60], which strongly affect the microporous surface area (Table 2).

On the contrary, in the 8- and 16-min-treated sepiolite samples, in general the St_{micro} and S_{BET} values are notably higher when these samples are degassed at 200 °C. This result suggests that the SSA values of these treated sepiolite samples are not affected by any structural transformation. In acid-treated sepiolite, a higher temperature allows the complete degassing of the sample and therefore the obtaining of SSA values closer to reality [39].

On the other hand, the evolution of St_{micro} and St_{ext} of bentonite and kerolite samples after the MAT does not show erratic behaviour. For this reason, we only analysed the samples after degas at the same temperature (200 °C). These values can be slightly enhanced by a thermal activation but the low SSA value observed in the starting and treated-Bm samples shows that the thermal activation at 200 °C has a very small influence in the SSA values.

The starting sepiolite samples (Sp1 and Sp2) show different SSA values (288 and 320 m²/g, respectively). In the first 8 min, the SSA value in Sp1 increases by 41% (from 288 to 406 m²/g), increasing an additional 39.6% at 16 min of treatment (from 406 to 519 m²/g). Regarding the Sp2 sample, an increase of only 16% is observed in the first 8 min of treatment (from 320 to 371 m²/g) followed by an additional increase of 21% at 16 min (from 371 to 438 m²/g). In the case of Sp1, SSA increases from 288 to 519 m²/g ($\Delta S_{BET} = 231 \text{ m}^2/\text{g}$; $\Delta \Delta S_{BET} = 80$), and in the case of K2 from 320 to 438 m²/g ($\Delta S_{BET} = 118 \text{ m}^2/\text{g}$; $\Delta \Delta S_{BET} = 37$). The larger increase in SSA in the Sp1 sample appears to be related to the higher textural modifications shown in the SEM micrographs (Figures 10 and 11). In both cases, with the assistance of microwaves during the acid treatment, it is possible to obtain, after 16 min of treatment, materials with an SSA value equivalent to those obtained after 48 h with conventional heating methods [17,29].

For an appropriate comparison, Table 3 summarizes the main modifications obtained with the different clay mineral studied after 16 min of microwave-assisted acid treatment. This table shows the increments of SSA (ΔS_{BET} ; $\Delta \Delta S_{BET}$) as the difference between the S_{BET} of the 16 min acid-treated sample and the corresponding starting clay. This table shows that higher increments are obtained with Sp1 and Bs samples that also display a higher loss of octahedral cations. It is interesting to note that both kerolitic clay materials have a similar loss of octahedral cations and similar S_{BET} increments.

This result is not consistent with the different structural modification observed by XRD. This result suggests that, contrary to what was observed with saponite and stevensite, the layered structure of kerolite is roughly maintained despite the loss of about 36% of the octahedral cations.

Type of Clay	Sample	Main Characteristic	XRD Modification	FTIR Amorphous Si–O Bands	% Octahedral Cations Loss	ΔS_{BET} (m ² /g)	ΔS_{BET}
Bentonite -	Bs	trioctahedral smectite	020 and 060 reflections disappears	YES	71	198	155
	Bm	dioctahedral smectite	020 and 060 reflections appears to be unaffected	NO	3	13	27
Kerolitic clay	K1	stevensite > kerolite	020 and 060 reflections strongly decreased	YES	49	161	61
	K2	stevensite < kerolite	020 and 060 reflections slightly decreased	YES	47	156	56
Sepiolite _	Sp1	fibres: thinner, longer	reflections between 18 and 32° 2θ disappear and replaced with a broad band	YES	93	231	80
	Sp2	fibres: wider, shorter	reflections between 18 and 32° 2θ slightly decrease	YES	60	118	37

Table 3. Principal modifications obtained after 16 min of microwave-assisted acid treatment (MAT).

4. Conclusions

This study shows the effectiveness of microwave-assisted acid treatment in several 2:1 clay minerals. With these treatments, extremely high specific surface area (SSA) increases are achieved in just 16 min. The highest SSA increments were obtained with a sepiolite (up to $231 \text{ m}^2/\text{g}$ in 16 min). Nevertheless, notably higher SSA increments were obtained with Mg–smectites (up to $198 \text{ m}^2/\text{g}$ in 16 min) and kerolitic clays (up to $161 \text{ m}^2/\text{g}$). This is a significant breakthrough compared to conventional acid treatments that reached smaller increments after 48 h of treatment. The increase in the specific surface area obtained with the microwave-assisted acid treatment can be used as an indicator of the extension of the activation process or the reactivity of the 2:1 clay minerals to this type of treatment. This reactivity appears to depend on several factors including octahedral sheet compositions and the structure of the specific 2:1 clay minerals.

