

Article

## Quantifying Reaeration Rates in Alpine Streams Using Deliberate Gas Tracer Experiments

Andrew Benson <sup>1</sup>, Matthew Zane <sup>2</sup>, Timothy E. Becker <sup>2</sup>, Ate Visser <sup>3</sup>,  
Stephanie H. Uriostegui <sup>2,3</sup>, Elizabeth DeRubeis <sup>4</sup>, Jean E. Moran <sup>4</sup>, Bradley K. Esser <sup>3</sup>  
and Jordan F. Clark <sup>1,2,\*</sup>

<sup>1</sup> Program of Environmental Studies, University of California, Santa Barbara, CA 93106, USA;  
E-Mail: abvg6691@gmail.com

<sup>2</sup> Department of Earth Science, University of California, Santa Barbara, CA 93106, USA;  
E-Mails: mattzane@msn.com (M.Z.); timothy.e.becker@gmail.com (T.E.B.);  
stephanieuriostegui@umail.ucsb.edu (S.H.U.)

<sup>3</sup> Chemical Science Division, Lawrence Livermore National Laboratory, Livermore, CA 94551,  
USA; E-Mails: visser3@llnl.gov (A.V.); diaz42@llnl.gov (S.H.U.); esser1@llnl.gov (B.K.E.)

<sup>4</sup> Department of Earth and Environmental Sciences, California State University East Bay, Hayward,  
CA 94542, USA; E-Mails: ederubeis75@gmail.com (E.D.); jean.moran@csueastbay.edu (J.E.M.)

\* Author to whom correspondence should be addressed; E-Mail: jfclark@geol.ucsb.edu;  
Tel.: +1-805-893-7838; Fax: +1-805-893-2314.

Received: 20 January 2014; in revised form: 1 April 2014 / Accepted: 9 April 2014 /

Published: 22 April 2014

---

**Abstract:** Gas exchange across the air-water interface is a critical process that maintains adequate dissolved oxygen (DO) in the water column to support life. Oxygen reaeration rates can be accurately measured using deliberate gas tracers, like sulfur hexafluoride (SF<sub>6</sub>) or xenon (Xe). Two continuous release experiments were conducted in different creeks in the Sierra Nevada of California: Sagehen Creek in September, 2009, using SF<sub>6</sub> and Martis Creek in August, 2012, using both SF<sub>6</sub> and Xe. Measuring gas loss along the creek, which was approximated with the one-dimensional advection-dispersion equation, allows for the estimation of the SF<sub>6</sub> or Xe reaeration coefficient ( $K_{SF6}$ ,  $K_{Xe}$ ), which is converted to DO reaeration ( $K_{DO}$  or  $K_2$ ) using Schmidt numbers. Mean  $K_{SF6}$  for upper and lower Sagehen and Martis Creeks were, respectively, 34 day<sup>−1</sup>, 37 day<sup>−1</sup> and 33 day<sup>−1</sup>, with corresponding  $K_{DO}$ s of 61 day<sup>−1</sup>, 66 day<sup>−1</sup> and 47 day<sup>−1</sup>. In Martis Creek,  $K_{Xe}$  was slightly higher (21%) than  $K_{SF6}$ , but the calculated  $K_{DO}$  from SF<sub>6</sub> agreed with the calculated  $K_{DO}$  from Xe within about 15%; this difference may be due to bubble-enhanced gas transfer. Established

empirical equations of  $K_{DO}$  using stream characteristics did a poor job predicting  $K_{DO}$  for both creeks.

**Keywords:** reaeration; field tracer experiment; sulfur hexafluoride; xenon; alpine stream; dissolved oxygen

---

## 1. Introduction

Aquatic life requires adequate levels of dissolved oxygen (DO) for survival. Therefore, DO content is a standard monitoring tool to determine the health of fresh water systems. A major consumer of DO in aquatic systems is the microbial degradation of organic matter. A stream's ability to make up this oxygen deficit and return to its solubility equilibrium with the atmosphere is vital. This is accomplished primarily through gas exchange at the air-water interface, where oxygen is either lost or reabsorbed into the stream, due to the concentration gradient between the atmosphere and water. If levels of demand exceed reaeration into the system, a body of water will struggle to support life and may reach hypoxic conditions; in extreme cases, the surface water will become anoxic.

Reaeration coefficients vary widely due to their dependence on turbulence at the air water interface, which is poorly understood and hard to measure. Previous work has related reaeration coefficients to stream characteristics, such as mean depth, current velocity, stream channel slope and discharge (e.g., [1–5]). A number of equations have been postulated, with no one equation appropriate for every channel. The empirical equations also disagree significantly within the same channel. For instance, the equations respond differently to increasing flow: velocity-depth equations predict decreasing reaeration rates, while energy-dissipation models (utilizing channel slope and velocity) predict increasing rates (e.g., [3]). When applied to wastewater management, an underestimation in the reaeration coefficient would result in overly restrictive regulations. With overestimation, a stream or creek will be less resilient to wastewater discharge than predicted, risking hypoxia and collapse of the aquatic ecosystem.

In addition to DO studies, high quality reaeration coefficients are needed for understanding most biogeochemical cycles within streams and other surface water bodies. Many of the gases involved in these cycles, such as  $\text{CO}_2$ ,  $\text{N}_2\text{O}$  and  $\text{CH}_4$ , are greenhouse gases. They are also needed for interpreting  $^{222}\text{Rn}$  distributions that are commonly used to investigate groundwater-surface water interactions (e.g., [6–8]). For instance, the results of the tracer experiment described here provided an estimate of the loss rate of  $^{222}\text{Rn}$  from Martis Creek, which allowed for the quantification of the groundwater influx when combined with  $^{222}\text{Rn}$  measurements in the stream and an estimate of the  $^{222}\text{Rn}$  concentration in shallow groundwater [8].