Author Contributions: M.P. provided the samples for this work. F.F. and M.P. conceived and designed the experimental research and reviewed background research. F.F. supervised data acquisition and compilation. L.P. and J.A.C. carried out the experiments, data reduction and calculations. E.B. carried out the X-ray fluorescence analysis. All authors contributed to the interpretation of the results as well as writing and editing the paper. The authors have approved the submitted version and agree to be personally accountable for their own contributions and for ensuring that questions related to the accuracy or integrity of any part of the work, even one in which the author was not personally involved, are appropriately investigated, resolved, and documented in the literature.

Acknowledgments: The authors are very grateful for the inestimable help of the reviewers who have significantly improved the article and to Petra Gannholm for the careful review of English. The authors also wish to thank the project BIA2017-82391-R for the financial support. The preparation of this paper is included within the activities of the RNM-199 (Universidad de Málaga) and C-144 (Universidad Autónoma de Madrid, Geomaterials and Geological Processes) Research Groups.

Conflicts of Interest: The authors declare no conflict of interest. The founding sponsors had no role in the design of the study; in the collection, analyses, or interpretation of data; in the writing of the manuscript, and in the decision to publish the results.

References

 Guggenheim, S.; Adams, J.M.; Bain, D.C.; Bergaya, F.; Brigatti, M.F.; Drits, V.A.; Formoso, M.L.L.; Galán, E.; Kogure, T.; Stanjek, H. Summary of recommendations of nomenclature committees relevant to clay mineralogy: Report of the Association Internationale pour l'Etude des Argiles (AIPEA) Nomenclature Committee for 2006. *Clays Clay Miner.* 2006, *54*, 761–772. [CrossRef]

- Guggenheim, S.; Adams, J.M.; Bain, D.C.; Bergaya, F.; Brigatti, M.F.; Drits, V.A.; Formoso, M.L.L.; Galán, E.; Kogure, T.; Stanjek, H. Summary of recommendations of nomenclature committees relevant to clay mineralogy: Report of the Association Internationale pour l'Etude des Argiles (AIPEA) Nomenclature Committee for 2006. *Clay Miner.* 2006, *41*, 863–878. [CrossRef]
- Brindley, G.W.; Brown, G. X-ray diffraction procedures for clay mineral identification. In *Crystal Structures of Clay Minerals and Their X-Ray Identification*; Brindley, G.W., Brown, G., Eds.; Mineralogical Society: London, UK, 1980; pp. 305–356, ISBN 9780903056373.
- 4. Eberl, D.D.; Jones, B.F.; Khoury, H.N. Mixed layer kerolite-stevensite from the Amargosa Desert, Nevada. *Clays Clay Miner.* **1982**, *30*, 321–326. [CrossRef]
- 5. Martin de Vidales, J.L.; Pozo, M.; Alia, J.M.; Garcia Navarro, F.; Rull, F. Kerolite-stevensite mixed-layers from the Madrid Basin, Central Spain. *Clay Miner.* **1991**, *26*, 329–342. [CrossRef]
- 6. Pozo, M.; Casas, J. Origin of kerolite and associated Mg clays in palustrine-lacustrine environments. The Esquivias deposit (Neogene Madrid Basin, Spain). *Clay Miner.* **1999**, *34*, 395–418. [CrossRef]
- Dekov, V.M.; Cuadros, J.; Shanks, W.C.; Koski, R.A. Deposition of talc-kerolite-smectite-smectite at seafloor hydrothermal vent fields: Evidence from mineralogical, geochemical and oxygen isotope studies. *Chem. Geol.* 2008, 247, 171–194. [CrossRef]
- 8. Steudel, A.; Friedrich, F.; Schuhmann, R.; Friedrich, R.; Ruf, U.; Sohling, U.; Emmerich, K. Characterization of a fine-grained interstratification of turbostratic talc and saponite. *Minerals* **2017**, *7*, 5. [CrossRef]
- 9. Miller, A.K.; Guggenheim, S.; Van Groos, A.F.K. Bond energy of adsorbed and interlayer water kerolite dehydration at elevated pressures. *Clays Clay Miner*. **1991**, *39*, 127–130. [CrossRef]
- Brindley, G.W.; Sempels, R.E. Preparation and properties of some hydroxy-aluminium beidellites. *Clay Miner*. 1977, 12, 229–237. [CrossRef]
- 11. Murray, H.H. Traditional and new applications for kaolin, smectite, and palygorskite: A general overview. *Appl. Clay Sci.* **2000**, *17*, 207–221. [CrossRef]
- 12. González-Pradas, E.; Socias-Viciana, M.; Saifi, M.; Ureña-Amate, M.D.; Flores-Cespedes, F.; Fernández-Pérez, M.; Villafrenca-Sánchez, M. Adsorption of chloridazon from aqueous solution on heat and acid treated sepiolites. *Water Res.* **2005**, *39*, 1849–1857. [CrossRef] [PubMed]
- 13. Tan, Ö.; Yilmaz, L.; Zaimoglu, A.S. Variation of some engineering properties of clays with heat treatment. *Mater. Lett.* **2004**, *58*, 1176–1179. [CrossRef]
- 14. Kara, M.; Yuzer, H.; Sabah, E.; Celik, M.S. Adsorption of cobalt from aqueous solutions onto sepiolite. *Water Res.* **2003**, *37*, 224–232. [CrossRef]
- 15. Dekany, I.; Turi, L.; Fonseca, A.; Nagy, B.J. The structure of acid treated sepiolites: Small-angle X-ray scattering and multi MAS-NMR investigations. *Appl. Clay Sci.* **1999**, *14*, 141–160. [CrossRef]
- 16. Balci, S. Thermal decomposition of sepiolite and variations in pore structure with and without acid pre-treatment. *J. Chem. Technol. Biotechnol.* **1996**, *66*, 72–78. [CrossRef]
- Jiménez-López, A.; López-González, J.D.; Ramírez-Sáenz, A.; Rodríguez-Reinoso, F.; Valenzuela-Calahorro, C.; Zurita-Herrera, L. Evolution of surface area in a sepiolite as a function of acid and heat treatments. *Clay Miner*. 1978, 13, 375–385. [CrossRef]
- 18. Inukai, K.; Miyawaki, R.; Tomura, S.; Shimosaka, K.; Irkec, T. Purification of Turkish sepiolite through hydrochloric acid treatment. *Appl. Clay Sci.* **1994**, *9*, 11–29. [CrossRef]
- 19. Aznar, A.J.; Gutiérrez, E.; Díaz, P.; Álvarez, A.; Poncelet, G. Silica from sepiolite: Preparation, textural properties, and use as support to catalyst. *Microporous Mater.* **1996**, *6*, 105–114. [CrossRef]
- 20. Sabah, E.; Turan, M.; Çelik, M.S. Adsorption mechanism of canonic surfactants onto acid and heat-activated sepiolites. *Water Res.* **2002**, *36*, 3957–3964. [CrossRef]
- 21. Vico, L.I. Acid-base behaviour and Cu²⁺ and Zn²⁺ complexation properties of the sepiolite/water interface. *Chem. Geol.* **2003**, *198*, 213–222. [CrossRef]
- 22. Kendall, T. Smectite Clays. In *Industrial Clays;* Kendall, T., Ed.; Industrial Minerals Information Ltd.: London, UK, 1996; pp. 1–12.
- 23. Čičel, B.; Komadel, P. Structural formulae of layer silicates. In *Quantitative Methods in Soil Mineralogy*; Luxmoore, R.J., Ed.; Soil Science Society of America, Inc.: Madison, WI, USA, 1994; pp. 114–136.
- 24. Kaviratna, H.; Pinnavaia, T.J. Acid hydrolysis of octahedral Mg²⁺ Sites in 2:1 layered silicates: An assessment of edge attack and gallery access mechanisms. *Clays Clay Miner.* **1994**, *42*, 717–723. [CrossRef]