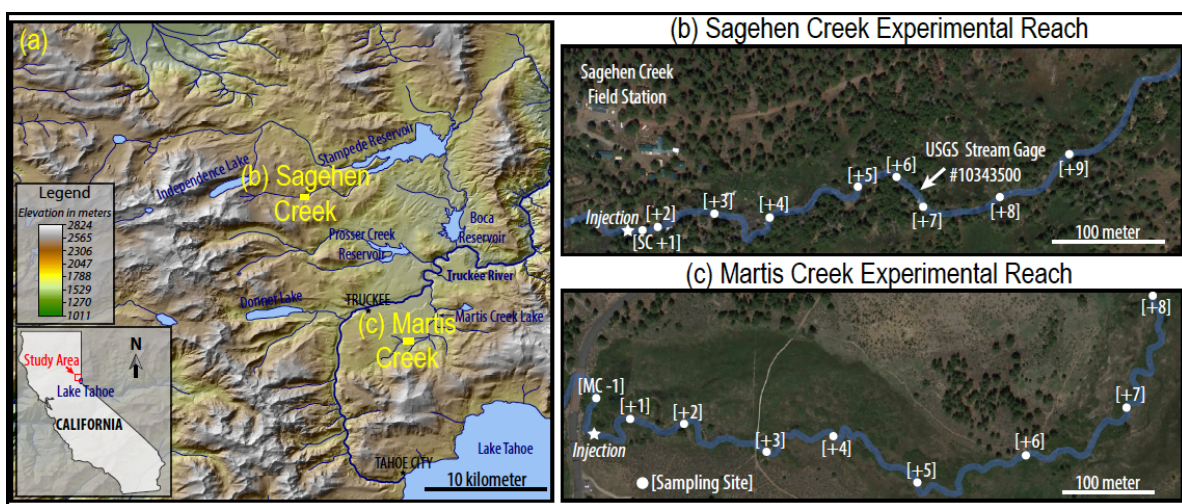
Gas tracers can be used to accurately estimate a stream's reaeration coefficient through point-source injection as either a continuous release (as is the case here) or a single pulse that includes a conservative ion, dye or a second gas tracer [9–12]. Tracer concentrations are measured at specific intervals downstream from injection after the experimental reach has been flushed. Tracer loss from the water to the atmosphere can be used as a proxy for gas exchange across the interface, ultimately allowing for a calculation of reaeration rates.

Two trace gases, sulfur hexafluoride ( $\text{SF}_6$ ) and the noble gas, xenon ( $\text{Xe}$ ), were selected for the tracer experiments conducted in two headwater creeks in the Sierra Nevada of California.  $\text{SF}_6$  was selected, because it has been used in numerous earlier experiments [9–11,13]. As discussed below, it is a very strong greenhouse gas and is currently regulated in California, so during the second experiment, a second gas tracer was tested to determine if it could be used as an alternative tracer for these experiments. The additional objectives of this study are to compare the reaeration rates between the two creeks and with empirical relationships, as others have done [1–5].

### 1.1. Study Area

The first experiment used only  $\text{SF}_6$  and was conducted in Sagehen Creek, a headwater catchment to the Truckee River about 35 km northwest of Lake Tahoe. The second included both  $\text{SF}_6$  and  $\text{Xe}$  and was conducted in another tributary to the Truckee River, Martis Creek, about 15 km southeast of Sagehen near the town of Truckee, CA, USA (Figure 1a). The Sagehen experiment occurred along an approximately 500-m reach of the creek in 2009, adjacent to the Sagehen Creek Field Station (SCFS). The Martis experiment occurred along an approximately 1000-m reach in 2012 (prior to the implementation of the  $\text{SF}_6$  regulations that began 1 January 2013). SCFS is part of the University of California (UC) Natural Reserve system and is managed by UC Berkeley.

**Figure 1.** (a) The study area is located northwest of Lake Tahoe within the Sierra Nevada. Main base map from the United States Geologic Survey (USGS) 10-m National Elevation Dataset [14]. Experimental stream reaches of (b) Sagehen and (c) Martis Creeks.



Both creeks are on the eastern side of the Sierra Nevada at elevations above 1800 m and flow through glacial till deposits (derived from andesite and granodiorite basement rocks) to the Truckee River. They are shallow (mean reach depths of *ca.* 10 cm), meandering streams with riffle and pool morphology and receive discharge perennially from shallow aquifers. At the time of the tracer experiments, the mean water temperatures were 8.3 °C (Sagehen) and 15.8 °C (Martis), though diurnal temperature fluctuations were observed. Both watersheds receive more than 80 cm of precipitation per year, most of which falls as snow between October and April. Peak and minimum discharge typically occur, respectively, during the month of May and September (as baseflow).

## 2. Materials and Methods

### 2.1. One-Dimensional Advection-Dispersion Equation

To obtain the reaeration coefficients for the gas tracers ( $\text{SF}_6$  and  $\text{Xe}$ ), dissolved gas transport was approximated using a one-dimensional (1D) advection-dispersion equation assuming first order decay of a continuously released solute in a river; gas exchange is assumed to be a first order chemical reaction. The 1D transport equation for a pollutant or tracer in a stream is written as (e.g., [11,15,16]):

$$\frac{\partial C}{\partial t} + U \frac{\partial C}{\partial x} = E_x \frac{\partial^2 C}{\partial x^2} - KC \quad (1)$$

Where  $C$  is the concentration of the dissolved gas in excess of its solubility equilibrium value (*i.e.*,  $C = C_{ob} - C_{eq}$ , where  $C_{ob}$  and  $C_{eq}$  are, respectively, the observed gas concentration and the solubility equilibrium value);  $t$  is time;  $U$  is the mean stream velocity;  $x$  is the distance downstream;  $E_x$  is the longitudinal dispersion coefficient; and  $K$  is the reaeration coefficient.

The concentration gradient ( $\partial C/\partial x$ ) is small, because the tracer is being continuously released, and thus, the flux due to dispersion is negligible when compared to the flux due to advection (e.g., [15,16]). Once the stream reach has been flushed with tracer, the stream is assumed to be at a steady state (*i.e.*,  $\partial C/\partial t = 0$ ), and samples can be collected. With these conditions, the advection-dispersion equation can be simplified to:

$$U \frac{dC}{dx} = -KC \quad (2)$$

Solving this differential equation gives the relationship between concentration and sampling distance,  $x$ , from injection:

$$C(x) = C_0 \exp\left[-\frac{K}{U}x\right] \quad (3)$$

where  $C_0$  is the initial concentration of gas tracer at the injection point ( $x = 0$ ).

Injecting the  $\text{SF}_6$  and  $\text{Xe}$  into the stream introduces a problem when calculating the appropriate initial concentration ( $C_0$ ), because the injection rate is not known perfectly, especially when injecting by bubbling, as was done during the 2009 Sagehen Creek experiment.  $C_0$  is obtained from the downstream data. Fortunately,  $C_0$  is not needed to determine the reaeration coefficient.

To be useful for other applications and comparisons to earlier works, the reaeration coefficient for the gas tracers must be converted to analogous values for other gases, such as  $\text{DO}$ ,  $\text{CO}_2$ ,  $\text{Rn}$  and  $\text{N}_2\text{O}$ . This is done using Schmidt numbers ( $Sc_i$ , dimensionless ratios of the kinematic viscosity of water and the diffusion coefficient of the gas in question). The reaeration rate of  $\text{SF}_6$  ( $K_{\text{SF}_6}$ ) or  $\text{Xe}$  ( $K_{\text{Xe}}$ ) is related to other gases ( $K_i$ ) by [1,9,17]:

$$K_i = \left(\frac{Sc_i}{Sc_j}\right)^{-0.5} K_j \quad (4)$$

where the  $Sc_i$  for each gas is defined at a given temperature usually using temperature relations developed by Wanninkhof *et al.* [18] and Raymond *et al.* [1] and  $j$  refers to the gas tracer.