- 25. Komadel, P.; Madejová, J. Acid activation of clay minerals. In *Handbook of Clay Science*; Bergaya, F., Theng, B.K.G., Lagaly, G., Eds.; Elsevier: Amsterdam, The Netherlands, 2006; Volume 1, pp. 263–287.
- 26. Rice, N.M.; Strong, L.W. The leaching of lateritic nickel ores in hydrochloric acid. *Can. Metall. Quart.* **1974**, *b13*, 485–493. [CrossRef]
- 27. Komadel, P.; Schmidt, D.; Madejová, J.; Čičel, B. Alteration of smectites by treatments with hydrochloric acid and sodium carbonate solutions. *Appl. Clay Sci.* **1990**, *5*, 113–122. [CrossRef]
- 28. Komadel, P.; Bujdák, J.; Madejová, J.; Šucha, V.; Elsass, F. Effect of non-swelling layers on the dissolution of reduced-charge montmorillonite in hydrochloric acid. *Clay Miner.* **1996**, *31*, 333–345. [CrossRef]
- 29. Vicente-Rodríguez, M.A.; López González, J.D.; Bañares-Muñoz, M.A. Preparation of microporous solids by acid treatment of a saponite. *Microporous Mater.* **1995**, *4*, 251–264. [CrossRef]
- 30. Tkáč, I.; Komadel, P.; Müller, D. Acid-treated montmorillonites-A study by ²⁹Si and ²⁷Al MAS NMR. *Clay Miner.* **1994**, *29*, 11–19. [CrossRef]
- 31. Breen, C.; Madejová, J.; Komadel, P. Characterization of moderately acid-treated, size-fractionated montmorillonites using IR and MAS NMR spectroscopy and thermal analysis. *J. Mater. Chem.* **1995**, *5*, 469–474. [CrossRef]
- 32. Adams, J.M.; McCabe, R.W. Clay Minerals as Catalysts. In *Handbook of Clay Science*; Bergaya, F., Theng, B.K.G., Lagaly, G., Eds.; Elsevier: Amsterdam, The Netherlands, 2006; Volume 1, pp. 541–581.
- 33. García Carmona, J.; Rodriguez Clemente, R.; Gómez Morales, J. Comparative preparation of microporous VPI-5 using conventional and microwave heating techniques. *Zeolites* **1997**, *18*, 340–346. [CrossRef]
- 34. Choudhary, V.R.; Tillu, V.H.; Narkhede, V.S.; Borate, H.B.; Wakharkar, R.D. Microwave assisted solvent-free synthesis of dihydropyrimidinones by Biginelli reaction over Si-MCM-41 supported FeCl₃ catalyst. *Catal Commun.* **2003**, *4*, 449–453. [CrossRef]
- 35. Deksnys, T.P.; Menezes, R.R.; Fagury-Neto, E.; Kiminami, R.H.G.A. Synthesis of Al₂O₃/SiC in microwave oven: Study of the processing parameters. *Cerâmica* **2005**, *51*, 343–348. [CrossRef]
- 36. Deng, S.G.; Lin, Y.S. Microwave heating synthesis of supported sorbents. *Chem. Eng. Sci.* **1997**, *52*, 1563–1575. [CrossRef]
- 37. Roy, R.; Komarneni, S.; Yang, J.L. Controlled microwave heating and melting of gels. *J. Am. Ceram. Soc.* **1985**, 68, 392–395. [CrossRef]
- 38. Korichi, S.; Elias, A.; Mefti, A. Characterization of smectite after acid activation with microwave irradiation. *Appl. Clay Sci.* **2009**, *42*, 432–438. [CrossRef]
- Franco, F.; Pozo, M.; Cecilia, J.A.; Benítez-Guerrero, M.; Pozo, E.; Martín Rubí, J.A. Microwave assisted acid treatment of sepiolite: The role of composition and "crystallinity". *Appl. Clay Sci.* 2014, 102, 15–27. [CrossRef]
- Franco, F.; Pozo, M.; Cecilia, J.A.; Benítez-Guerrero, M.; Lorente, M. Effectiveness of microwave assisted acid treatment on dioctahedral and trioctahedral smectites. The influence of octahedral composition. *Appl. Clay Sci.* 2016, 120, 70–80. [CrossRef]
- Martín de Vidales, J.L.; Pozo, M.; Medina, J.A.; Leguey, S. Formación de sepiolita-paligorskita en litofacies lutítico-carbonáticas en el sector de Borox-Esquivias (Cuenca de Madrid). *Estud. Geol.* 1988, 44, 7–18. [CrossRef]
- 42. Santiago Buey, C.; Suárez Barrios, M.; García Romero, E.; Doval Montoya, M. Mg-Rich Smectite "Precursor" Phase in the Tagus Basin, Spain. *Clays Clay Miner.* **2000**, *48*, 366–373. [CrossRef]
- 43. Pozo, M.; Calvo, J.P. Madrid Basin (Spain): A natural lab for the formation and evolution of magnesian clay minerals. In *Magnesian Clays: Characterization Origin and Applications;* Pozo, M., Galán, E., Theng, B.K.G., Lagaly, G., Eds.; AIPEA Pub.: Rio de Janeiro, Brazil, 2015.
- 44. Galán, E.; Pozo, M. Palygorskite and sepiolite deposits in continental environments. Description, genetic patterns and sedimentary settings. In *Developments in Palygorskite–Sepiolite Research. A New Outlook on These Nanomaterials. Developments in Clay Science*; Galán, E., Singer, A., Eds.; Elsevier: Amsterdam, The Netherlands, 2011; Volume 3, pp. 125–173.
- 45. Lippens, B.C.; De Boer, J.H. Studies on pore systems in catalysts: V. The *t* method. *J. Catal.* **1965**, *4*, 319–323. [CrossRef]
- 46. Brunauer, S.; Emmett, P.H.; Teller, E. Adsorption of gases in multimolecular layers. *J. Am. Ceram. Soc.* **1938**, 60, 309–319. [CrossRef]

- 47. Galán, E.; Aparicio, P. Methodology for the identification and characterization of magnesian clays. In *Magnesian Clays: Characterization, Origin and Applications*; AIPEA Educational, Series; Pozo, M., Galán, E., Eds.; Digilabs: Bari, Italy, 2015; Volume 2, pp. 63–121.
- 48. Sánchez del Río, M.; García Romero, E.; Martínez Criado, G. Variability in sepiolite: Diffraction studies. *Am. Miner.* **2011**, *96*, 1443–1454. [CrossRef]
- Pozo, M.; Calvo, J.P.; Pozo, E.; Moreno, A. Genetic constraints on crystallinity, thermal behaviour and surface area of sepiolite from the Cerro de los Batallones deposit (Madrid Basin, Spain). *Appl. Clay Sci.* 2014, *91*, 30–45. [CrossRef]
- 50. Van der Marel, H.W.; Beutelspacher, H. *Atlas of Infrared Spectroscopy of Clay Minerals and Their Admixtures*; Elsevier: Amsterdam, The Netherlands, 1976; Volume 396.
- 51. Kloprogge, J.T.; Frost, R.L. Study of the thermal behavior of rectorite by in-situ infrared emission spectroscopy. *N. Jb. Miner. Mh.* **2000**, *4*, 145–157.
- 52. Bisio, C.; Boccaleri, E.; Marchese, L.; Superti, G.B.; Pastore, H.O.; Thommes, M. Understanding physico-chemical properties of saponite synthetic clays. *Microporous Mesoporous Mater.* **2008**, *107*, 90–101. [CrossRef]
- 53. Janin, A.; Maache, M.; Lavalley, J.C.; Joly, J.F.; Raatz, F.; Szydlowski, N. FT-IR study of the silanol groups in dealuminated HY zeolites: Nature of the extraframe debris. *Zeolites* **1991**, *11*, 391–396. [CrossRef]
- 54. Weitkamp, J. Zeolites and catalysis. Solid State Ion. 2000, 131, 175–188. [CrossRef]
- 55. Zviagina, B.B.; McCarty, D.K.; Środoń, J.; Drits, V.A. Interpretation of infrared spectra of dioctahedral smectites in the region of OH-stretching vibrations. *Clays Clay Miner.* **2004**, *52*, 399–410. [CrossRef]
- 56. Madejová, J. FTIR techniques in clay mineral studies. *Vib. Spectrosc.* 2003, *31*, 1–10. [CrossRef]
- 57. Farmer, V.C. The Infrared Spectra of Minerals; Mineralogical Society: London, UK, 1974; p. 539.
- 58. Anderson, T.H.; Weaver, F.W.; Owen, N.L. Anomalies in diffuse reflectance infrared spectra of wood and wood polymers. *J. Mol. Struct.* **1991**, 249, 257–275. [CrossRef]
- Madejová, J.; Bujdák, J.; Janek, M.; Komadel, P. Comparative FT-IR study of structural modifications during acid treatment of dioctahedral smectites and hectorite. *Spectrochim. Acta A Mol. Biomol. Spectrosc.* 1998, 54, 1397–1406. [CrossRef]
- 60. Valentín, J.L.; López-Manchado, M.A.; Rodríguez, A.; Posadas, P.; Ibarra, L. Novel anhydrous unfolded structure by heating of acid pre-treated sepiolite. *Appl. Clay Sci.* **2007**, *36*, 245–255. [CrossRef]
- 61. Frost, R.L.; Locos, O.B.; Ruan, H.; Kloprogge, J.T. Near-infrared and mid-infrared spectroscopic study of sepiolites and palygorskites. *Vib. Spectrosc.* **2001**, *27*, 1–13. [CrossRef]
- 62. Frost, R.L.; Cash, G.A.; Kloprogge, J.T. 'Rocky Mountain leather', sepiolite and attapulgite-an infrared emission spectroscopic study. *Vib. Spectrosc.* **1998**, *16*, 173–184. [CrossRef]



© 2018 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (http://creativecommons.org/licenses/by/4.0/).