The method used to estimate the reaeration coefficient here is significantly different from most methods reported in the literature, but similar to that used by Cook *et al.* [6]. This work uses a

continuous gas tracer source of SF<sub>6</sub> and allows the concentration distribution to come to equilibrium. A more common approach is to use instantaneous point sources of a gas tracer paired with either a solute tracer or a second gas (e.g., [9–12]). The main advantages of using a continuous source are that the solution to the differential equation (1) is much simpler than the solution for a point source, only one gas tracer is needed and fewer samples need to be analyzed. Having a concentration in terms of distance downstream makes it easy to calculate the reaeration coefficient graphically. The slope of the line created when the log of the concentration is plotted *versus* distance is equal to  $-K/U$ , and the y-intercept is  $C_0$ . Most works present reaeration values normalized to a Schmidt number of 600, which is similar to CO<sub>2</sub> at 20 °C in seawater or oxygen at 17.5 °C ( $K_{DO}$  or  $K_2$ ) in freshwater; we will also use this convention.

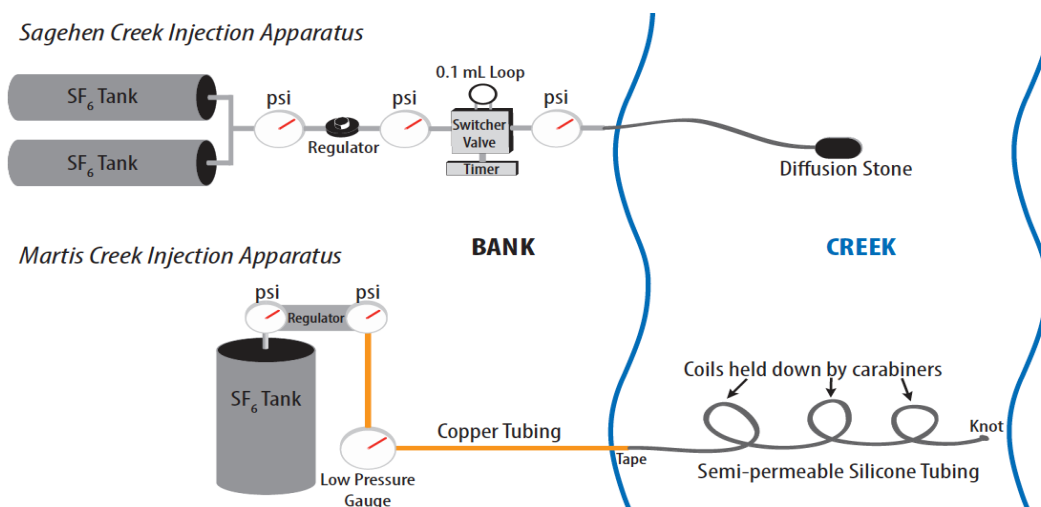
## 2.2. Tracer Injection

SF<sub>6</sub> and Xe were utilized as the deliberate gas tracer in these experiments for a number of reasons. SF<sub>6</sub> is: (1) non-reactive; (2) non-toxic [19]; (3) man-made, so the background concentrations are negligible, especially away from urban areas; (4) detectable at very low levels using gas chromatography, so little gas needs to be released; and (5) relatively inexpensive. However, SF<sub>6</sub> emissions are being regulated in California, because it is a strong greenhouse gas (~24,000 times stronger than CO<sub>2</sub> on a per molecule basis over a 100-y period [20]). Therefore, a second gas, Xe, was released during the Martis Creek experiment to evaluate its potential as a deliberate tracer for gas exchange. Xe, being a noble gas, is also non-reactive and non-toxic. The natural atmospheric concentration is low (87 ppt by volume) and well established. Xe can be measured at atmospheric abundance levels using membrane inlet mass spectrometry [21], so little gas needs to be released. While the cost per liter is similar for the two gases, Xe is not a greenhouse gas and, therefore, not regulated.

Different methods were used to inject the gas tracers during the two experiments. During the September 2009, Sagehen experiment, an injection system (Figure 2) was built and placed next to the stream. Approximately a 1:10 SF<sub>6</sub>:N<sub>2</sub> mixture was created (so that the concentration would be within the calibration range of the analytical method) and released semi-continuously using a switcher valve set to switch every 30 s and bubble into the stream through a diffusion stone placed at the bottom of a small pool near mid-channel (see [13] for the details of the injection device). Injection began the evening before the collection of the first samples and continued until after the collection of the final set. The injection rate of the SF<sub>6</sub> mixture was 0.34 mL/min for the first three sets and then increased to 0.68 mL/min for the final set.

Based on 100% dissolution of the bubbles and the injection rate, the initial concentrations of SF<sub>6</sub> should have been  $4.3 \times 10^6$  pmol/L for the Sagehen Creek experiment. Because the gas was being bubbled into the stream, most of it was lost to the atmosphere and only a small amount (less than 0.01%) of it actually dissolved into the stream. In the Martis Creek Experiment, tracer introduction via gas permeable silicon tubing achieved a gas introduction efficiency of 100%.

**Figure 2.** The diagram of the sulfur hexafluoride ( $\text{SF}_6$ ) injectors used during the Sagehen Creek (top) and Martis Creek (bottom) experiments.



The Martis Creek injection began at 14:15 on 14 August 2012. Injection occurred through 1/8" inner diameter gas permeable silicon tubing fully submerged in the stream (Figure 2) following methods similar to those of Cook *et al.* [6] and Clark [22]. The tubing was coiled and weighted down with carabiners across the creek. The end of the tubing was knotted tightly, so that no large gas bubbles could escape (and none were observed). Copper tubing was attached to the submerged silicon tubing and fastened to the gas tracer tank on the creek bank. Separate setups of silicone tubing were used for the  $\text{SF}_6$  and Xe tracers with 6 m of tubing submerged for  $\text{SF}_6$  and 11 m for Xe. The Xe pressure in the silicone tubing was set to about 2 psi using a pressure regulator. A pressure gauge in line with the copper tubing recorded a constant pressure of 15 psi over the course of the 3-day injection in the  $\text{N}_2$ - $\text{SF}_6$  line. Once again, a 1:10 diluted  $\text{SF}_6$  mixture was used during this experiment.

### 2.3. Creek Sampling

For both  $\text{SF}_6$  experiments, eight or nine sampling locations were selected every 15–100 m downstream of the injection apparatus (Figure 1). At each location, five samples were collected in pre-weighed 10 mL Vacutainers™ evenly distributed across the channel 5–10 cm below the air-water interface. The samples were collected by submerging the Vacutainers™, piercing the septa with a needle, until 2–5 mL of stream water were collected. During the Sagehen experiment, samples were collected during the morning, afternoon and evening of 11 September 2009. A fourth set was collected the following morning on 12 September 2009, after the injection rate was changed.

The Xe samples were collected at the same locations as the  $\text{SF}_6$  samples during the Martis Creek experiment. Duplicate samples were collected from the center of the stream by submersing an empty 40-mL glass vial upside down and filling it at 5–10 cm below the surface. The vial was capped under water. At two downstream locations (Martis Creek (MC) + 1 and MC + 8), additional cross-section samples were collected (in duplicate) at the left and right bank and in between the banks and the center.

During the Martis Creek experiment, sampling began about 1 h after the injection started. It was interrupted by a 2-h rain event that dropped considerable water onto the field site. While this precipitation event may have changed the creek flow, stream gauging shows that the discharge was the same on

Days 2 and 3, so if any changes occurred, they were short lived. The first full set was collected 18 h after injection, with six complete sets collected over the next three days. Four sets were collected on 15 August 2012 (morning, afternoon with duplicates and evening), and one set each was collected during the mornings of 16 August 2012, and 17 August 2012. Xe samples were collected on the morning of 17 August 2012. Background sampling occurred at one location upstream of the injector in Martis Creek. This paper will discuss only the 17 August 2012, data, when both gases were analyzed.

#### 2.4. Creek Measurements

To quantify channel geometry in Sagehen Creek, stream transect data were collected at 27 locations along the creek starting approximately 41 m upstream of the injection point and continuing approximately 550 m downstream. At each transect, the stream width and the depth every 20 cm (from the north bank to the south bank) were measured. In addition, the grain size of the streambed was descriptively characterized at each depth determination (for details, see [23]). A United States Geological Survey (USGS) river gauging station located about 370 m downstream from the injection point provided continuous 15-min discharge data, which showed little variation and averaged 43 L/s during the experiment. The gauge also represents the place where the mean channel geometry changes. In the upper reach, the average measured depth, cross-sectional area and flow velocity were, respectively, 0.11 m,  $0.41 \text{ m}^2$  and 0.11 m/s, while below the gauge, they were, respectively, 0.15 m,  $0.49 \text{ m}^2$  and 0.09 m/s.

The flow velocity in Martis Creek was determined by gauging with an FP111 Global Water™ flow probe at four stations (MC – 1, MC + 2, MC + 5, and MC + 8). Cross-sectional data were also measured at these stations. The average depth, cross-sectional area, flow velocity and discharge were, respectively, 0.11 m,  $0.20 \text{ m}^2$ , 0.29 m/s and 57 L/s. Flow velocity varied from a low average of 0.14 m/s at MC + 2 to a high average of 0.65 m/s at MC + 5; the average velocity at a given cross-sectional location did not vary significantly over the course of the experiment.

The channel gradient in Sagehen Creek was measured using a combination of GPS coordinates and a 1-m digital elevation model (DEM) provided by the Sagehen Creek Field Station that was created from aerial LiDAR flown in September 2005. The gradient was then calculated by dividing the measured center channel length by the difference in elevation based on the DEM and was found to be 0.0143; there was little change in slope between the upper and lower reaches. The channel slope for Martis Creek was determined in a similar fashion, but using Google Earth rather than ArcGIS and LiDAR elevations. It was determined to be 0.0145, very similar to Sagehen Creek, and it also did not vary substantially along the creek.

At Martis Creek, two deep pools along the experimental reach were sampled at three depths (top, middle and bottom) to assess vertical mixing. Little stratification was suspected, due to the stream's generally shallow depth. One pool was ~3 m downstream of MC + 4 with a maximum depth of 0.56 m; the other was ~1 m upstream of MC + 7 with a maximum depth of 0.66 m. Temperature and oxygen measurements were also made using a YSI™ 556 multi-meter (Yellow Springs, OH, USA) and averaged 16.3 °C and 15.4 °C for the two pools. This temperature difference was likely due to the typical diurnal fluctuation found in most shallow creeks rather than an indication of the groundwater discharge of cooler water.



### 2.5. Laboratory Analysis

SF<sub>6</sub> samples were processed using the modified headspace method of Clark *et al.* [24], which employs Vacutainers™. In order to minimize contamination, the containers were kept separate from the injection equipment, during the transport to, from and while at the field sites. The laboratory procedure was as follows: (1) each sample was weighed in order to calculate the volume of water collected; (2) the headspace of each vial was filled with ultra-high purity N<sub>2</sub> to atmospheric pressure in the lab (about 1 atm); (3) vials were agitated by gently shaking in order to allow the gasses in the headspace to become well-mixed; and (4) gases in the headspace were then fed into a gas chromatograph system via displacement with water. SF<sub>6</sub> standards of known concentrations (1.95 parts per billion by volume (ppbv) and 10.0 ppbv) purchased from and certified by Scott-Marrin, Inc. (Riverside, CA, USA) were run every ~10 samples to ensure accurate instrument calibration. Ultra-high purity nitrogen was used to flush the system of residual SF<sub>6</sub> between runs. The analytical uncertainty of the Vacutainer™ method is typically better than ±5%.

Xe samples were analyzed three days after collection on 20 August 2012, using a noble gas membrane inlet mass spectrometer (NG-MIMS) [21]. The NG-MIMS system consists of a membrane inlet, a dry ice water trap, a liquid nitrogen carbon-dioxide trap, two getters, a gate valve, a turbomolecular pump and a quadrupole mass spectrometer equipped with an electron multiplier. Vials were opened, and the sample was withdrawn from the bottom of the vial through the membrane inlet for 5 min at 0.5 mL/min. The last two minutes of measurements were averaged to reduce measurement noise. Dissolved Xe concentrations are determined from measurements made every 10 seconds at mass over charge ratios of 124 and 132, by comparing against air equilibrated water (AEW) standards, with a laboratory determined uncertainty of ±8%. The linearity of the NG-MIMS was demonstrated across a wide range of noble gas concentrations [21].

### 3. Results and Discussion

As expected, average SF<sub>6</sub> and Xe concentrations at each sampling site decreased downstream (Table 1, Figures 3 and 4). The standard deviation is used as a measure of horizontal mixing. When the deviation of bank-to-bank concentrations at each sampling location stabilizes, the tracer is assumed to be well-mixed across the creek. For all sets collected during the Sagehen Creek experiment, the standard deviation of the first sample location (SC + 1) ranges from 20% to 28% of the cross-sectional mean, while, with a few exceptions, the standard deviation for the rest of the samples is generally less than 5% of the mean, equivalent to analytical uncertainty. The first sample location is close to the injector (15 m downstream), and SF<sub>6</sub> had yet to become well mixed across the stream. By the second sampling location (44 m downstream), it had. The first samples were collected further downstream during the Martis Creek Experiment (69 m vs. 15 m) and showed a lower standard deviation (<15%) across the creek. The standard deviation of Xe concentrations in the first (MC + 1) and last (MC + 8) cross-sections was 14% and 10%, respectively. The duplicate reproducibility was 9% for the Xe samples.

In Sagehen Creek, at the location immediately upstream (SC + 6) and downstream (SC + 7) of the USGS stream gauge (about 370 m downstream injector), the drop in concentrations of SF<sub>6</sub> is quite large over a very short distance. This drop indicates a higher reaeration rate, due to the waterfall created by the gauge, as is commonly observed at spillways [25].



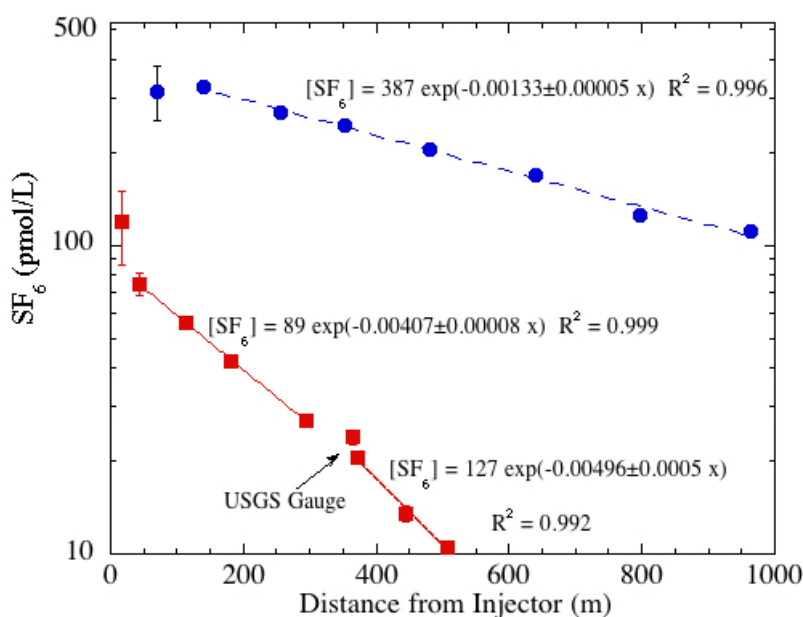
Furthermore, in the large pool immediately upstream of the gauge, the concentration suddenly increased for samples from Sets 1–3. After further inspection, the increases in concentration were coincident with a method change that required the use of different standards. It is important to note that after the offset in standards was found, the fourth set was analyzed using the same method (high standard) from the top to the bottom of the reach. In this paper, only the fourth set is discussed.

**Table 1.** Average tracer concentrations and standard deviations (SD) at each cross-section during the 12 September 2009, Sagehen Creek (SC) and 17 August 2012, Martis Creek (MC) experiments.

Station	Distance (m)	[SF <sub>6</sub> ] (pmol/L)	SD (%)	Station	Distance (m)	[SF <sub>6</sub> ] (pmol/L)	SD (%)	[Xe] (nmol/L)	SD (%)
SC + 1	16	118	27	MC − 1		0	0.0	0.46	-
SC + 2	44	74.9	7.8	MC + 1	69	315	20	43.5	14
SC + 3	114	55.9	2.8	MC + 2	139	325	2	40.2	-
SC + 4	180	42.0	4.3	MC + 3	257	270	3	34.1	-
SC + 5	294	27.1	2.8	MC + 4	352	244	2	29.9	-
SC + 6	364	23.9	5.5	MC + 5	481	203	2	24.4	-
SC + 7	372	20.5	4.9	MC + 6	641	169	2	19.9	-
SC + 8	445	13.5	5.6	MC + 7	798	125	3	13.5	-
SC + 9	507	10.5	2.4	MC + 8	964	111	3	10.8	10

The semi-log plot of the concentration for Sagehen Creek shows that the slope ( $-K/U$ ) for the lower reach is generally steeper than that of the upper reach (Figure 3). In fact, it is about 22% larger. Detailed analysis of the channel geometry, as noted above, revealed that the mean depth of the lower portion (0.15 m) was about 35% larger than the upper section (0.11 m).

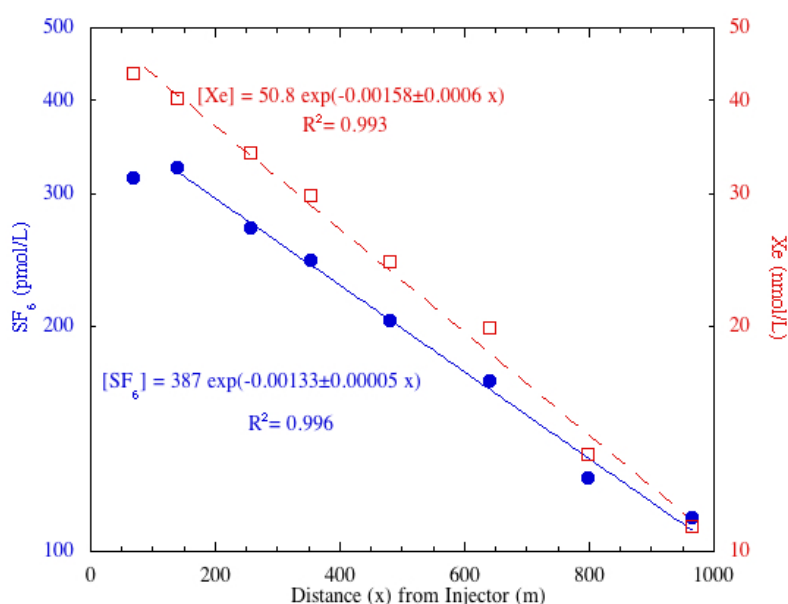
**Figure 3.** Semi-log plot of SF<sub>6</sub> concentrations with least squares fit trend lines showing the different slopes ( $-K/U$ ) for the upper and lower reaches in Sagehen Creek (red squares from 12 September 2009) and for Martis Creek (blue circles from 17 August 2012). In the equations, “x” is the distance downstream.



In Martis Creek, relatively more tracer degassing occurred per meter between stations MC + 6 and MC + 7 than elsewhere along the creek. A few locations of riffles with white water that accelerate gas transfer across the air-water interface were noted between these stations. A comparison of SF<sub>6</sub> and Xe tracers show more similarities than differences (Figure 4). Both gases show reaeration, with SF<sub>6</sub> degassing at a faster rate. This is most apparent when examining the reaeration coefficient normalized for DO (Table 2).

Reaeration rates are calculated from the semi-log plot of gas concentration *versus* sampling distance (Figures 3 and 4). The slope of the exponential best-fit equals  $-K/U$ , where  $K$  is the reaeration rate and  $U$  is mean velocity. Table 2 presents mean stream characteristics and reaeration rates by sampling date and location.

**Figure 4.** Semi-log plot of SF<sub>6</sub> (filled blue circles) and Xe (open red squares) concentrations with trend lines for the Martis Creek 17 August 2012, sampling event. The station closest to the injector was not used for the least squares fit and that only center samples were used for calculating the Xe distribution. In the equations, “x” is the distance downstream.



**Table 2.** Reaeration ( $K$ ) values for the gas tracers ( $K_{SF6 \text{ or } Xe}$ ) and normalized to dissolved oxygen (DO) ( $K_{DO}$ ) using temperature-dependent Schmidt numbers of 1875 and 1192 for SF<sub>6</sub> (Sagehen and Martis Creek), 1230 for Xe (Martis Creek) and 600 for DO (Sagehen and Martis Creek) and Equation (4).

Location	Sampling Date	Tracer	Discharge (L/s)	Velocity $U$ (m/s)	Depth (m)	$-K/U$ (m <sup>-1</sup> )	$K_{SF6 \text{ or } Xe}$ (day <sup>-1</sup> )	$K_{DO}$ (day <sup>-1</sup> )
Sagehen-Upper	12 September 2009	SF <sub>6</sub>	43	0.11	0.11	0.00407	34	61
Sagehen-Lower	12 September 2009	SF <sub>6</sub>	43	0.09	0.15	0.00496	37	66
Martis Creek	17 August 2012	SF <sub>6</sub>	57	0.29	0.11	0.00133	33	47
Martis Creek	17 August 2012	Xe	57	0.29	0.11	0.00156	40	57

For comparison,  $K_{DO}$  values for Sagehen and Martis Creeks were calculated using empirical relationships for the experimental reaches (Table 3, Figure 5). Predictive equations show a very wide

range of values for  $K_{DO}$  (6 to 142  $\text{day}^{-1}$ ) with a mean of  $43 \pm 34 \text{ day}^{-1}$  for Sagehen Creek and (22 to 544  $\text{day}^{-1}$ ) with a mean of  $117 \pm 131 \text{ day}^{-1}$  for Martis Creek. In both creeks, the Thyssen-Jeppesen [26] relationship appears to be an outlier. Removing that relationship from the data set reduces the range and mean to 6–74  $\text{day}^{-1}$  ( $35 \pm 20 \text{ day}^{-1}$ ) and 22–147  $\text{day}^{-1}$  ( $84 \pm 48 \text{ day}^{-1}$ ), respectively, for Sagehen and Martis Creeks.

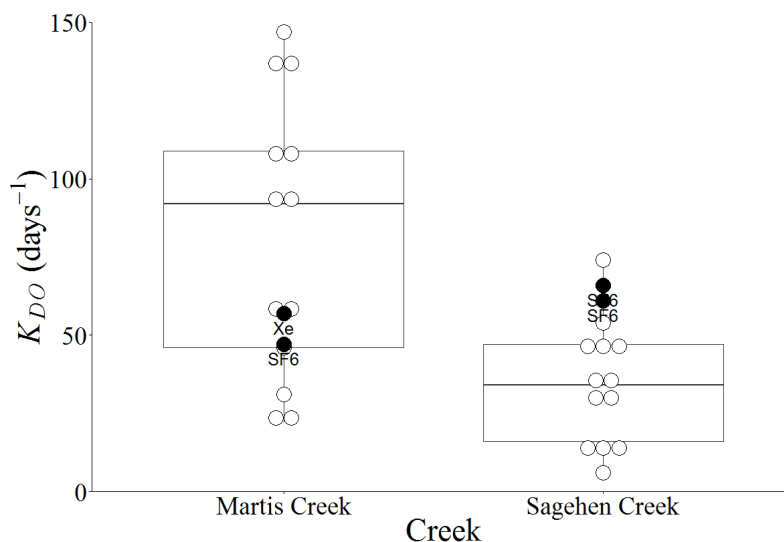
**Table 3.** Predictive equations for reaeration ( $K_{DO}$ ) compiled by Raymond *et al.* [1] and Cox [4]. See these publications for the citations to the original papers. Please note: Raymond *et al.* [1] models calculate gas transfer velocities ( $\text{m day}^{-1}$ ) rather than reaeration coefficients, so the equations listed have been converted, and only the equations that matched their metadata set with a  $R^2 > 0.7$  are presented.

Reference	Predictive Equation	Sagehen Creek $K_{DO} (\text{day}^{-1})$	Martis Creek $K_{DO} (\text{day}^{-1})$
[27]	$K_{DO} = 3.93U^{0.5}H^{-1.5}$	28	58
[28]	$K_{DO} = 5.026UH^{-1.67}$	16	59
[29]	$K_{DO} = 5.35UV^{0.67}H^{-1.85}$	54	139
[30]	$K_{DO} = 5.5773U^{0.607}H^{-1.689}$	47	109
[31]	$K_{DO} = 3170S$	45	46
[32]	$K_{DO} = 22,700SU$	32	95
[33]	$K_{DO} = 8784S^{0.93}U^{0.734}H^{-0.42}$	74	147
[34]	$K_{DO} = 596(SU)^{0.528}Q^{-0.136}$	12	22
[35]	$K_{DO} = 23.04(1 + 0.17F^2)(US^{0.375})H^{-1}(1.0212)^{(T-20)}$	16	27
[36]	$K_{DO} = 23,000(1 + 0.17F)^{2.66}S^{1.13}H^{-0.6}$	142	544
[1] #1	$K_{DO} = 5037(SU)^{0.89}D^{-0.46}$	34	107
[1] #2	$K_{DO} = 5937(1 - 2.54F)(SU)^{0.89}D^{-0.42}$	37	92
[1] #7	$K_{DO} = 4725(SU)^{0.86}Q^{-0.14}D^{-0.34}$	48	135
Mean and Standard Deviation *		$43 \pm 34$	$117 \pm 131$
SF <sub>6</sub> tracer		61, 66 <sup>†</sup>	47
Xe tracer			57

Notes: velocity ( $U$ ,  $\text{m/s}$ ); mean depth ( $H$ ,  $\text{m}$ ); stream gradient ( $S$ ,  $\text{m/m}$ ); Froude number ( $F$ ); Temperature ( $T$ ); \* by removing Thyssen-Jeppesen [36], the mean and standard deviation become  $35 \pm 20$  and  $84 \pm 48$  for Sagehen and Martis Creeks; <sup>†</sup> for upper and lower Sagehen Creek.

Despite the fact that most predictive equations produced wildly differing reaeration rates, underestimation is better than overestimation when considering wastewater discharge to streams. Over predictions could lead to assumptions of a higher level of resiliency and less stringent and potentially dangerous management strategies. Under-predictions, on the other hand, would lead to overly precautionous management, which, while inaccurate, would not threaten stream health.

**Figure 5.** A comparison of the field  $K_{DO}$  (filled circles) listed in Table 2 with the model predicted  $K_{DO}$  (open circles) listed in Table 3, with the exception that the Thyssen-Jeppesen [36] model values are not shown or used in the calculation of median or quartile values.



#### 4. Implications and Conclusions

Sagehen and Martis Creeks are very similar in terms of their morphology (both have pool and riffle characteristics and similar channel slopes) and hydrology (with much of the late summer flow derived from groundwater discharge and a similar mean depth at the time of the experiments). As shown in Table 2, the discharge and current velocity in Martis Creek were higher (respectively, 57 vs. 43 L/s and 0.3 vs. 0.1 m/s). The  $\text{SF}_6$  tracer experiments revealed that the reaeration coefficients corrected to a Schmidt number of 600 ( $K_{DO}$ ) differed substantially, with Sagehen Creek ( $>60 \text{ day}^{-1}$ ) being larger than Martis ( $<50 \text{ day}^{-1}$ ). Interestingly, the results of the Xe tracer experiment ( $57 \text{ day}^{-1}$ ) agree better with the  $\text{SF}_6$  result from Sagehen Creek ( $61\text{--}66 \text{ day}^{-1}$ ) than from Martis Creek ( $47 \text{ day}^{-1}$ ). This difference between the  $\text{SF}_6$  and Xe results could be due to the bubble-enhanced gas transfer [26] that generally removes lower solubility gases faster ( $\text{SF}_6$  is less soluble than Xe).

Most empirical formulas of reaeration rates, expressed as  $K_{DO}$ , rely on an assessment of the current velocity. The actual reaeration rate over a stretch of river or stream is directly observed as a change in tracer concentration. Established empirical formulas of  $K_{DO}$  using stream characteristics did a poor job of predicting  $K_{DO}$  for the two creeks. Although some agree with the field experiments in either Sagehen or Martis Creek, no one relationship worked well for both of them.

Dissolved oxygen is one of the most important indicators of a stream's biologic potential and health. The United States Environmental Protection Agency's criteria for dissolved oxygen are designed to protect freshwater aquatic life under "worst case conditions"; the suggested DO limit expressed as a seven-day mean minimum for cold water is 4.0 mg/L [37]. In this study, DO concentrations in Martis Creek ranged between 5 mg/L and 9 mg/L throughout the summer. Considering that minimum criteria estimations are conservative, DO content in this stream reach is sometimes close to threatening levels. This exemplifies the need for accurate measurements of reaeration to ensure a

healthy aquatic environment. Future expansion of development in the Tahoe and Truckee areas could also increase biological oxygen demand on the creek and lead to noncompliance.

With mounting concerns and subsequent regulations of greenhouse gasses, the future of SF<sub>6</sub> as a hydrologic tracer is in doubt. Noble gasses, such as helium and Xe, are on their way to replacing SF<sub>6</sub>, and we demonstrate that Xe is a viable replacement option for reaeration studies, as shown by others [12]. Future SF<sub>6</sub> use may depend on more efficient injection methods. Traditional bubbling results in an immediate loss of most of the gas to the atmosphere, requiring more SF<sub>6</sub> for water tagging. The semi-permeable silicon tubing, used in the Martis Creek study, is very promising in this regard and results in very high injection efficiency.

## Acknowledgments

We would like to thank Tom Gleeson (McGill University) and Andrew Manning (USGS, Denver, CO, USA) for their help with collecting samples and measuring stream transects. We are grateful to Sagehen Creek Field Station; especially, station manager Jeff Brown, for technical and logistical assistance. Gina Lee helped with some concepts and GIS. We are also grateful for the time and effort of the two anonymous reviewers. Their comments helped to improve the paper substantially. Financial support for this project was provided by the California Energy Commission Public Interest Energy Research (PIER) Project PIR-08-010, the WaterReuse Research Foundation project WRF 09-11, the State of California Groundwater Ambient Monitoring & Assessment (GAMA) Special Studies Program, and the Lawrence Livermore National Laboratory Lawrence (LLNL) Scholar Program. Parts of this project were performed by LLNL under Contract DE-AC52-07NA27344.

## Conflicts of Interest

The authors declare no conflict of interest.

## References

1. Raymond, P.A.; Zappa, C.J.; Butman, D.; Bott, T.L.; Potter, J.; Mulholland, P.; Laursen, A.E.; McDowell, W.H.; Newbold, D. Scaling the gas transfer velocity and hydraulic geometry in streams and small rivers. *Limnol. Oceanogr. Fluids Environ.* **2012**, *2*, 41–53.
2. Moog, D.B.; Jirka, G.H. Analysis of Reaeration Equations Using Mean Multiplicative Error. In: *Air-Water Gas Transfer*; Jähne, B., Monahan, E.C., Eds.; AEON Verlag & Studio: Hanau, Germany, 1995; pp. 101–111.
3. St. John, J.P.; Gallagher, T.W.; Paquin, P.R. The Sensitivity of the Dissolved Oxygen Balance to Predictive Reaeration Equations. In *Gas Transfer at Water Surfaces*; Brutsaert, W., Jirka, G.H., Eds.; D. Reidel Publishing Company: Dordrecht, The Netherlands, 1984; pp. 577–588.
4. Cox, B.A. A review of dissolved oxygen modelling techniques for lowland rivers. *Sci. Total Environ.* **2003**, *314*, 303–334.
5. Haider, H.; Ali, W.; Haydar, S. Evaluation of various relationships of reaeration rate coefficient for modeling dissolved oxygen in a river with extreme flow variations in Pakistan. *Hydrol. Process.* **2013**, *27*, 3949–3963.

6. Cook, P.G.; Lamontagne, S.; Berhane, D.; Clark, J.F. Quantifying groundwater discharge to Cockburn River, southeastern Australia, using dissolved gas tracers  $^{222}\text{Rn}$  and  $\text{SF}_6$ . *Water Resour. Res.* **2006**, *42*, 12.
7. Cartwright, I.; Hofmann, H.; Sirianos, M.A.; Weaver, T.R.; Simmons, C.T. Geochemical and  $^{222}\text{Rn}$  constraints on baseflow to the Murray River, Australia, and timescales for the decay of low-salinity groundwater lenses. *J. Hydrol.* **2011**, *405*, 333–343.
8. DeRubeis, E.A. Radon as a Tracer of Groundwater-Surface Water Interaction in Martis Valley. Master's Thesis, California State University, East Bay, CA, USA, 4 October 2013.
9. Wanninkhof, R.; Mulholland, P.J.; Elwood, J.W. Gas exchange rates for a first-order stream determined with deliberate and natural tracers. *Water Resour. Res.* **1990**, *26*, 1621–1630.
10. Clark, J.F.; Wanninkhof, R.; Schlosser, P.; Simpson, H.J. Gas-exchange rates in the tidal Hudson River using a dual tracer technique. *Tellus Ser. B Chem. Phys. Meteorol.* **1994**, *46*, 274–285.
11. Hibbs, D.E.; Parkhill, K.L.; Gulliver, J.S. Sulfur hexafluoride gas tracer studies in streams. *J. Environ. Eng. Asce* **1998**, *124*, 752–760.
12. Reid, S.E.; Mackinnon, P.A.; Elliot, T. Direct measurements of reaeration rates using noble gas tracers in the River Lagan, Northern Ireland. *Water Environ. J.* **2007**, *21*, 182–191.
13. Clark, J.F.; Hudson, G.B.; Avisar, D. Gas transport below artificial recharge ponds: Insights from dissolved noble gases and a dual gas ( $\text{SF}_6$  and  $^3\text{He}$ ) tracer experiment. *Environ. Sci. Technol.* **2005**, *39*, 3939–3945.
14. Gesch, D.B. The National Elevation Dataset. In *Digital Elevation Model Technologies and Applications: The DEM Users Manual*, 2nd ed.; Maune, D., Ed.; American Society for Photogrammetry and Remote Sensing: Bethesda, MD, USA, 2007; pp. 99–118.
15. Rutherford, J.C. *River Mixing*; Wiley: New York, NY, USA, 1994; p. 347.
16. Thomann, R.V.; Mueller, J.A. *Principles of Surface Water Quality Modeling and Control*; Harper & Row: New York City, NY, USA, 1987; p. 644.
17. Jähne, B.; Heinz, G.; Dietrich, W. Measurement of the diffusion coefficients of sparingly soluble gases in water. *J. Geophys. Res. Oceans* **1987**, *92*, 10767–10776.
18. Wanninkhof, R. Relationship between wind-speed and gas-exchange over the ocean. *J. Geophys. Res. Oceans* **1992**, *97*, 7373–7382.
19. Lester, D.; Greenberg, L.A. The toxicity of sulfur hexafluoride. *Arch. Ind. Hyg. Occup. Med.* **1950**, *2*, 348–349.
20. Houghton, J.T.; Filho, L.G.M.; Bruce, J.P.; Lee, H.; Callander, B.A.; Haites, E.F.; Harris, N.; Maskell, K. *Climate Change 1994: Radiative Forcing of Climate Change and an Evaluation of the IPCC 1992 IS92 Emission Scenarios*; Cambridge University Press: Cambridge, UK, 1995; p. 339.
21. Visser, A.; Singleton, M.J.; Hillegonds, D.J.; Velsko, C.A.; Moran, J.E.; Esser, B.K. A membrane inlet mass spectrometry system for noble gases at natural abundances in gas and water samples. *Rapid Commun. Mass Spectrom.* **2013**, *27*, 2472–2482.
22. Clark, J.F. Defining Transport near ASR Operations Using Sulfur Hexafluoride Gas Tracer Experiments. In *Management of Aquifer Recharge for Sustainability (4th International Symposium on Artificial Recharge of Groundwater)*; Dillon, P.J., Ed.; A.A. Balkema Publishers: Exton, PA, USA, 2002; pp. 257–260.

23. Zane, M. Reaeration of Sagehen Creek near Truckee, CA. Bachelor's Thesis, University of California, Santa Barbara, CA, USA, 11 June 2010.
24. Clark, J.F.; Hudson, G.B.; Davisson, M.L.; Woodside, G.; Herndon, R. Geochemical imaging of flow near an artificial recharge facility, Orange County, California. *Ground Water* **2004**, *42*, 167–174.
25. Rindels, A.J.; Gulliver, J.S. Oxygen Transfer at Spillways. In Proceedings of the Second International Symposium on Gas Transfer at Water Surfaces, Minneapolis, MN, USA, 11–14 September 1990; American Society of Civil Engineers: New York, NY, USA, 2001; pp. 524–533.
26. Ho, D.T.; Asher, W.E.; Bliven, L.F.; Schlosser, P.; Gordan, E.L. On mechanisms of rain-induced air-water gas exchange. *J. Geophys. Res. Oceans* **2000**, *105*, 24045–24057.
27. O'Conner, D.J.; Dobbins, W.E. Mechanisms of reaeration in natural streams. *Trans. Am. Soc. Civil Eng.* **1956**, *123*, 641–667.
28. Churchill, M.A.; Elmore, H.L.; Buckingham, R.A. Prediction of stream reaeration rates. *Int. J. Air Water Pollut.* **1962**, *6*, 467–504.
29. Owens, M.; Edwards, R.; Gibbs, J. Some reaeration studies in streams. *Int. J. Air Water Pollut.* **1964**, *8*, 469–486.
30. Bennett, J.P.; Rathburn, R.E. *Reaeration in Open-Channel Flow*, USGS Professional Paper: No. 737; United States Geological Survey: Washington, DC, USA, 1972.
31. Tsiovoglou, E.C.; Neal, L.A. Tracer measurements of reaeration: III—Predicting the reaeration capacity of inland streams. *J. Water Pollut. Control. Fed.* **1976**, *48*, 2669–2689.
32. Grant, R.S. *Reaeration-Coefficient Measurement of 10 Small Streams in Wisconsin Using Radioactive Tracers with a Section of the Energy Dissipation Model*; U.S. Geological Survey Water-Resources Investigations 76–96; United States Geological Survey: Madison, WI, USA, 1976.
33. Thyssen, N.; Erlandsen, M.; Jeppesen, E.; Ursin, C. Reaeration of oxygen in shallow, macrophyte rich streams: I—Determination of the reaeration rate coefficient. *Int. Revue Ges. Hydrobiol.* **1987**, *72*, 405–429.
34. Melching, C.S.; Flores, H.E. Reaeration equations derived from US Geological Survey database. *J. Environ. Eng.* **1999**, *125*, 407–414.
35. Parkhurst, J.D.; Pomeroy, R.D. Oxygen absorption of streams—State of the art. *J. Sanit. Eng. Div. ASCE* **1972**, *98*, 101–124.
36. Thyssen, N.; Jeppesen, E. Reaeration in small water courses. *Vatten* **1980**, *36*, 231–248.
37. US Environmental Protection Agency. *Quality Criteria for Water 1986: The Gold Book*; US Environmental Protection Agency: Washington, DC, USA, 1987; p. 477.